

1 **Large-Eddy-Simulation Study on the Effects of**
2 **Building-Height Variability on Turbulent Flows over**
3 **an Actual Urban Area**

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7 **Abstract** A large-eddy simulation (LES) was conducted to investigate the ef-
8 fects of building-height variability on turbulent flows over an actual urban area,
9 the city of Kyoto, which was reproduced using a 2-m resolution digital surface
10 dataset. Comparison of the morphological characteristics of Kyoto with those of
11 European, North American, and other Japanese cities indicates a similarity to
12 European cities but with more variable building heights. The performance of the
13 LES model is validated and found to be consistent with turbulence observations
14 obtained from a meteorological tower and Doppler lidar. We conducted the follow-
15 ing two numerical experiments: a control experiment using Kyoto buildings, and
16 a sensitivity experiment in which all the building heights are set to the average
17 height over the computational region h_{all} . The difference of Reynolds stress at

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18 height $z = 2.5h_{all}$ between the control and sensitivity experiments is found to
19 increase with the increase in the plan-area index (λ_p) for $\lambda_p > 0.32$. Thus, values
20 of λ_p of around 0.3 can be regarded as a threshold for distinguishing the effects of
21 building-height variability. The quadrant analysis reveals that sweeps contribute to
22 the increase in the Reynolds stress in the control experiment at height $z = 2.5h_{all}$.
23 The exuberance in the control experiment at height $z = 0.5h_{all}$ is found to de-
24 crease with an increase in the building-height variability. Although the extreme
25 momentum flux at height $z = 2.5h_{all}$ in the control experiment appears around
26 buildings, it contributes little to the total Reynolds stress and is not associated
27 with coherent motion.

28 **Keywords** Actual urban building · Large-eddy simulation · Atmospheric
29 turbulence · Roughness parameter · Reynolds stress · Quadrant analysis

30 1 Introduction

31 Atmospheric processes over urban areas are affected not only by meteorological
32 disturbances, such as thunderstorms, fronts, and cyclones, but also by the rough-
33 ness and thermal effects of buildings and man-made structures. The geometrical
34 features of buildings and structures determine the roughness effects of an urban
35 area, while human activities and the material characteristics of buildings play a
36 role in defining the thermal effects of such areas. The complex geometrical nature
37 of urban surfaces results in highly complex turbulent flows. To properly under-
38 stand the physical processes of momentum and heat transfer in urban areas and
39 develop parametrizations for urban environments in numerical weather prediction

40 models, it is important to reveal relationships between the effects of actual urban
41 buildings and turbulent flows.

42 The characteristics of turbulent flows over urban surfaces have been examined
43 in numerous previous studies. Oke (1988) categorized the airflow over roughness
44 obstacles as a function of obstacle density as isolated flows, wake-interference flows,
45 and skimming flows. Macdonald et al. (1998) derived a theoretical relation for the
46 aerodynamic roughness length z_0 and displacement height d for flows over rough-
47 ness blocks. While these studies examined turbulent flows over roughness blocks
48 with constant height and regular distribution, the recent focus has shifted to the
49 effects on turbulent flows of roughness blocks with variable height and inhom-
50 geneous arrangement. Wind-tunnel experiments conducted by Cheng and Castro
51 (2002) demonstrated that the roughness sublayers over block arrays with random
52 height are thicker than those over uniform-height arrays. Xie et al. (2008) con-
53 ducted a large-eddy simulation (LES) of turbulent flows over block arrays with
54 random height, and found that the tall blocks significantly contribute to the total
55 drag of such arrays. Nakayama et al. (2011) performed LES investigations over
56 building arrays with different height variability and found that the vertical pro-
57 files of the mean velocity and Reynolds stress depend significantly on the building-
58 height variability. Zaki et al. (2011) performed wind-tunnel experiments with block
59 arrays of buildings with variable height distributed randomly, and showed that the
60 drag coefficient C_d increases with the building density and the standard deviation
61 of the building height for high building densities. Numerical simulations of plume
62 dispersion over urban surfaces have revealed that the turbulence is significantly
63 affected by the source location and wind direction because of the strong depen-
64 dence on the building height and distribution. (Xie and Castro 2009; Xie 2011;

65 Nakayama et al. 2016). The parametrizations of z_0 and d have been improved by
66 taking into account roughness parameters associated with actual urban buildings,
67 such as the maximum, standard deviation, and skewness of the building height
68 (Nakayama et al. 2011; Kanda et al. 2013; Zhu et al. 2017). Giometto et al. (2016)
69 suggested that the dispersive flux derived from spatial variations of temporal mean
70 flows around buildings should be considered to improve conventional urban-canopy
71 parametrizations.

72 To fully understand the effects of roughness obstacles on turbulent flows, it
73 is helpful to investigate the relationships between turbulent organized structures
74 and obstacles, because organized structures are associated with downwards mo-
75 mentum transfer in the form of ejection and sweep events based on a quadrant
76 analysis for the turbulent momentum flux. The results of wind-tunnel experiments
77 on flows over rough surfaces conducted by Raupach (1981) indicate that sweeps
78 are dominant for the total momentum flux near surfaces, and that the contribu-
79 tion of ejection to the momentum flux increases with height. Studies in which
80 turbulence was observed over actual urban areas have revealed the characteristics
81 of momentum transfer and coherent motion. Oikawa and Meng (1995) observed
82 turbulent structures associated with ejections and sweeps over an urban area, and
83 found that turbulent structures correlate with heat transfer within and above the
84 urban canopy. Christen et al. (2007) analyzed field experimental data obtained
85 from sonic-anemometer measurements within and above a street canyon in Basel,
86 Switzerland, and found that sweeps are mostly dominant up to a height of approxi-
87 mately twice the average building height in a street canyon. Numerical simulations
88 of flows over building arrays have revealed the spatial characteristics of turbulent
89 organized structures. Kanda et al. (2004) carried out LES investigations of tur-

90 bulent flows over uniform-height block arrays to investigate turbulent organized
91 structures over such arrays. They found low-speed streaks and streamwise vortices
92 similar to those in flows over flat-wall boundary layers. Kanda (2006) indicated
93 that streak structures are a common feature over various types of block arrays.
94 Using direct numerical simulations, Coceal et al. (2007a,b) revealed that hairpin
95 vortices associated with ejections and sweeps are generated over uniform block
96 arrays, and that the low-speed streaks identified above such arrays are composed
97 of large numbers of hairpin vortices aligned in the streamwise direction. Park et al.
98 (2015) used LES results to analyze turbulent-flow structures over an actual urban
99 area in Seoul, Korea, and showed that turbulent structures behind high-rise build-
100 ings are characterized by streamwise vortices with strong ejections. They focused
101 on small regions containing high-rise buildings, and demonstrated the significant
102 influence of high-rise buildings on wake flows. The majority of studies presented
103 thus far have focused on the characteristics of turbulent flows over idealized or
104 specific buildings, while only a few have examined the urban-scale effects on the
105 characteristics of turbulent momentum transfer produced by the complex geomet-
106 rical features of actual urban surfaces.

107 The geometrical characteristics of actual urban surfaces can be reproduced
108 from digital surface datasets. Ratti et al. (2002) calculated the roughness param-
109 eters of North American and European cities, and found that parameters differ
110 significantly by city. Bou-Zeid et al. (2009) indicated that turbulent flows are
111 dependent on the building representation over the actual urban surface. To un-
112 derstand the characteristics of turbulent flows over urban areas, it is therefore
113 important to use the geometry of actual buildings in simulations and experiments.

114 We investigate here the effects of building-height variability in an actual urban
115 area on turbulent flows at an urban scale, focusing on the airflow within and above
116 an urban-canopy layer, where turbulent flows are strongly influenced by individual
117 buildings.

118 We simulate the turbulent flow over the urban area of Kyoto, which is charac-
119 terized by the presence of both business districts with high buildings and densely
120 built residential districts. Furthermore, a meteorological observation tower owned
121 by Kyoto University and located in the southern part of the city can be used for
122 the validation of simulations. In Sect. 2, the building morphological characteris-
123 tics of Kyoto are evaluated using roughness parameters. The details of our LES
124 model are described in Sect. 3. The study area of the LES investigation is defined
125 to include the meteorological tower site at which turbulence was measured by a
126 sonic anemometer and Doppler lidar, so that LES results may be compared with
127 the observations (see Sect. 4). Along with a control simulation, we conduct a sen-
128 sitivity test assuming a constant building height to reveal the effects of building
129 height–height variability, with the differences between the control and sensitivity
130 experiments examined in Sect. 5. Finally, Sect. 6 concludes the paper.

131 **2 Building Morphological Characteristics of Kyoto**

132 Our study area covered both business districts and suburban areas in Kyoto. Fig-
133 ure 1 shows the area of interest in Kyoto, which extends 11 km in a north–south
134 direction and by 2 km in an east–west direction. A digital surface model (Koku-
135 sai Kogyo Co., Ltd.) was used to reproduce the actual urban buildings within a
136 numerical model. The original 2-m-resolution data are smoothed and converted

137 to a 4-m resolution, which is used as the horizontal grid spacing of the numerical
138 experiments as described in Sect. 3.2.

139 Figure 2a shows the height of the actual buildings in the analysis area. The
140 north–south and west–east directions are referred to as the x and y directions,
141 respectively. The region with $x = 0 - 4$ km corresponds to the city centre of
142 Kyoto. The heights of almost all buildings in the region are up to 50 m, and there
143 are no high-rise building clusters of the type seen in the centre of Tokyo. The
144 region for $x = 7 - 11$ km is primarily occupied by suburban areas and rivers.

145 The difference between the building heights over these two regions is clearly
146 indicated in Fig. 2b, which shows the frequency distributions of building heights
147 over the entire analysis area and in the $x = 0 - 4$ km and $x = 7 - 11$ km regions. In
148 calculating the frequency distributions, all buildings are defined as having heights
149 of at least 1 m to distinguish between the buildings and the ground. It is seen that
150 most of the buildings taller than 25 m are located in the former region.

151 To quantitatively indicate the morphological characteristics of buildings in Ky-
152 oto, we use roughness parameters such as the average building height H_{ave} , the
153 standard deviation of the building height σ_H , the plan-area index λ_p (the ratio of
154 the plan area occupied by buildings to the total surface area), and the frontal-area
155 index λ_f (the ratio of the frontal area of buildings to the total surface area). These
156 parameters are calculated for each 1 km by 1 km area following the analysis of
157 Kanda et al. (2013). Figure 3a shows λ_p calculated in the areas of 1 km by 1 km
158 for the buildings shown in Fig. 2a, with the values of the roughness parameters in
159 the 1 km by 1 km areas summarized in Fig. 3b. The average values of H_{ave} , σ_H ,
160 λ_p , and λ_f over the $x = 0 - 4$ km region are 10.8, 7, 0.41, and 0.25, respectively,
161 while the corresponding averages over $x = 7 - 11$ km are 9.8, 5.3, 0.2, and 0.16,

162 respectively. Thus, the $x = 0 - 4$ km region is more densely built than the $x = 7 -$
 163 11 km. Using building data from Tokyo and Nagoya, Japan, Kanda et al. (2013)
 164 derived the following empirical relationships between λ_p and λ_f , and between H_{ave}
 165 and σ_H ,

$$\lambda_f = 1.42\lambda_p^2 + 0.4\lambda_p, \quad (1)$$

$$\sigma_H = 1.05H_{ave} - 3.7. \quad (2)$$

166 Figure 3c and d indicates the respective relationships between λ_p and λ_f , and
 167 between H_{ave} and σ_H , based on the data given in Fig. 3b. Also shown in the
 168 panels are the empirical relationships of Kanda et al. (2013) and the data for North
 169 American and European cities found in Ratti et al. (2002). For $\lambda_p > 0.3$, the λ_f
 170 values for Kyoto tend to be smaller than in the empirical profile. This feature
 171 of Kyoto appears to be similar to those seen in European cities, and indicates
 172 that the fraction of high buildings in Kyoto is limited relative to those in major
 173 metropolitan cities in Japan and North America. The relationship between H_{ave}
 174 and σ_H for Kyoto is in good agreement with those of Tokyo and Nagoya, but
 175 differs from those of European cities. Finally, the magnitudes of H_{ave} and σ_H in
 176 Kyoto are smaller than those of Los Angeles by a factor of 5 – 10.

177 According to these results, Kyoto can be morphologically characterized as hav-
 178 ing densely distributed buildings with widely varying heights. The Kyoto dataset
 179 was used for the numerical simulations described in the next section.

3 Numerical Model and Experimental Design

3.1 Numerical Model

Our LES model is effectively the same as the one used in Nakayama et al. (2011), except that it neglects the molecular viscosity term, and employs a bottom boundary condition based on Monin–Obukhov similarity theory, as described later. In Nakayama et al. (2011), the performance of the LES model reproducing turbulent statistics was validated using data obtained from wind-tunnel experiments; as a close agreement was found, the model developed by Nakayama et al. (2011) has subsequently been applied to simulate turbulent flows over actual urban cities. Nakayama et al. (2012) conducted LES investigations of turbulent flows over Tokyo by coupling their model with a mesoscale meteorological model, and found that observed gust factors are accurately reproduced by the model. The model was also used to successfully reproduce the wind speeds and directions at the ground level in the Fukushima Daiichi Nuclear Power Plant during the Great East Japan Earthquake and its aftermath in March 2011 (Nakayama et al., 2015). Nakayama et al. (2016) further applied their LES model for the simulation of turbulent flows and plume dispersion over Oklahoma City, and showed that the observed characteristics of turbulence and dispersion are reproduced despite the fact that small differences in wind direction caused by the building distribution significantly influenced the plume dispersion. Thus, our LES model has been widely tested and is applicable for the analysis of the turbulent flow over Kyoto.

The LES model solves the filtered continuity and Navier–Stokes equations in Cartesian coordinates with the subgrid-scale stress parametrized by the standard Smagorinsky model (Smagorinsky, 1963). The governing equations are

$$\frac{\partial \tilde{u}_i}{\partial x_i} = 0, \quad (3)$$

$$\frac{\partial \tilde{u}_i}{\partial t} + \tilde{u}_j \frac{\partial \tilde{u}_i}{\partial x_j} = -\frac{1}{\rho} \frac{\partial \tilde{p}^*}{\partial x_i} - \frac{\partial \tau_{ij}}{\partial x_j} + f_i, \quad (4)$$

$$\tau_{ij} - \frac{1}{3} \delta_{ij} \tau_{kk} = -2(C_s \Delta)^2 (2\tilde{S}_{ij} \tilde{S}_{ij})^{1/2} \tilde{S}_{ij}, \quad (5)$$

$$\tilde{S}_{ij} = \frac{1}{2} \left(\frac{\partial \tilde{u}_i}{\partial x_j} + \frac{\partial \tilde{u}_j}{\partial x_i} \right), \quad (6)$$

204 where t denotes time, \tilde{u}_i is the filtered air velocity in the direction i , $\tilde{p}^* =$
 205 $\tilde{p} + \frac{1}{3} \rho \tau_{kk}$ is the modified pressure, \tilde{p} is the filtered pressure, ρ is the density of air,
 206 τ_{ij} is the subgrid-scale stress, δ_{ij} is the Kronecker delta, \tilde{S}_{ij} is the filtered stress
 207 tensor, and f_i is the external force exerted by roughness obstacles. The parameter
 208 x_i represents the coordinate system, with components $i = 1, 2,$ and 3 referring
 209 to the streamwise (x), spanwise (y) and vertical (z) directions, respectively. In
 210 addition, $\Delta = (\Delta_x \Delta_y \Delta_z)^{1/3}$ is the filter width, where $\Delta_x, \Delta_y,$ and Δ_z are the
 211 streamwise, spanwise, and vertical grid spacings, respectively. The Smagorinsky
 212 coefficient C_s is set to 0.14. Note that the viscous term is neglected because our
 213 target is the simulation of turbulent flow with a high Reynolds number.

214 The external force f_i is used to simulate the effects of buildings on the flow,
 215 for which we employ the feedback forcing by Goldstein et al. (1993) who give

$$f_i = \alpha \int_0^t u_i(t') dt' + \beta u_i(t), \quad \alpha < 0, \beta < 0, \quad (7)$$

216 where α and β are negative constants. The stability limit is given by $\Delta t <$
 217 $\frac{-\beta - \sqrt{(\beta^2 - 2\alpha k)}}{\alpha}$, where k is a constant of order one. Following Nakayama et al.
 218 (2011), these constants are set as $\alpha = -10,$ $\beta = -1,$ and $k = 1.$

219 The governing equations are discretized on a staggered-grid system. The veloc-
 220 ity and pressure fields are solved using a coupling method based on the marker-and-

221 cell method (Chorin, 1967). The successive over-relaxation method is used to solve
222 the Poisson equation for pressure, and the Adams–Bashforth scheme is adopted for
223 the time integration. A second-order accurate, central-differencing scheme is em-
224 ployed for spatial discretization. The code is parallelized using a Message Passing
225 Interface library to reduce the computational time.

226 3.2 Experimental Design

227 The governing equations are numerically solved in two computational domains:
228 the driver region, which features regularly arrayed obstacles, and the main region,
229 which contains the actual buildings of Kyoto. To ensure the flow field of the LES
230 is turbulent, a turbulent flow is generated in the driver region and imposed as
231 the inflow at the boundary of the main region. The concept involved in setting
232 the driver and the main regions is demonstrated in Fig. 4. The size of the driver
233 region is 6 km (streamwise) \times 2.4 km (spanwise) \times 1.015 km (vertical), with a
234 grid spacing of 4 m in the horizontal directions, and a grid spacing stretched with
235 increasing altitude from 1 m to 16 m in the vertical direction. The total number
236 of grid points is $1500 \times 600 \times 105$. In the driver region, there is one rectangular
237 block aligned in the spanwise direction, and an array of roughness blocks staggered
238 with $\lambda_p = 0.04$. The individual rectangular and roughness block sizes are $50 \text{ m} \times$
239 $2400 \text{ m} \times 50 \text{ m}$ and $16 \text{ m} \times 16 \text{ m} \times 10 \text{ m}$, respectively. The purpose of setting
240 the size of the rectangular block is to enhance perturbations near the inlet of the
241 driver region. The λ_p value chosen for the block array is set to be a little larger
242 than that in Nakayama et al. (2014) to reduce the generation of turbulence. The

243 height of the blocks is chosen according to the mean building height in the main
 244 region.

245 A uniform flow with a velocity magnitude of 5 m s^{-1} is imposed at the inflow
 246 boundary of the driver region. The Sommerfeld radiation condition is imposed at
 247 the outflow boundary, while a periodic condition is set at the lateral boundaries.
 248 At the top boundary, free-slip and zero-speed conditions are imposed for the hor-
 249 izontal and vertical velocity components, respectively. At the ground, a boundary
 250 condition based on Monin–Obukhov similarity theory is employed. The stress at
 251 the first vertical grid $\tau_{i3}(x, y, t)$ is calculated as (Stoll and Porté-Agel, 2006)

$$\tau_{i3}(x, y, t) = - \left[\frac{\tilde{u}_r(x, y, z_s, t)\kappa}{\ln(z_s/z_0)} \right]^2 \frac{\tilde{u}_i(x, y, z_s, t)}{\tilde{u}_r(x, y, z_s, t)}, \quad (8)$$

252 where $\tilde{u}_r(x, y, z_s, t) = [\tilde{u}_1(x, y, z_s, t)^2 + \tilde{u}_2(x, y, z_s, t)^2]^{1/2}$ is the instantaneous re-
 253 solved velocity magnitude, z_s is the altitude at the first vertical grid, z_0 is the
 254 roughness length, and κ is the von Kármán constant. Here, $z_0 = 0.1 \text{ m}$ (Bou-Zeid
 255 et al., 2009) and $\kappa = 0.4$.

256 The ratio of the boundary-layer height δ of the generated outflow to the rough-
 257 ness block height in the driver region is 27.9. Note that, here, δ is defined as the
 258 height at which the mean streamwise velocity component at the outflow indicates a
 259 peak value. In Nakayama et al. (2011), the ratio of δ to the roughness block height
 260 in the driver region is 13. In addition, we confirmed that the vertical profiles of the
 261 standard deviation of each velocity component and Reynolds stress are in reason-
 262 able agreement with those obtained from wind-tunnel experiments, although the
 263 LES results underestimate the spanwise and vertical components and Reynolds-
 264 stress values relative to the wind-tunnel results (see Online Resource 1, Figure 1).

265 These results suggest that well-developed, deep turbulent flows are generated in
266 the driver region.

267 In the main region, the domain size and the total number of grid points are 12
268 km \times 2.4 km \times 1.015 km and $3000 \times 600 \times 105$, respectively. The main region
269 includes the actual buildings and structures in Kyoto, as shown in Fig. 2a. For
270 computational purposes, we set a buffer area spanning 500 m and 200 m in the
271 streamwise and spanwise directions, respectively, surrounding the actual building
272 area in the main region (not shown in Fig. 2a). The streamwise width of the area
273 was determined based on Nakayama et al. (2012), who carried out an LES inves-
274 tigation of the airflow over Tokyo. Whereas Nakayama et al. (2012) did not set a
275 buffer area in the spanwise direction, we decided that a spanwise buffer is necessary
276 to avoid building discontinuities arising from the periodic boundary conditions. In
277 this buffer area, the same roughness blocks used in the driver region are applied to
278 maintain a turbulent flow over roughness surfaces. Note that the coordinates $x = 0$
279 km and $y = 0$ km are set to the northern and western boundaries, respectively, of
280 the actual building area in the main region. Correspondingly, the inflow boundary
281 condition provided by the driver region is set at $x = -500$ m in the main region.
282 Outside of the inflow boundary, the boundary conditions of the main region are
283 the same as those in the driver region, and all grid spacings are identical to those
284 in the driver region.

285 Hereafter, the simulation using the actual buildings in Kyoto is referred to as
286 the control experiment (CTL). To reveal the effects of building-height variability,
287 we conducted an additional experiment referred to as the uniform experiment
288 (UNI) in which all building heights are set to the average of the actual building
289 heights in the main region ($h_{all} = 10.3$ m). The integration time for each of the

two experiments is 7,200 s, with the results obtained from the last 1,800 s used for the analysis of turbulent statistics. In Sect. 4.3, we confirm that the flows were in equilibrium states during this analysis period, as shown Fig. 5. In addition, as seen in Fig. 1 of Online Resource 1, the second-order moments of the inflow profiles are relatively small compared with those of the wind-tunnel experiments, which possibly influences the results presented here. However, as the same inflow condition was applied in both the CTL and UNI experiments, we can assume that any differences in the respective experimental results are unaffected by this issue.

4 Comparison with Observations

4.1 Observational Setting

The observations were performed at the Ujigawa Open Laboratory of the Disaster Prevention Research Institute, Kyoto University, during the period from 12 January to 12 February 2016. The laboratory is located in the southern part of Kyoto, and is surrounded by low-rise buildings and structures. The location of the observation site is shown in Fig. 1, which includes a meteorological observation tower of height 55 m. This tower is a unique facility first deployed in 1978 (Nakajima et al., 1979), and is currently one of the few meteorological towers operating in Japan.

A sonic anemometer (DA-600, Kaijo Co.) installed on the tower at a 25-m height measures the three velocity components as well as the air temperature at a 10-Hz sampling rate. The surrounding area up to 500 m north of the tower has only low building heights (< 25 m), enabling the assumption that observations taken by the sonic anemometer are not influenced by the strong wakes of tall buildings.

313 We also installed a Doppler lidar (WINDCUBE WLS-7, Leosphere) at the
 314 ground near the tower, from which we obtained three-component velocity measure-
 315 ments at heights ranging from 40 m to 200 m with a 20-m interval at a sampling
 316 rate of 1 Hz.

317 4.2 Data Selection

318 The observation site was included in the main region assessed in the numerical
 319 experiment for the purpose of directly comparing the LES results in the CTL
 320 experiment with the observations. As the sonic anemometer installed on the tower
 321 faces northwards, we analyzed data for dominant northerly wind directions to
 322 minimize the interference from the tower. To extract suitable periods from the
 323 observational data, we imposed two criteria for sorting values obtained from the
 324 sonic anemometer. First, a northerly flow condition was adopted by classifying 10-
 325 min averaged wind directions into 16 classes and extracting periods when northerly
 326 wind directions ($348.25^\circ - 360^\circ, 0^\circ - 11.25^\circ$) were sustained for at least 30 min.
 327 Note that the time period for the analysis of the LES data was also 30 min. Second,
 328 a neutrally stratified condition was chosen based on the Monin–Obukhov stability
 329 parameter

$$\frac{z}{L} = -\frac{(g/\bar{T})\overline{w'T'}}{u_*^3/\kappa z}, \quad (9)$$

330 so that the assumption of turbulent flows under a neutrally stratified condition in
 331 the LES model is valid. Here, L is the Obukhov length (m), g is the acceleration
 332 due to gravity (m s^{-2}), T is the air temperature (K), $\overline{w'T'}$ is the sensible heat flux
 333 (K m s^{-1}), and u_* is the friction velocity (m s^{-1}). An overbar and prime denote

334 a temporal average and fluctuation, respectively. A period for $|z/L| \leq 0.05$ (Roth,
335 2000) is regarded as fulfilling the neutrally stratified condition.

336 By imposing the above conditions on the observational data, we obtained the
337 following four 30-min periods: 0720 – 0750 LT (local time = UTC + 9 h) 22
338 January; 1650 – 1720 LT 30 January; 0740 – 0810 LT 2 February; and 1830 –
339 1900 LT 10 February, which are referred to as the D1, D2, D3, and D4 periods,
340 respectively. The wind directions for each period calculated from the averaged
341 horizontal velocity components are 4.9° , 358.8° , 353.8° , and 351.5° for the D1 to
342 D4 periods, respectively.

343 To compare the LES results with the observations, it is necessary to use airflows
344 observed at the Ujigawa Open Laboratory coming from the northern boundary of
345 the analysis region of Kyoto passing through the analysis region, and not from the
346 western or eastern boundaries. Because of the periodic conditions at the western
347 and eastern boundaries, the flow through these lateral boundaries is unlikely to
348 be accurately simulated by the LES model. This condition requires that wind
349 directions be within a range of between approximately 355° and 5° based on the
350 streamwise length and half the spanwise length of the analysis region (i.e., $\arctan(1$
351 $\text{km}/11 \text{ km})$). Overall, the wind directions in the periods D1 – D4 are almost
352 within the range of this condition, although those in the periods D3 and D4 are
353 slightly shifted westwards from the condition. We confirmed that the area within
354 at least 1 km westwards from the analysis region is dominated by land-use and
355 building types similar to those in the analysis region. Thus, we concluded that the
356 anemometer data taken during the four periods described above are appropriate
357 for comparison with the LES results. However, the wind directions measure by the
358 Doppler lidar deviate from those recorded by the sonic anemometer. The directions

359 of the Doppler lidar in the D1 and D3 periods become more westerly with height,
360 reaching 330° at a height of 200 m, while those in the D2 and D4 periods are
361 relatively constant with altitude and within a range between approximately 350°
362 and 0° . We discuss the possible influences of the variation of wind direction in
363 Sect. 4.3. As explained above, none of the observed wind directions were oriented
364 in a truly northerly fashion. Correspondingly, we rotated the streamwise directions
365 to the mean of the wind directions measured by the sonic anemometer and the
366 Doppler lidar.

367 4.3 Results

368 Figure 5a and b shows the time series of streamwise and spanwise velocity com-
369 ponents produced by the LES model and measured by the sonic anemometer at a
370 25-m height, respectively. To avoid interference from the tower on the wind-speed
371 profiles, the LES results are shown for a grid point 16 m north of the tower. It
372 is seen that the LES turbulent fluctuations in both the streamwise and spanwise
373 directions are quite comparable to those from the anemometer. Note that average
374 spanwise velocity components are nearly zero, as indicated in Fig. 5b. The stream-
375 wise velocity component is stronger in the D2 period than in the other periods.
376 Comparison of the respective weather charts for the four time periods reveals the
377 stronger wind speeds in the D2 period to be caused by a large low-pressure system
378 passing through the northwest Pacific Ocean off the coast of the Japanese Islands.

379 Figure 5c shows a comparison of the LES and observed vertical profiles of the
380 mean streamwise velocity component. Both datasets are averaged over time, and
381 the time-averaged LES data are averaged horizontally over a 16 m by 16 m area to

382 the north of the tower to increase the representativeness of the simulated flows for
383 the observation site. Note that, given the logarithmic scales used on both axes, the
384 slopes of the mean streamwise velocity component in Fig. 5c suggest a power-law
385 profile. According to Counihan (1975), the slopes of suburban and urban areas
386 range between 0.21 and 0.28, making a power-law exponent of $1/4$ suitable for ref-
387 erence, where it is seen that the slopes of the observations and the LES results are
388 very similar to this value. We also examined the respective vertical profiles of the
389 mean streamwise velocity component normalized by the mean streamwise velocity
390 component at the 25-m height (see Online Resource 1, Figure 2) and found that
391 the LES and observed mean streamwise velocity components are quantitatively
392 consistent. We conclude that this result is also good evidence for the reasonable
393 performance of our LES model. In contrast, the slopes above approximately the
394 150-m height in the D1, D2, and D3 periods appear to deviate from the reference
395 slope. In the case of the D1 and D3 periods, we assume this occurs because of the
396 change in wind direction from northerly to westerly, as described in the previous
397 subsection. Another possible explanation for the deviation at the higher levels is
398 that the stability conditions may not have been neutral at these heights during
399 the observed periods. Because there were no observational data available to clas-
400 sify the stability condition above height of the sonic anemometer at 25 m, it is
401 impossible to quantitatively reveal the stability above that height.

402 The vertical profiles of Reynolds stress in both the observations and the LES
403 results are shown in Fig. 5d. Note that the Reynolds stress is normalized by the
404 mean streamwise velocity components at each height. The Reynolds stress of the
405 LES data is averaged horizontally over the same 16 m by 16 m area used for the
406 mean streamwise velocity component. It is seen that the vertical profile of the LES

407 data is within the range of differences found in the observation periods, which is a
408 feature similar to that of the profiles normalized by the mean streamwise velocity
409 component at the 25-m height (see Online Resource 1, Figure 2). However, it is
410 necessary to be careful in comparing the LES results with the Doppler lidar data
411 because the latter might include some errors in representing perturbations of the
412 wind speed as discussed below.

413 We now compare the results for the turbulence intensity, which is the ratio
414 of the standard deviation of each velocity component σ_i to the mean streamwise
415 velocity component. As previously mentioned, the turbulence intensity was also
416 averaged horizontally in the 16 m by 16 m area. Figure 6 compares the vertical
417 profiles of turbulence intensity in the LES results and observations with the em-
418 pirical form of the ESDU (1985), which is a database providing the turbulence
419 characteristics of a neutrally stratified atmospheric boundary layer based on var-
420 ious field measurements from around the world. In Fig. 6, all sonic-anemometer
421 components fall within the rough-surface category given by the ESDU, which in-
422 dicates suburban areas with z_0 between 0.1 and 0.5. Each component simulated
423 by the LES model appears to capture the vertical distribution of that obtained by
424 the ESDU within or around its upper and lower limits, at least below about the
425 height of 150 m, while being slightly smaller than those of the sonic-anemometer
426 observations. In fact, the values obtained from the sonic anemometer lie near the
427 upper limit of the ESDU profile, suggesting that the LES results within the ESDU
428 range are generally more favourable.

429 In contrast, there appears to be large discrepancies between the Doppler lidar
430 observations and the LES results in terms of the u and v velocity components.
431 The turbulence intensities for these components measured by the Doppler lidar

432 are even larger than the upper limits of the ESDU, suggesting that the measure-
433 ments may include an overestimating bias for the turbulence intensities. It is in
434 fact commonly understood that the Doppler lidar measurements overestimate the
435 turbulence intensities for the streamwise component. This characteristic was noted
436 in Cañadillas et al. (2011), who showed that the results produced by the Doppler
437 lidar observations are larger than those of sonic anemometers at various wind
438 speeds and altitudes, and the deviations become larger with the decrease in wind
439 speed. A close look at Figs. 5c and Fig. 6a indicates that the difference between
440 Doppler lidar and the ESDU in terms of the streamwise turbulence intensity be-
441 low 100 m decreases as the streamwise velocity component increases, in apparent
442 confirmation of the finding of Cañadillas et al. (2011). For the Doppler lidar data
443 above 100 m, changes in the wind direction and uncertainty in the stability, as
444 revealed in the mean streamwise velocity component, may contribute to this over-
445 estimation. An overestimating tendency in the lidar data can also be found for the
446 spanwise component, which has a mean value of nearly zero.

447 The vertical component produced by the LES results appears to be consistent
448 with both the lidar data and the ESDU profile, but the lidar tends to underestimate
449 the vertical turbulence intensity, particularly in weaker wind-speed conditions.

450 Figure 7 shows the power spectra of the time series of each velocity component
451 obtained from the LES results and the sonic anemometer at a height of 25 m. The
452 spectra were calculated from the time series shown in Fig. 5, and the frequency
453 f and velocity spectra $E(f)$ are normalized in dimensionless form. The figure
454 includes the empirical reference from Kaimal et al. (1972) derived from observa-
455 tions over a rural region. A close agreement is seen between the sonic anemometer
456 and the reference results for all three components. The spectra from the sonic

457 anemometer clearly represent an inertial subrange with a $2/3$ slope. Comparison
458 of the LES spectra with the observations and empirical reference reveals that the
459 spectra of the u and v components of the LES data are similar to those of the sonic
460 anemometer data except in the highest frequency range. The lower frequency por-
461 tion of the inertial subrange appears to be well reproduced for these components
462 in the LES results.

463 However, the LES model is able to reproduce the vertical velocity components
464 in only the lowest frequency portion of the inertial subrange. It is possible that
465 the grid spacing used in our modelling is insufficient for resolving the smallest
466 eddies and their corresponding vertical motion. Further increases in the vertical
467 resolution may be required to represent the small-scale vertical motion likely to be
468 induced at the edges of buildings. However, we note that the spectral peak of the
469 w component in the LES results agrees well with that of the sonic anemometer.

470 From the above comparisons, we conclude that the use of our LES model leads
471 to a reasonable reproduction of the turbulent boundary-layer flow over actual
472 buildings under a neutral stability condition, at least up to a height of about 150
473 m. We emphasize that, in general, the results produced by our LES model agree
474 favourably with the observations within the range of differences among the chosen
475 periods (D1 – D4), even though our inflow condition employed an idealized turbu-
476 lent flow generated in the driver region without realistic meteorological conditions.
477 These results are sufficient here because our analysis of building-height variabil-
478 ity focuses on altitudes below approximately 25 m (i.e., at height $z = 2.5h_{all}$),
479 where the LES results show an especially close agreement with the observations,
480 as shown in Figs. 6 and 7.

481 5 Sensitivity to Building-Height Variability

482 5.1 General characteristics of turbulent flows

483 We now focus on the overall characteristics of turbulent flows in the CTL and UNI
484 experiments, starting with the differences between the respective experiments.

485 Figure 8a and b shows the vertical profiles of the space- and time-averaged
486 streamwise velocity component $\langle \bar{u} \rangle_{all}$ and Reynolds stress $-\langle \overline{u'w'} \rangle_{all}$ over
487 the entire main region for the CTL and UNI experiments, respectively. Here, the
488 angled brackets denote a spatial average, while the subscript *all* refers to the overall
489 main region. Note that the values are normalized by the mean streamwise velocity
490 component U_∞ at the height of the boundary-layer (δ). The mean streamwise ve-
491 locity components above height $z = h_{all}$ (i.e., above the canopy layer) are lower in
492 the CTL experiment than in the UNI experiment. In contrast, the velocities below
493 height $z = h_{all}$ for the CTL experiment are higher than in the UNI experiment.
494 The Reynolds stress above height $z = h_{all}$ in the CTL experiment is larger than
495 that in UNI experiment. Furthermore, the level of peak Reynolds stress is higher
496 in the CTL experiment than in the UNI experiment.

497 These differences between the CTL and UNI results can be attributed to the
498 effects of building-height variability. Using the LES results of flows over idealized
499 arrays of roughness blocks, Nakayama et al. (2011) showed that the mean velocity
500 above the building height decreases with increasing building-height variability, and
501 that the magnitude and height of the peak of the Reynolds stress both increase
502 with building-height variability. Our results in terms of the streamwise velocity
503 component and Reynolds stress are consistent with the results of Nakayama et al.
504 (2011).

505 Xie et al. (2008) carried out an LES investigation over block arrays with ran-
 506 dom and uniform heights, and found that both types of arrays produced similar
 507 turbulent kinetic energies below the average building height. The Reynolds stresses
 508 produced in the CTL and UNI experiments below height $z = h_{all}$ are consistent
 509 with their results. From Fig. 8b, it is seen that the Reynolds stress in the UNI
 510 experiment sharply increases around height $z = h_{all}$, which is likely caused by
 511 the presence of the uniform tops of buildings in the UNI experiment, resulting in
 512 sharp wind shear and the generation of turbulence.

513 In Coceal et al. (2006), the velocity components u_i were decomposed as

$$u_i = \langle \bar{u}_i \rangle + \bar{u}_i'' + u_i', \quad (10)$$

514 where $\langle \bar{u}_i \rangle$ are the time- and space-averaged velocities, \bar{u}_i'' is the spatial vari-
 515 ation of the time-averaged velocity, and u_i' is the turbulent fluctuation. Coceal
 516 et al. (2006) showed that dispersive flux, which is defined as $\langle \bar{u}'' \bar{w}'' \rangle$, signifi-
 517 cantly contributes to the total momentum flux in the canopy layer in which the
 518 time-averaged velocities are spatially inhomogeneous. The vertical profiles of the
 519 dispersive flux normalized by U_∞ in the CTL and UNI experiments are shown
 520 in Fig. 8c. Although the dispersive fluxes for both experiments have peaks just
 521 below height $z = h_{all}$, the magnitude of the peak in the UNI experiment is larger
 522 than that in the CTL experiment. The UNI profile decreases sharply with height
 523 above the height of the peak. Above height $z = h_{all}$, the dispersive flux in the
 524 CTL experiment is larger than that in the UNI experiment up to about height
 525 $z = 3.5h_{all}$. Xie et al. (2008) performed an LES investigation to compare the
 526 dispersive flux in random and uniform block arrays. Their results suggest that
 527 both types of dispersive flux have peaks near the average building height, that

528 the peaks obtained from uniform block arrays are stronger than those for random
 529 block arrays, and that the dispersive flux of uniform block arrays decreases much
 530 more abruptly with increasing height above the height of the peak than that of
 531 random block arrays. These characteristics are qualitatively consistent with our
 532 results. The dispersive fluxes in both the CTL and UNI experiments appear not to
 533 decrease linearly with height because the time-averaged velocities are not spatially
 534 homogeneous at heights above the canopy layer. Based on the results shown in Fig.
 535 8a and b, we focus on the height of $z = 0.5h_{all}$ at which the difference between
 536 the CTL and UNI experiments is small, and the height of $z = 2.5h_{all}$ where clear
 537 differences are seen between the respective experiments.

538 Figure 9 shows the fields of time-averaged streamwise velocity component nor-
 539 malized by U_∞ for the CTL and UNI experiments over an upstream region ($x =$
 540 $1 - 5$ km) in which the business districts are located. The difference between
 541 the respective experimental results for the region appears to be small at height
 542 $z = 0.5h_{all}$ except in areas along a major street around $y = 1.3$ km. This is likely
 543 caused by a stronger convergence of the streamwise velocity components on the
 544 street in the UNI experiment owing to enhancements arising from the presence of
 545 uniform-height buildings (i.e., in the UNI experiment, all lower building heights are
 546 raised to $z = h_{all}$). The velocity-deficit regions are reproduced at height $z = 2.5h_{all}$
 547 behind buildings in the CTL experiment, which contrasts to the smooth field of
 548 time-averaged streamwise velocity components at height $z = 2.5h_{all}$ in the UNI
 549 experiment.

550 Figure 10 shows the fields of Reynolds stress normalized by U_∞ for the CTL
 551 and UNI experiments over the upstream region. While the features are quite similar
 552 at height $z = 0.5h_{all}$, the field at height $z = 2.5h_{all}$ in the CTL results has larger

553 values behind the buildings than in the UNI results, which indicates the important
554 role of sparsely and randomly distributed buildings at and above height $z = 2.5h_{all}$
555 in generating turbulence in the CTL experiment.

556 5.2 Analysis of Roughness Parameter

557 To quantitatively reveal the effects of building-height variability, we examined the
558 relationships between the turbulent statistics and roughness parameters. The plan-
559 area index λ_p is used for this analysis because the CTL and UNI experiments have
560 the same values for this parameter. Turbulent statistics were derived in each 1 km
561 by 1 km area in a manner similar to that used to find the roughness parameters
562 in Sect. 2.

563 5.2.1 Reynolds stress

564 Figure 11 shows how the Reynolds stress normalized by U_∞ in the CTL and
565 UNI experiments changes as a function of λ_p at the heights of $z = 0.5h_{all}$ and
566 $z = 2.5h_{all}$. The brackets with subscript 1 km^2 indicate spatial averaging over a 1
567 km by 1 km area. The Reynolds stress at height $z = 0.5h_{all}$ is very similar for the
568 two experiments, which is consistent with the features shown in Fig. 10a and b.
569 By contrast, the values at height $z = 2.5h_{all}$ in the CTL experiment increase with
570 λ_p , while those in the UNI experiment are nearly independent of λ_p . In addition,
571 the differences between the CTL and UNI results at height $z = 2.5h_{all}$ are more
572 apparent when $\lambda_p > 0.32$.

573 As shown in Figs. 9 and 10, the difference between the CTL and UNI exper-
574 iments in terms of building distributions at height $z = 2.5h_{all}$ has a significant

575 effect on the turbulent flow results. To interpret this difference, we calculated the
 576 respective plan-area indices λ_p at this altitude; i.e., for each experiment, if the
 577 building height in a grid cell is below $z = 2.5h_{all}$, the grid cell is regarded as
 578 having no buildings. Figure 12a and b shows λ_p at heights $z = 0.5h_{all}$ (denoted
 579 by $\lambda_{p, 0.5h_{all}}$) and $z = 2.5h_{all}$ ($\lambda_{p, 2.5h_{all}}$), respectively, plotted against λ_p at the
 580 surface for both the CTL and UNI experiments. Note that, in the UNI experiment,
 581 the value of $\lambda_{p, 0.5h_{all}}$ is the same as that of λ_p at the surface, and that $\lambda_{p, 2.5h_{all}}$
 582 is zero for this experiment. The difference between the CTL and UNI experiments
 583 in terms of $\lambda_{p, 0.5h_{all}}$ is not very large, confirming the similarity of the respec-
 584 tive Reynolds stresses at height $z = 0.5h_{all}$ in Fig. 10. In the CTL experiment,
 585 $\lambda_{p, 2.5h_{all}}$ rapidly increases if λ_p exceeds 0.32, which appears to be consistent with
 586 the Reynolds-stress feature in the CTL experiment at height $z = 2.5h_{all}$ as seen
 587 in Fig. 10. Based on these results, we suggest that the Reynolds stress from the
 588 CTL experiment at height $z = 2.5h_{all}$ becomes stronger at $\lambda_p > 0.32$ because
 589 some building clusters are still present at height $z = 2.5h_{all}$ in this experiment.

590 The frontal-area index λ_f is another important parameter for describing the ge-
 591 ometrical characteristics of urban areas. Here we examine the frontal area of build-
 592 ings above the height h_{all} . Figure 12c shows λ_f above height $z = h_{all}$ ($\lambda_{f, h_{all}}$)
 593 plotted against λ_p for the CTL experiment (the figure does not include the corre-
 594 sponding values for the UNI experiment owing to the absence of buildings at that
 595 altitude). It is seen that $\lambda_{f, h_{all}}$ increases with λ_p , and sharply increases when
 596 $\lambda_p > 0.32$. These features agree well with the characteristics determined above
 597 for $\lambda_{p, 2.5h_{all}}$ and the Reynolds stress. According to these results, the effects of
 598 building-height variability on the Reynolds stress increase with λ_p when λ_p is

599 greater than 0.32, and are closely linked to the higher values of λ_p , $2.5h_{all}$ and
600 λ_f , h_{all} at such values of λ_p .

601 Interestingly, Zaki et al. (2011) found that the drag coefficient C_d in wind-
602 tunnel experiments, which is relevant to the Reynolds stress, increases with λ_p
603 when $\lambda_p > 0.32$ in flows over block arrays with random heights. A similar feature
604 can also be found for the Reynolds stress and C_d in the LES investigation by
605 Nakayama et al. (2011). According to Zaki et al. (2011), this is because taller
606 buildings, which contribute largely to the total drag in a block array (Xie et al.,
607 2008), tend to be sparsely distributed and, therefore, despite the increase in λ_p ,
608 the flow pattern does not enter a skimming flow regime (Oke, 1988). Based on
609 these previous studies and our results, $\lambda_p \approx 0.3$ can be regarded as a threshold
610 at which the effects of building-height variability on the turbulent flow become
611 apparent in various cities.

612 5.2.2 Momentum transfer according to a quadrant analysis

613 As described in Sect. 1, turbulent coherent structures over urban surfaces are re-
614 lated to the physical process of turbulent momentum transfer. A quadrant analysis
615 is a useful method for identifying the characteristics of the momentum transfer as-
616 sociated with coherent structures, and has been used in numerous studies of wall
617 turbulence (Wallace, 2016). This method divides the Reynolds stress into four
618 components based on the signs of u' and w' : outwards interaction (quadrant 1,
619 $u' > 0$, $w' > 0$); ejection (quadrant 2, $u' < 0$, $w' > 0$); inwards interaction
620 (quadrant 3, $u' < 0$, $w' < 0$) and sweep (quadrant 4, $u' > 0$, $w' < 0$). Raupach
621 (1981) introduced a conditional averaging using the threshold H to investigate the

622 contribution to the Reynolds stress from the i th quadrant as

$$\langle u'w' \rangle_{i, H} = \lim_{T \rightarrow \infty} \frac{1}{T} \int_0^T u'(t)w'(t)I_{i, H}[u'(t), w'(t)]dt, \quad (11)$$

623 where the trigger indicator $I_{i, H}$ is defined as

$$I_{i, H}(u', w') = \begin{cases} 1, & \text{if } (u', w') \text{ is in quadrant } i \text{ and if } |u'w'| \geq H|\overline{u'w'}|, \\ 0, & \text{otherwise.} \end{cases} \quad (12)$$

624 The fraction of stress exceeding the threshold, which indicates the relative quantity

625 of the i th quadrant, is

$$S_{i, H} = \langle u'w' \rangle_{i, H} / \overline{u'w'}. \quad (13)$$

626 It is noted that the relationship

$$S_{1, 0} + S_{2, 0} + S_{3, 0} + S_{4, 0} = 1 \quad (14)$$

627 holds only for $H = 0$. When the Reynolds stress is negative (as is normally seen in

628 the boundary layer), $S_{2, 0}$ and $S_{4, 0}$ are positive, while $S_{1, 0}$ and $S_{3, 0}$ are negative.

629 Ejections and sweeps contribute to the downwards momentum flux, and are

630 considered to be associated with organized turbulent motions as indicated in Sect.

631 1. Thus, the magnitude of ejections and sweeps is a good indicator for determining

632 the characteristics of turbulent flows.

633 To further reveal the relative roles of ejections and sweeps in vertical momen-

634 tum transfer, we introduce the two parameters

$$\Delta S_0 = S_{4,0} - S_{2,0}, \quad (15)$$

$$Ex = (S_{1,0} + S_{3,0}) / (S_{2,0} + S_{4,0}), \quad (16)$$

635 where ΔS_0 is the difference between sweeps and ejections, and Ex , which is called

636 the exuberance (Shaw et al., 1983), is the ratio of unorganized (S_1 and S_3) mo-

637 tions to organized (S_2 and S_4) motions. The exuberance indicates the efficiency

638 of the vertical momentum flux. Christen et al. (2007) used these parameters to
639 investigate vertical momentum exchange in an urban district and elucidated the
640 roles of coherent structures in momentum transport.

641 Figure 13 shows the vertical profiles of ΔS_0 and Ex for the CTL and UNI
642 experiments, which are averaged temporally and spatially in a manner similar to
643 the profiles in Fig. 8, where ΔS_0 in the CTL experiment is generally larger than in
644 the UNI experiment, except for heights around h_{all} . This feature of ΔS_0 contrasts
645 to the vertical profile of the Reynolds stress shown in Fig. 8b, which indicates that
646 the Reynolds stress is nearly identical in the CTL and UNI experiments below
647 height $z = 0.5h_{all}$. This suggests that, despite the similarities in the Reynolds
648 stress seen in the two experiments, the building-height variability in the CTL
649 experiment changes the ratio of ejections to sweeps within the building canopy
650 layer. In the upper layer from heights $z = 2.5h_{all}$ to $z = 10h_{all}$, both ΔS_0 and the
651 Reynolds stress are larger in the CTL experiment than in the UNI experiment.
652 We consider that the increased Reynolds stress in this upper layer in the CTL
653 experiment is caused by a sweep-dominated vertical flux.

654 Figure 13b shows the value of Ex below $z = 2.5h_{all}$ in the CTL experiment
655 to be smaller than that in the UNI experiment. Below $z = 0.5h_{all}$, the decrease
656 in Ex appears to be more pronounced in the CTL experiment than in the UNI
657 experiment even though the respective Reynolds stresses are similar, as shown in
658 Fig. 8b. This indicates that the efficiency of the vertical momentum flux in the
659 canopy layer is reduced by building-height variability. In contrast, the values of
660 Ex in the CTL and UNI experiments are very similar at altitudes above height
661 $z = 2.5h_{all}$, indicating that the efficiency of the momentum flux above these
662 altitudes is similar for both experiments.

Based on the differences between the vertical profiles shown in Fig. 13, we focus on the heights $z = 0.5h_{all}$ and $z = 2.5h_{all}$ to reveal the relationship between building-height variability and turbulent-flow characteristics. Figure 14a and b shows variations in ΔS_0 against λ_p in the CTL and UNI experiments at these two altitudes. At height $z = 0.5h_{all}$, sweeps are dominant among the contributions to the Reynolds stress for both experiments, which is consistent with previous results showing a stronger contribution of sweeps to the total momentum flux than ejections near and below the tops of block arrays (Raupach 1981; Coceal et al. 2007a). From Fig. 13a, it is seen that the contribution of sweeps in the CTL experiment is larger than in the UNI experiment. By contrast, the value of ΔS_0 at height $z = 0.5h_{all}$ appears to be independent of λ_p in both experiments. However, at height $z = 2.5h_{all}$, the value of ΔS_0 in the CTL experiment increases with λ_p when $\lambda_p > 0.32$, while in the UNI experiment, it is independent of λ_p . The increase in ΔS_0 in the CTL experiment is consistent with the Reynolds-stress results shown Fig. 11b, thus suggesting that sweeps contribute to the increase in Reynolds stress for $\lambda_p > 0.32$. Similar results were noted by Kanda (2006).

Figure 14c and d shows Ex plotted against λ_p in the CTL and UNI experiments at heights $z = 0.5$ and $2.5h_{all}$, respectively, where the difference in the value of Ex at height $z = 0.5h_{all}$ increases with λ_p , suggesting the dominance of unorganized structures as λ_p increases. As shown in Fig. 13b, at height $z = 2.5h_{all}$ the values of Ex in both experiments are practically independent of λ_p .

By setting H in Eq. 12 to a value larger than zero, we evaluate the extent to which extreme instantaneous momentum fluxes contribute to the total Reynolds stress in a certain period. We define the percentage contribution to the Reynolds stress of a value of $u'w'$ larger than the Reynolds stress by a factor of H using

$$E_H \equiv \sum_{i=1}^4 S_{i, H} = \sum_{i=1}^4 \langle u'w' \rangle_{i, H} / \overline{u'w'}. \quad (17)$$

688 Unlike in Raupach (1981) in which each component of the momentum flux was
 689 evaluated, all of the components in Eq. 17 are added to assess the total of the
 690 extreme momentum fluxes. We set $H = 20$ here to extract extreme values of $u'w'$,
 691 with qualitatively similar results also found with $H = 15$ and $H = 10$. Thus, $H =$
 692 20 is assumed to be a representative value.

693 Figure 14e and f shows the variations of E_{20} with λ_p at heights $z = 0.5$ and
 694 $z = 2.5h_{all}$, respectively, for both experiments. The results at height $z = 0.5h_{all}$
 695 reveal small differences between the CTL and UNI experiments and are, in general,
 696 larger than those at height $z = 2.5h_{all}$. This indicates that the flow is highly
 697 turbulent at height $z = 0.5h_{all}$, and that the extreme values of the momentum
 698 flux contribute more significantly to the total momentum flux at this altitude.
 699 However, as the magnitude of $u'w'$ itself is low at height $z = 0.5h_{all}$, the effects of
 700 the fluctuation itself may not be very strong. It is seen that, at height $z = 2.5h_{all}$,
 701 the value of E_{20} in the CTL experiment increases with λ_p , but is independent of
 702 λ_p in the UNI experiment. Moreover, the shape of the relationship between E_{20}
 703 at height $z = 2.5h_{all}$ and λ_p in the CTL experiment appears to be quite similar to
 704 that between $\lambda_{p, 2.5h_{all}}$ and λ_p shown in Fig. 12b. This suggests that increasing
 705 the number of buildings at height $z = 2.5h_{all}$ generates highly turbulent flows at
 706 higher values of λ_p .

707 The increase in the contribution from extreme values of $u'w'$ to the Reynolds
 708 stress at height $z = 2.5h_{all}$ in the CTL experiment occurs because the building-
 709 height variability in this experiment leads to a higher momentum flux at this

710 altitude as clearly indicated in Fig. 15a, which shows the horizontal cross-section
711 of E_{20} over a 1 km by 1 km area within one of the business districts. It is seen
712 that high values of E_{20} appear in areas around randomly and sparsely distributed
713 buildings. In contrast, areas with higher E_{20} values also correspond to areas with a
714 weak Reynolds stress and small value of Ex (see Fig. 15b and c), which indicates
715 the small contribution of the extreme momentum flux around buildings to the
716 total momentum flux, and is not related to organized turbulent motion. From
717 the features demonstrated in Figs. 14 and 15, it is seen that the turbulent flow
718 characteristics and contributions of extreme momentum fluxes are significantly
719 influenced by the presence of buildings with significant height variability.

720 We have shown the qualitative consistency of the Reynolds stress and quadrant
721 analysis results, if averaged both in time and space, with that over block arrays
722 with variable height. In contrast, the inhomogeneous profiles of the turbulent-flow
723 characteristics (Fig. 15) suggest that the local characteristics of the turbulent flow
724 over urban surfaces are significantly influenced by the inhomogeneity of actual
725 urban buildings, and would not be expected to be similar to that over idealized
726 block arrays.

727 **6 Summary and Conclusions**

728 An LES investigation of the turbulent flow over the city of Kyoto has been con-
729 ducted to investigate the effects of building-height variability on the turbulence
730 in the lower part of the urban boundary layer. A digital surface model data has
731 reproduced the actual buildings of Kyoto in the LES model.

732 We used roughness parameters such as H_{ave} , σ_H , λ_p , and λ_f to evaluate the
733 morphological characteristics of buildings, and compared these parameters with
734 those derived for Tokyo, Nagoya as well as for North American and European
735 cities. For $\lambda_p > 0.3$, the value of λ_f for Kyoto is small compared with the em-
736 pirical values for Tokyo and Nagoya, but similar to those obtained for European
737 cities. The relationship between H_{ave} and σ_H in Kyoto agrees closely with the em-
738 pirical profile. From these comparisons, the building morphological characteristics
739 of Kyoto indicate a dense distribution, and buildings with a variety of heights.

740 We compared the LES results with observations of atmospheric turbulence
741 obtained using a sonic anemometer and a Doppler lidar at the Ujigawa Open
742 Laboratory, which is an area included in the main region of the LES model. For
743 this comparison, certain periods were extracted from the total set of observations to
744 meet the weather conditions assumed in the LES model. The model is to reproduce
745 the observed characteristics of turbulence up to a height of about 150 m.

746 We carried out two experiments: one modelling the actual buildings of Kyoto
747 (CTL), and one (UNI) in which all building heights were set to the average building
748 height in the main region of the city h_{all} . We find small differences between the
749 CTL and UNI experiments in terms of the mean streamwise velocity component
750 and the Reynolds stress at height $z = 0.5h_{all}$, but large differences at height
751 $z = 2.5h_{all}$. The spatial fields of time-averaged streamwise velocity components
752 and Reynolds stresses produced in the CTL experiment indicate regions of reduced
753 velocity and strong Reynolds stress behind sparsely and randomly distributed
754 buildings at height $z = 2.5h_{all}$; this contrasts with the UNI results, in which
755 these fields at height $z = 2.5h_{all}$ are smooth. We investigated the relationships
756 between turbulent statistics and λ_p evaluated over 1 km by 1 km areas to reveal

757 the differences between the CTL and UNI experiments. The Reynolds stress in the
758 CTL experiment at height $z = 2.5h_{all}$ is larger than that in the UNI experiment
759 when $\lambda_p > 0.32$, while the Reynolds stress at height $z = 0.5h_{all}$ is similar for
760 both experiments. We suggest that the increase in the Reynolds stress at height
761 $z = 2.5h_{all}$ is caused by the presence of some building clusters at height $z = 2.5h_{all}$
762 in the CTL experiment, and that a value of λ_p of about 0.3 is the threshold above
763 which the effects of building-height variability become obvious over various urban
764 surfaces.

765 A quadrant analysis was used to investigate the characteristics of turbulent
766 coherent flows. Sweeps in the CTL experiment at height $z = 2.5h_{all}$ are found
767 to increase with λ_p for $\lambda_p > 0.32$, which is similar to that seen in the Reynolds
768 stress for $\lambda_p > 0.32$, suggesting the increase in Reynolds stress is caused by the
769 presence of sweeps. The vertical momentum flux in the CTL experiment is less
770 efficient than that in the UNI experiment at height $z = 0.5h_{all}$, which indicates
771 that the building-height variability in the CTL experiment reduces the efficiency
772 of the flux in the canopy layer.

773 The contributions of the extreme instantaneous momentum flux to the total
774 Reynolds stress were also investigated. The amount of extreme momentum flux
775 in the CTL experiment at height $z = 2.5h_{all}$ depends strongly on the presence
776 of buildings at this altitude. Examination of horizontal cross-sections reveals that
777 areas with extreme momentum fluxes are distributed around buildings. However,
778 the efficiency of the Reynolds stress and momentum flux are small in areas with
779 an extreme momentum flux, implying its negligible contribution around build-
780 ings to the net Reynolds stress, as well as the lack of association with coherent
781 turbulent motions. The relationships between turbulent coherent structures and

782 building-height variability were investigated through the use of space- and time-
783 averaged profiles. However, future research on turbulent coherent structures over
784 urban surfaces should focus on instantaneous and local structures, such as vortex
785 structures behind high, isolated buildings (Park et al., 2015), and flow patterns
786 in block arrays associated with coherent structures above blocks (Inagaki et al.,
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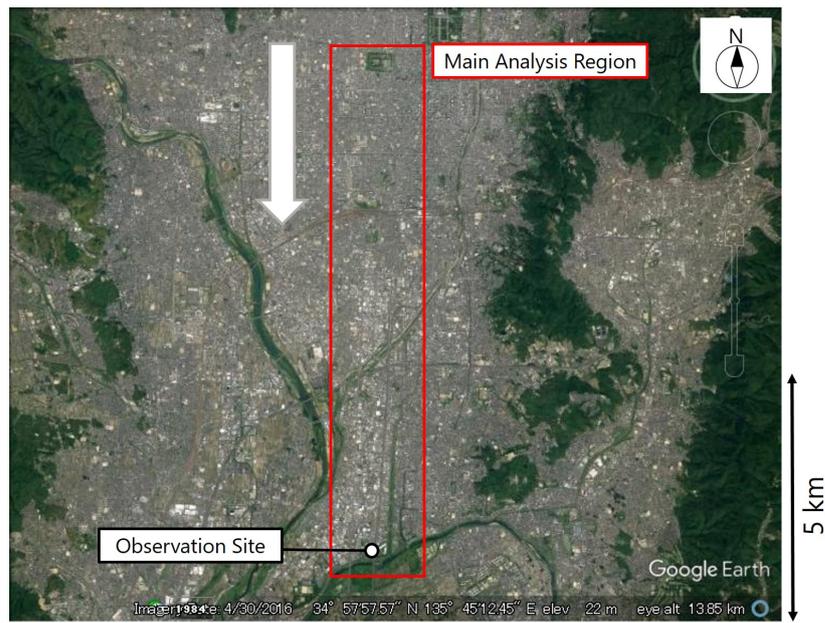


Fig. 1 The study area in which the LES model and observations were carried out is indicated by the red box. The observational site of the Disaster Prevention Research Institute, Kyoto University, is indicated by the white circle. The white arrow indicates the streamwise wind direction. The satellite picture is taken from Google Earth.

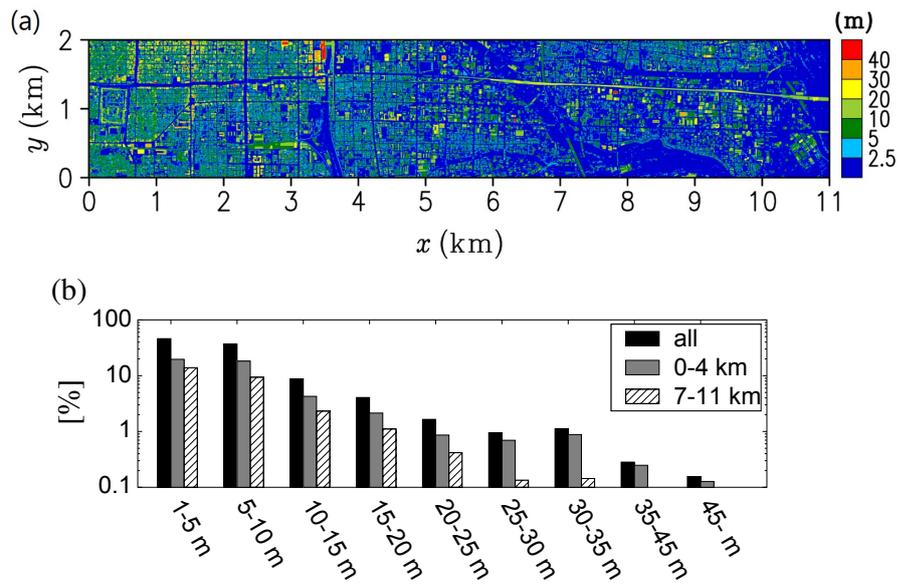


Fig. 2 (a) Distribution of building and structure heights in the analysis region of Kyoto. (b) Frequency distribution of building heights in the analysis region. The black bar indicates the frequency distribution of buildings in the overall region, while the grey and hatched bars indicate the frequency distributions of buildings in the regions with $x = 0 - 4$ km and with $x = 7 - 11$ km, respectively.

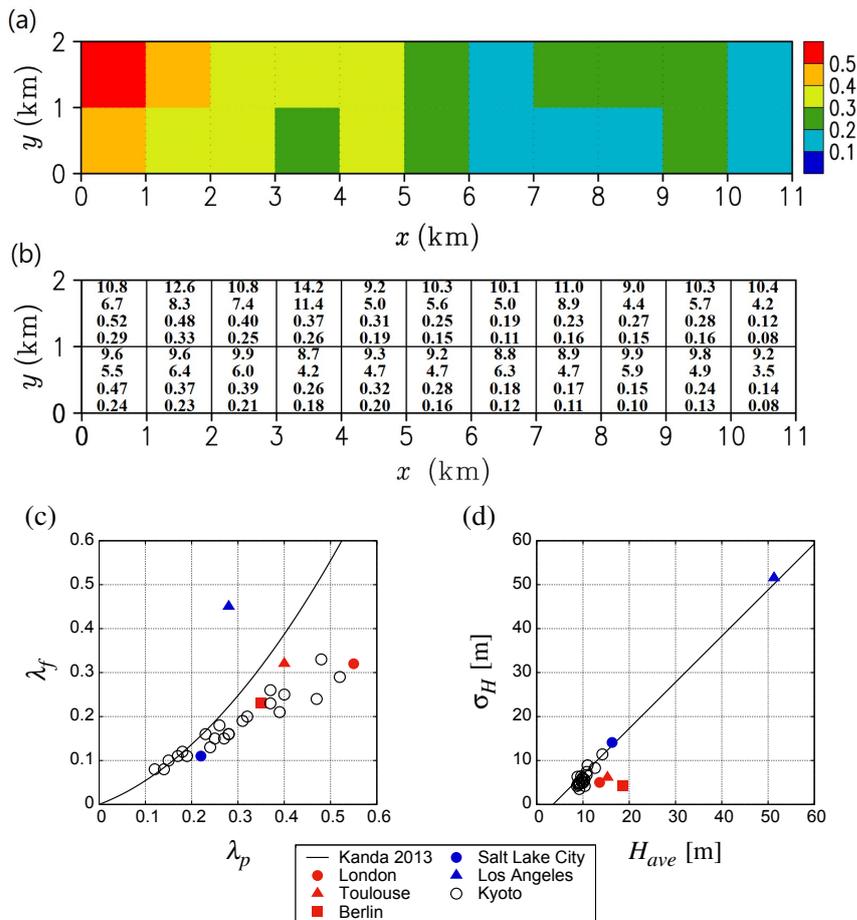


Fig. 3 (a) λ_p calculated for 1 km by 1 km areas over the analysis region. (b) Roughness parameters calculated for 1 km by 1 km areas over the analysis region. In each box, the first row is H_{ave} , the second is σ_H , the third is λ_p , and the fourth is λ_f . Scatter plots (c) between λ_p and λ_f and (d) between σ_H and H_{ave} . The black lines in (c) and (d) indicate the empirical relationships derived from Tokyo and Nagoya, respectively, by Kanda et al. (2013). The values of Salt Lake City and Los Angeles in North America and London, Toulouse, and Berlin in Europe, as indicated by the lower legend, are obtained from Ratti et al. (2002).

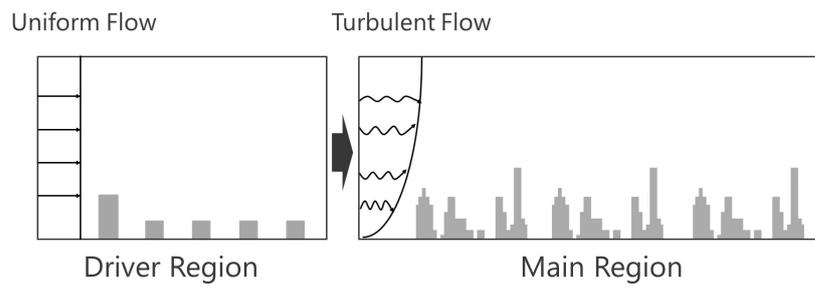


Fig. 4 Schematic of turbulent flows formed in the driver region and imposed on the main region as the inflow condition.

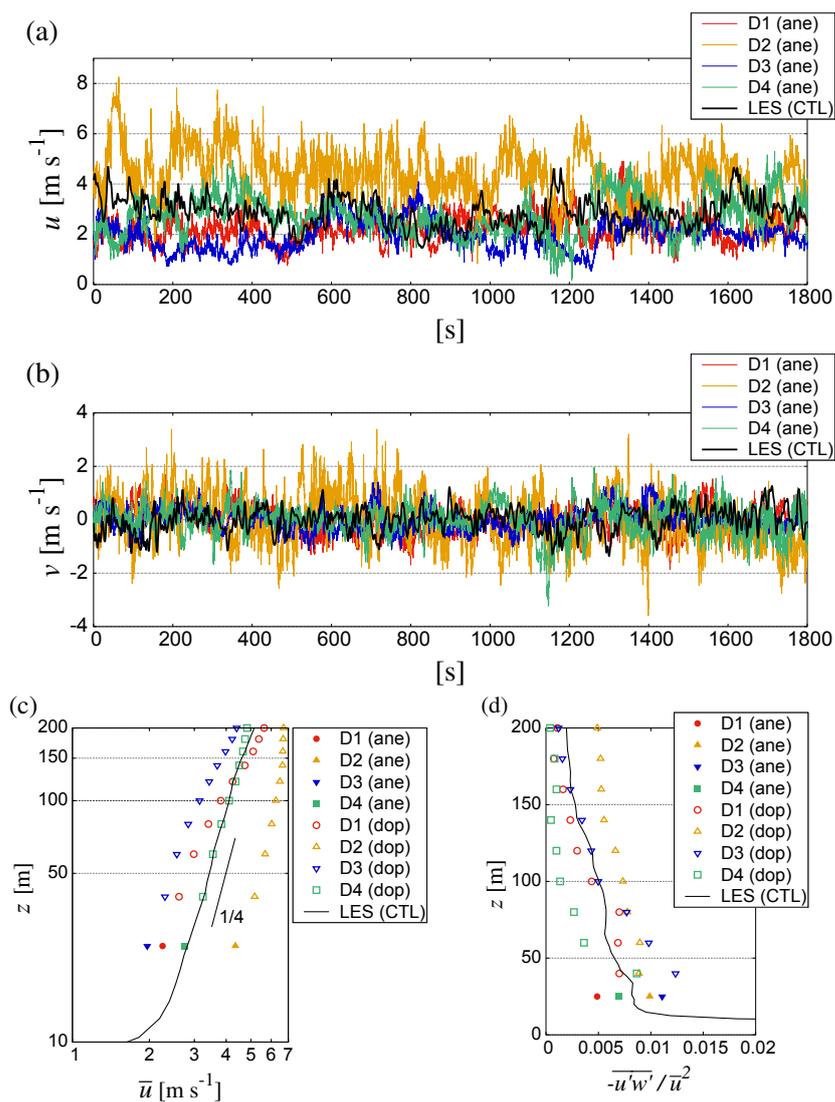


Fig. 5 Time series of (a) streamwise and (b) spanwise velocity components at 25-m height observed by the sonic anemometer during periods D1 – D4 and simulated by the LES model. Vertical profiles of (c) mean streamwise velocity component and (d) Reynolds stress normalized by the mean streamwise velocity component in the observations and the LES results. Note that the profiles in (c) are plotted on logarithmic axes. A line with a slope 1/4 is also plotted for reference. Here, ‘ane’ and ‘dop’ refer to the observations by the anemometer and Doppler lidar, respectively.

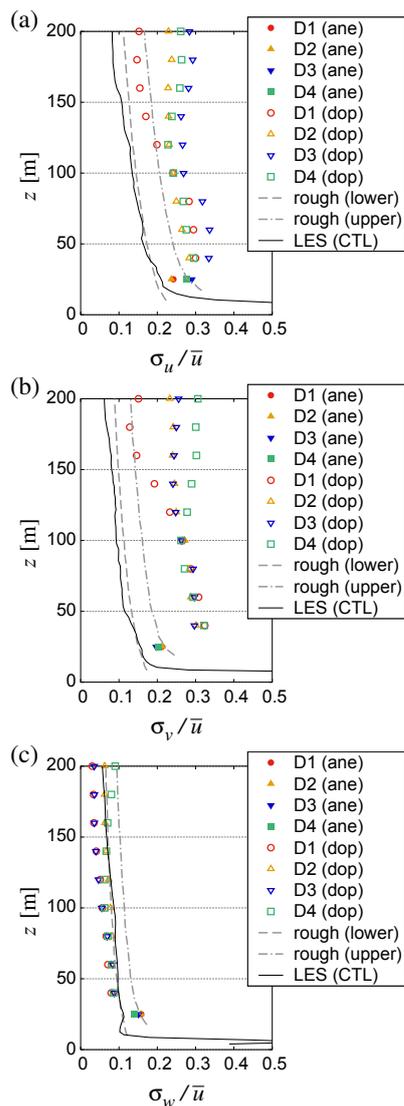


Fig. 6 Vertical profiles of turbulence intensity from the observations, the LES model, and the empirical profiles provided by the ESDU (1985) for the (a) u , (b) v , and (c) w components. The dashed and dashed-dotted lines indicate the upper and the lower limits, respectively, of the rough-surface class based on the ESDU (1985).

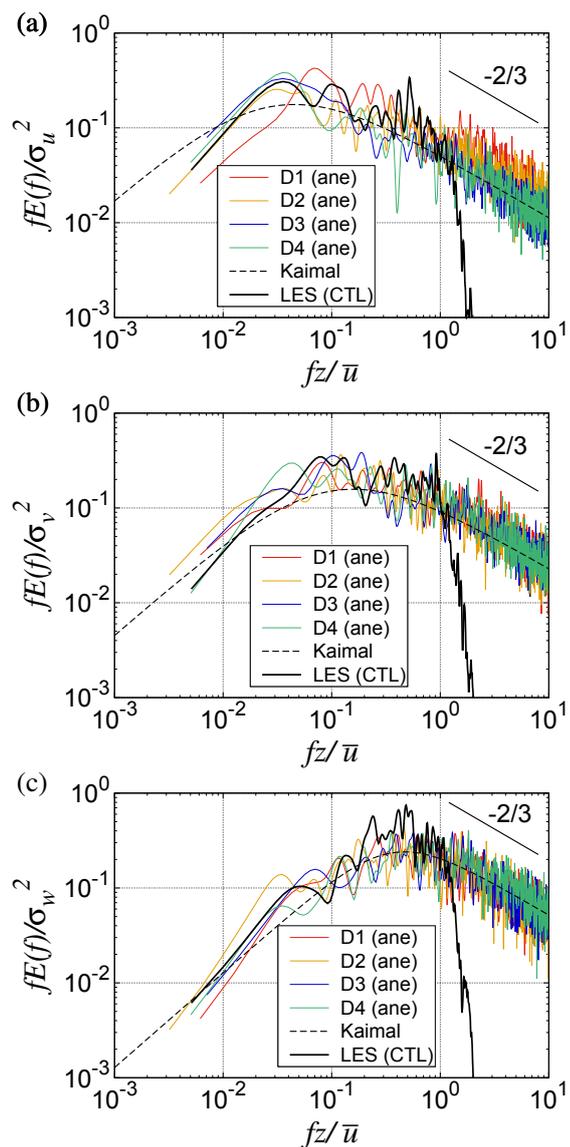


Fig. 7 Power spectra obtained from the sonic anemometer and the LES model at 25-m height plotted on logarithmic axes: (a) u , (b) v , and (c) w components. The dashed line indicates the empirical profile over a rural surface proposed by Kaimal et al. (1972).

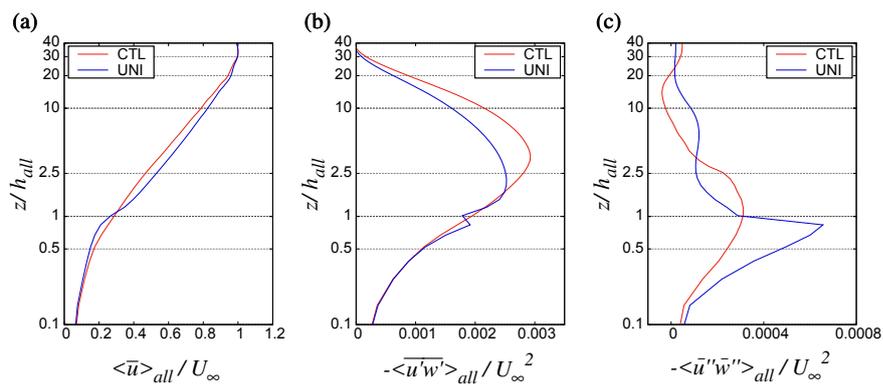


Fig. 8 Vertical profiles of (a) time-averaged streamwise velocity component, (b) Reynolds stress, and (c) dispersive flux averaged spatially over the main region. These values are normalized by U_{∞} . Red and blue lines denote the result of the CTL and UNI experiments, respectively. The vertical axis is normalized by $z = h_{all}$. Note that a logarithmic scale is used for the vertical axis.

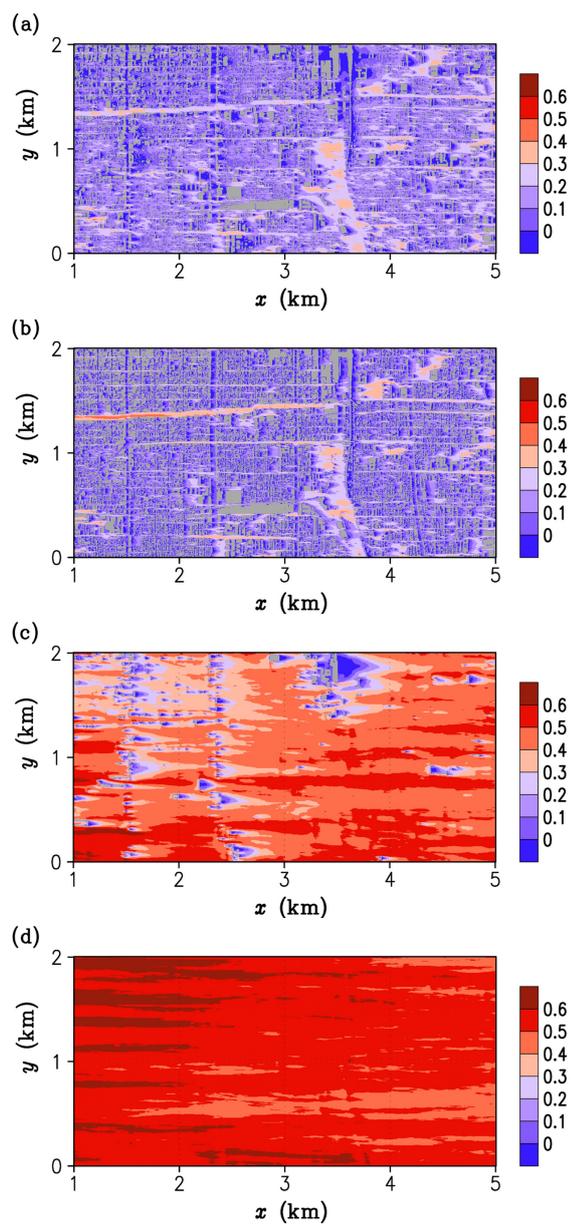


Fig. 9 Horizontal cross sections of the time-averaged streamwise velocity component normalized by U_∞ in (a) the CTL experiment at height $z = 0.5h_{all}$, (b) the UNI experiment at height $z = 0.5h_{all}$, (c) the CTL experiment at height $z = 2.5h_{all}$, and (d) the UNI experiment at height $z = 2.5h_{all}$. An upstream part of the main region is shown. The legend indicating the wind speed is present to the right of each panel. The grey shading indicates buildings.

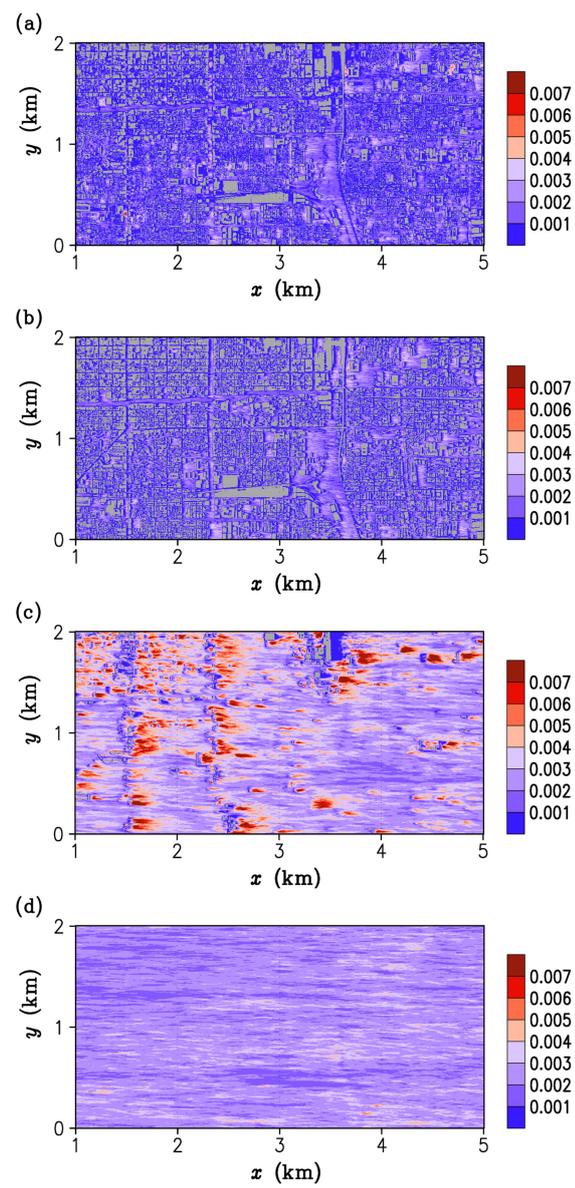


Fig. 10 As Fig. 9, except with the corresponding Reynolds-stress results.

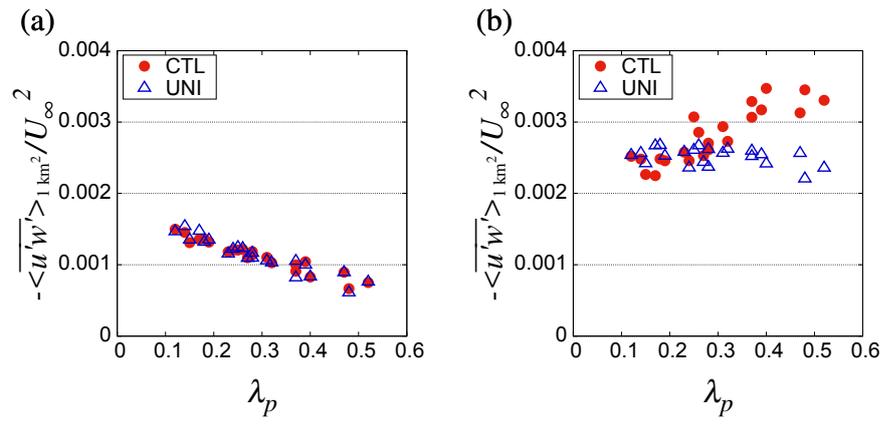


Fig. 11 Variations of Reynolds stress normalized by U_∞ with λ_p at heights (a) $z = 0.5h_{all}$ and (b) $z = 2.5h_{all}$.

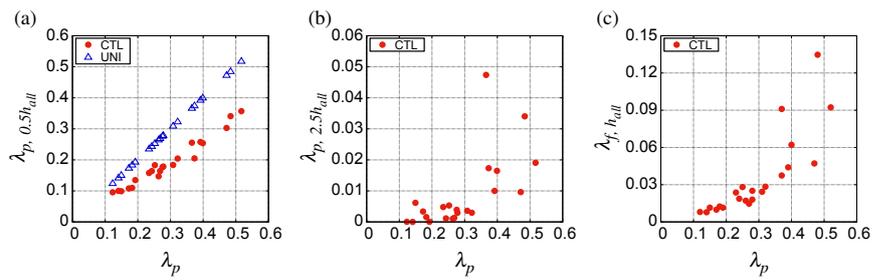


Fig. 12 Variations of λ_p calculated (a) at height $z = 0.5h_{all}$ ($\lambda_{p, 0.5h_{all}}$), (b) at height $z = 2.5h_{all}$ ($\lambda_{p, 2.5h_{all}}$), and (c) λ_f calculated at height $z = h_{all}$ ($\lambda_{f, h_{all}}$) with λ_p at the surface. Note that the scale of the vertical axis differs by panel.

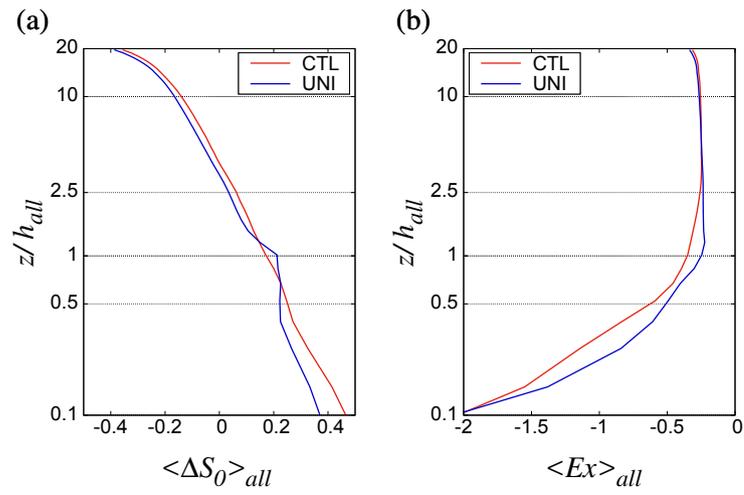


Fig. 13 Vertical profiles of (a) ΔS_0 and (b) Ex averaged spatially over the main region. Red and blue lines denote the results of the CTL and UNI experiments, respectively. The vertical axis is normalized by h_{all} . Note that a logarithmic scale is used for the vertical axis.

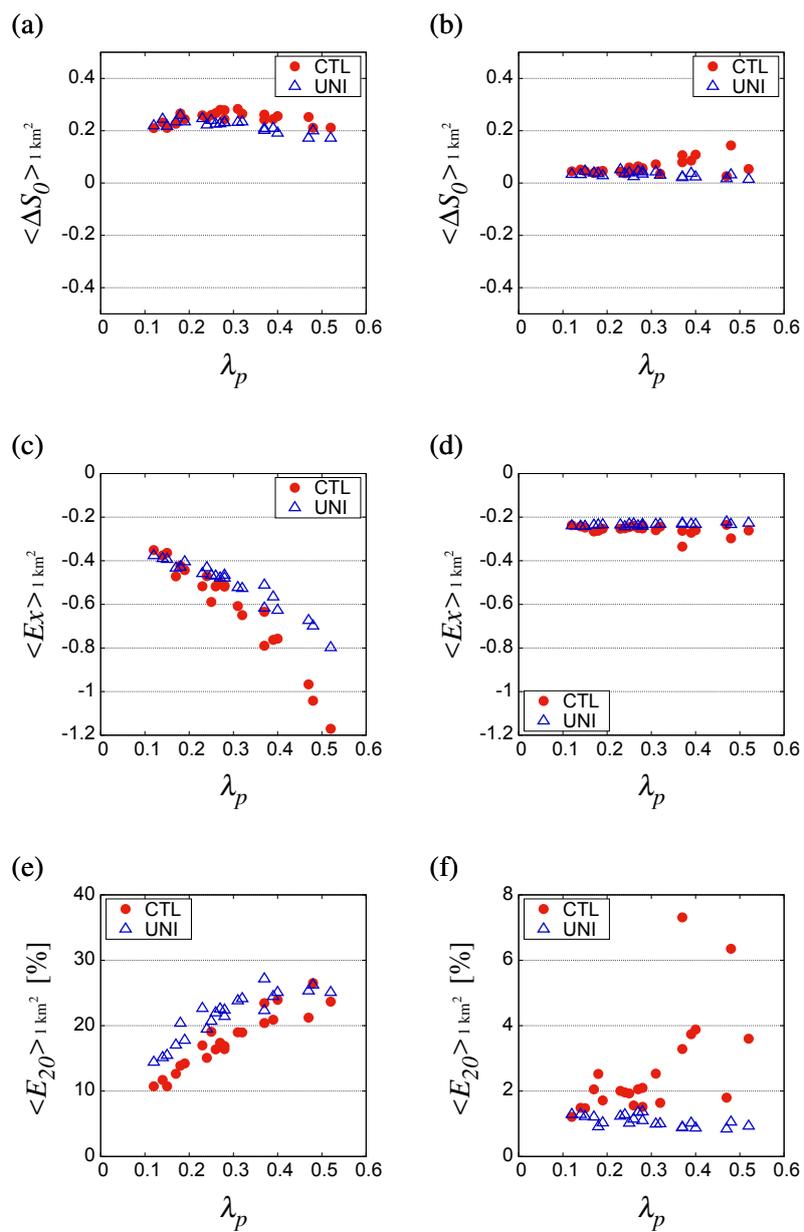


Fig. 14 Variations of ΔS_0 with λ_p at (a) heights $z = 0.5h_{all}$ and (b) $z = 2.5h_{all}$, Ex at (c) heights $z = 0.5h_{all}$ and (d) $z = 2.5h_{all}$, and E_{20} at (e) heights $z = 0.5h_{all}$ and (f) $z = 2.5h_{all}$.

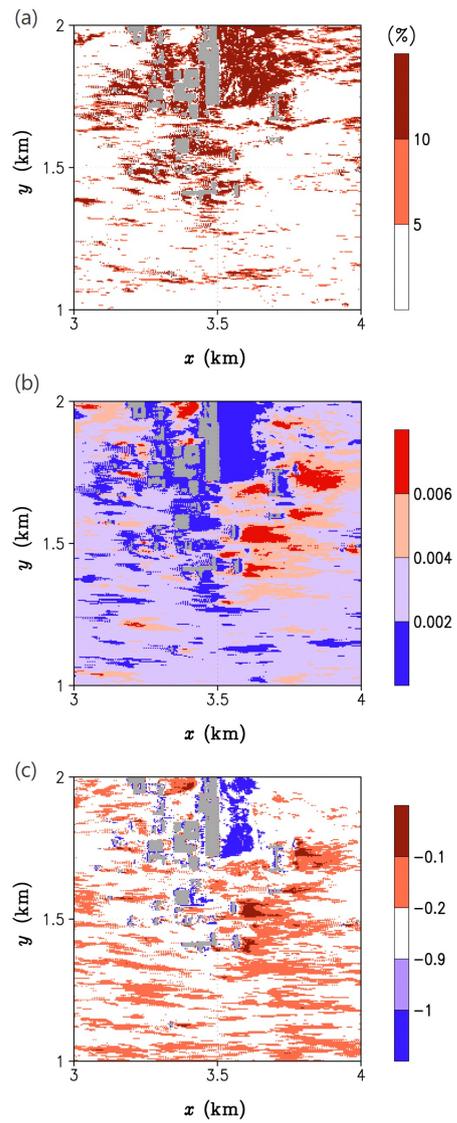


Fig. 15 Horizontal cross section of (a) E_{20} , (b) Reynolds stress normalized by U_∞ , and (c) Ex at height $z = 2.5h_{all}$ over a 1 km by 1 km area within the business district in the CTL experiment. The grey shading indicates buildings.