

Global gradient catastrophe and its unfolding in solutions of the Airy's model of shallow water waves.

Roberto Camassa

University of North Carolina at Chapel Hill, Carolina Center for Interdisciplinary Applied Mathematics, Department of Mathematics, Chapel Hill, NC 27599, USA

Executive summary

It is well known that under certain circumstances (see, e.g., [1, 2]), model equations such as the Airy's shallow water system,

$$\eta_t + (u\eta)_x = 0, \quad u_t + uu_x + \eta_x = 0, \quad x \in \mathbb{R}, \quad t \in \mathbb{R}^+, \quad (1.1)$$

here written in suitable nondimensional space-time (x, t) coordinates with η representing the water layer thickness and u the layer-averaged horizontal velocity, can capture some of the fundamental dynamics of the parent Euler equations for a free-surface, inviscid fluid under gravity, extending laterally to infinity and confined below by a flat bottom. Much of the model effectiveness at the qualitative and even quantitative level ultimately resides in the fact that the conservation laws of mass and horizontal momentum are captured either exactly (as in the first equation) or asymptotically in the limit of long waves, as in the second equation. With this in mind, we have recently found and studied exact solutions that can be used to shed some light on the peculiar features of the dynamics when the layer thickness vanishes (as set by initial data). In fact, it can be shown, even for the parent Euler equations, that “dry” points where $\eta(x, t) = 0$, so that the free surface touches the bottom of the fluid layer, tend to persist as long as the function $\eta(x, \cdot)$ remains sufficiently regular at these contact points. As a consequence, detachment of the free surface from the bottom can only happen through a loss of regularity of the solution.

Exact solutions

We consider the class of initial data

$$\eta = \gamma(t)x^2 + \mu(t), \quad u = \nu(t)x. \quad (1.2)$$

where the time evolution of the coefficients $\gamma(t)$, $\mu(t)$ and $\nu(t)$ is governed by the ODE's

$$\dot{\nu} + \nu^2 + 2\gamma = 0, \quad \dot{\gamma} + 3\nu\gamma = 0, \quad \dot{\mu} + \nu\mu = 0. \quad (1.3)$$

This system admits closed form solutions for generic initial data $\gamma(0) = \gamma_0$, $\mu(0) = \mu_0$ and $\nu(0) = \nu_0$. Of particular interest for our purposes is the case of contact $\eta(x, 0) = 0$ which occurs for $\mu_0 = 0$; then, the third equation in system (1.3) implies $\mu(t) = 0$ for all times $t > 0$ as long as the system's solution exists. This means that for as long as the parabolic form of $\eta(x, t)$ is maintained, the parabola's vertex at $x = 0$ stays at the bottom, or $\eta = 0$. Of course, this expression of the thickness of the water layer is unbounded for $|x| \rightarrow \infty$, and hence the solution (1.2) *per se* is unphysical. However, the parabola can be chopped at some elevation, $\eta = Q$ say, and spliced continuously with a constant background state $\eta(x, 0) = Q$ for $|x| > \sqrt{Q/\gamma_0}$, to form physically relevant initial data.

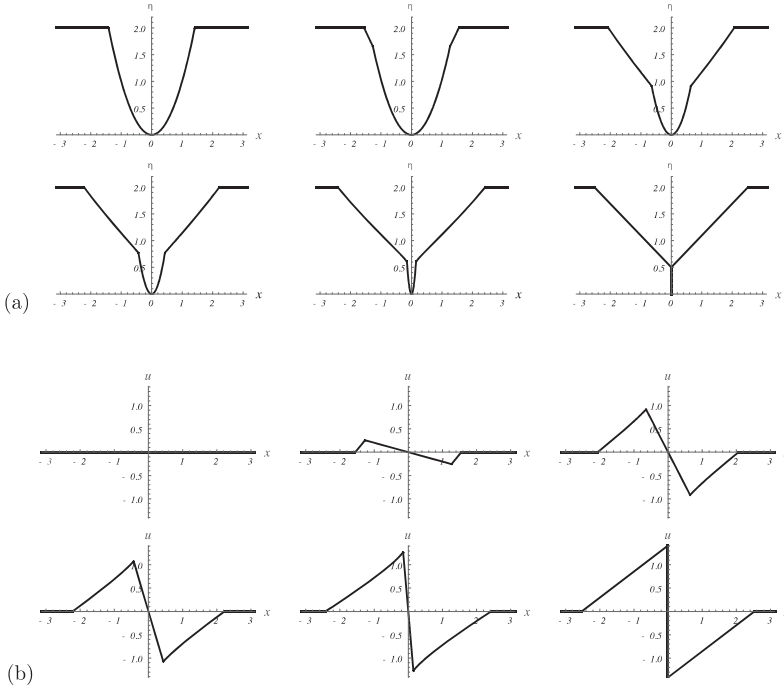


Figure 1: Layer thickness η (a), and velocity u (b) evolution for the Airy model when the initial surface has a dry (contact) point and zero initial velocities. The parameters are $Q = 2$, $\gamma_0 = 1$, $\mu_0 = 0$, yielding a collapse time $t_c = t_s \simeq 0.7854$. Time snapshots $t = 0, 0.01, 0.46, 0.57, 0.71, 0.78$.

It can be shown that the points $|x| = \sqrt{Q/\gamma_0}$, where the parabola joins the constant background, each split and evolve along distinct curves in the (x, t) -plane, which are among the characteristics of system (1.1), i.e., solutions of the ODE's

$$\dot{x}_{\pm} = \lambda_{\pm} \equiv u(x_{\pm}(t), t) \pm \sqrt{\eta(x_{\pm}(t), t)} \quad (1.4)$$

where the quantities (Riemann invariants)

$$R_{\pm} = u \pm 2\sqrt{\eta} \quad (1.5)$$

maintain their initial values. These curves emanating from the junction points bracket simple waves of system (1.1), that is, solutions for which η and u are functionally related. In the half domain $x \geq 0$ such simple wave solution, $\eta \equiv N(x, t)$ and $u \equiv V(x, t)$ say, can be expressed in closed form, albeit implicitly, through an auxiliary variable $\sigma_0(x, t) \in [1, \infty)$:

$$N(x, t) = \sigma_0 Q (\sqrt{\sigma_0} - \sqrt{\sigma_0 - 1})^2, \quad V(x, t) = 2\sqrt{N(x, t)} - 2\sqrt{Q}. \quad (1.6)$$

Here σ_0 as a function of x and t is defined by the solution of

$$x = \Lambda(\sigma_0) \left(t - \frac{\sqrt{\sigma_0 - 1} + \sigma_0 \arctan(\sqrt{\sigma_0 - 1})}{2\sigma_0 \sqrt{\gamma_0}} \right) + \frac{\sqrt{Q\sigma_0} - \sqrt{Q(\sigma_0 - 1)}}{\sigma_0 \sqrt{\gamma_0}}, \quad (1.7)$$

where we have used the shorthand notation Λ for the characteristic velocity λ_- ,

$$\Lambda(\sigma_0) = 3\sqrt{Q}\sigma_0\left(1 - \sqrt{1 - \frac{1}{\sigma_0} - \frac{2}{3\sigma_0}}\right), \quad (1.8)$$

for any given point (x, t) in the spatio-temporal half-plane $x > 0$ and $t > 0$. Symmetric and antisymmetric extensions, respectively for η and u , when $x < 0$ complete the exact solution form. A typical evolution of the initial data according of the exact solution constructed piecewise with the above results is depicted in figure 1.

As seen from the last panel in the figure, the solution develops a singularity in finite time, corresponding to $\gamma(t) \rightarrow \infty$, i.e., the parabolic core of the piecewise solution collapses onto a vertical segment of length $Q/4$. From the solution of the ODE's (1.3), and/or from the implicit expressions of the bracketing simple waves (1.7), (1.6) by finding the first time at which the partial derivative N_x becomes infinite, it can be shown that this verticality at the origin, or ‘‘global gradient catastrophe’’ (as the derivative of η and u become infinite not just at single points but for all points in an interval of the range of $\eta(x, \cdot)$), occurs at the time

$$t = t_c \equiv \frac{\pi}{4\sqrt{\gamma_0}}. \quad (1.9)$$

Evolution beyond the gradient catastrophe

As seen in the last panel of figure 1(a), the singularity of the dependent variables (η, u) at the time $t = t_c$ is that of a jump discontinuity for both fields. The segment connecting the free surface to the bottom at $x = 0$ represents colliding water masses at that point and time, and it is natural to remove it to obtain a connected domain for the water layer at times $t = t_c^+$. However, the discontinuity from the values \sqrt{Q} (left) to $-\sqrt{Q}$ (right) in the velocity u would remain at $x = 0$, and it is natural to expect that this jump would result in the instantaneous rising of the fluid in a neighbourhood of this location. This can be determined precisely by studying the initial value problem for system (1.1) corresponding to initial data obtained by the ‘‘wings’’ solutions $N(x, t_c)$ and $V(x, t_c)$ in (1.6), appropriately reflected across the origin.

Looking at the limit $x \rightarrow 0^+$ of $N(x, t_c)$, it can be seen that this function is not differentiable at the origin; more precisely

$$N(x, t_c) \sim \frac{Q}{4} + \frac{3^{2/3}\gamma_0^{1/3}Q^{2/3}}{8}x^{2/3} + o(x^{2/3}) \quad \text{as } x \rightarrow 0. \quad (1.10)$$

The consequences of this branch point singularity on the evolution of the ‘‘new’’ initial data after gradient catastrophe are mostly of technical nature, and it is worth studying a class of data that removes the fractional power obstacle, yet captures the main features of the time advancement past the $t = t_c$.

Consider the initial data obtained by splicing together two ‘‘Stoker’’ simple waves (see [1]) crossing at $x = 0$ and $\eta(0, 0) = \eta_0$,

$$N_S(x, t) = \begin{cases} Q, & x \geq x_Q \\ \frac{1}{9} \left(\frac{x - x_d}{t + t_c} + 2\sqrt{Q} \right)^2, & 0 < x < x_Q \\ \frac{1}{9} \left(-\frac{x + x_d}{t + t_c} + 2\sqrt{Q} \right)^2, & -x_Q < x < 0 \\ Q, & x \leq -x_Q \end{cases}, \quad V_S(x, t) = \begin{cases} 0, & x \geq x_Q \\ \frac{2}{3} \frac{x - x_d}{t + t_c} - \frac{2}{3} \sqrt{Q}, & 0 < x < x_Q \\ \frac{2}{3} \frac{x + x_d}{t + t_c} + \frac{2}{3} \sqrt{Q}, & -x_Q < x < 0 \\ 0, & x \leq -x_Q \end{cases}. \quad (1.11)$$

With

$$x_d = \frac{1}{4}\sqrt{\frac{3Q}{g_0}}, \quad t_c = \frac{1}{2}\sqrt{\frac{3}{g_0}}, \quad x_Q = \sqrt{Q}t + \frac{3}{4}\sqrt{\frac{3Q}{g_0}}. \quad (1.12)$$

one can obtain a configuration that closely resembles that of the pair $N(x, t_c), V(x, t_c)$ in (1.6) sketched in figure 2, dashed curves. As depicted in this figure, the evolution out of these initial data for times $t > 0$

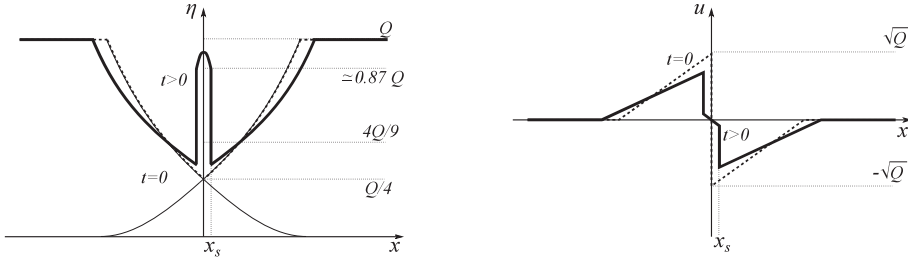


Figure 2: Schematics of the initial condition (dash) and evolution (solid) with shock development at short times $t > t_c$ for the Airy solutions (1.11). The thin segments of parabolae are removed from the initial data and the shocks develops from the initial discontinuity in the velocity (right panel).

(corresponding to times $t > t_c$ for the original case of (1.2)) involves the generation of two symmetrically placed shocks in both the η and u fields, moving away from the origin, with an intermediate region between the shock corresponding to fluid slowly filling the initial time “hole” at $x = 0$, $\eta = Q/4$. These shocks move over the background given by the evolution of the Stoker waves (1.11) for $t > 0$.

Let $x_s(t)$ be the position of the right-going shock. We introduce the “unfolding” coordinates (ξ, τ) to set the shock positions at fixed locations in time, e.g., $\xi = \pm 1$, which can be achieved by

$$\xi \equiv \frac{x}{x_s(t)}, \quad \tau \equiv \log(x_s(t)), \quad (1.13)$$

so that the Jacobian of this transformation reads

$$\partial_x = \frac{1}{x_s} \partial_\xi, \quad \partial_t = \frac{\dot{x}_s}{x_s} (\partial_\tau - \xi \partial_\xi). \quad (1.14)$$

With this mapping, the Airy’s system (1.1) assumes the form (with a little abuse of notation by maintaining the same symbols for dependent variables)

$$\frac{\partial \eta}{\partial \tau} - \xi \frac{\partial \eta}{\partial \xi} + \frac{1}{\dot{x}_s} \frac{\partial}{\partial \xi} (\eta u) = 0, \quad \frac{\partial u}{\partial \tau} - \xi \frac{\partial u}{\partial \xi} + \frac{1}{\dot{x}_s} \frac{\partial}{\partial \xi} \left(\frac{u^2}{2} + \eta \right) = 0. \quad (1.15)$$

Here \dot{x}_s is a placeholder for the expression that couples the evolution equation for the physical time t to the new evolution variable τ . The shock position evolves according to the equation that defines \dot{x}_s in terms of the jump amplitudes $[\eta]$ and $[u]$:

$$\frac{d(e^\tau)}{dt} = \dot{x}_s = \frac{[\eta u]}{[\eta]} = \frac{N(x_s(t), t) V(x_s(t), t) - \eta(1, \tau(t)) u(1, \tau(t))}{N(x_s(t), t) - \eta(1, \tau(t))}, \quad x_s(0) = 0, \quad (1.16)$$

so that, in the new variables,

$$\frac{dt}{d\tau} = \frac{e^\tau (N(e^\tau, t(\tau)) - \eta(1, \tau))}{N(e^\tau, t(\tau)) V(e^\tau, t(\tau)) - \eta(1, \tau) u(1, \tau)}, \quad t(\tau) \rightarrow 0 \quad \text{as } \tau \rightarrow -\infty. \quad (1.17)$$

Here and in the following, we have suppressed the subscript \cdot_S for the Stoker simple-waves, and adopted the usual square-bracket notation for jumps defined as the difference between values \pm at the right and left of the shock location, respectively (see e.g., [2]). From these expressions, the system governing the evolution of the “inner” solution between shocks consists of this ordinary differential equation together with the partial

differential equations (1.15), to be solved within the strip $\xi \in [0, 1]$ (by symmetry only half the ξ domain $[-1, 1]$ may be used) subject to the boundary conditions

$$u(0, \tau) = 0, \quad u(1, \tau) = V(e^\tau, t(\tau)) + \sqrt{\frac{(N(e^\tau, t(\tau)) - \eta(1, \tau))^2 (N(e^\tau, t(\tau)) + \eta(1, \tau))}{2N(e^\tau, t(\tau))\eta(1, \tau)}}. \quad (1.18)$$

The first equality is a consequence of the antisymmetry of the velocity, $u(\xi, \tau) = -u(-\xi, \tau)$. The second relation expresses, in terms of the new independent variables, the consistency condition for the shock speed

$$\dot{x}_s = \frac{[\eta u^2 + \eta^2/2]}{[\eta u]}, \quad (1.19)$$

which can be manipulated to

$$[u]^2 = \frac{[\eta]^2(\eta_+ + \eta_-)}{2\eta_+ \eta_-}, \quad (1.20)$$

whence the second relation in (1.18) follows.

The boundary-value problem (1.18) for the evolution equations (1.15) and (1.17), by depending on the unknown η at the boundary, $\eta(1, \tau)$, is reminiscent in its structure of the classical (irrotational) water-wave problem, wherein the unknowns, the free surface location and the velocity potential along it, determine and in turn are determined by the solution of a PDE (for water wave problem, the Laplace equation for the velocity potential in the fluid domain). Just as in that case, an additional equation has to be provided at the boundary, which has its analog in (1.17). Similarly to the water-wave problem, the resulting structure is highly nonlinear and hence hardly amenable to closed form solutions. To make progress, observe that the “initial” data as $\tau \rightarrow -\infty$, i.e., $t = 0$, are given by $u(\xi, -\infty) = 0$ and $\eta(\xi, -\infty) = Q^*$, where Q^* is the constant solution of the cubic equation

$$-2\sqrt{Q} + 2\sqrt{N(0, 0)} + \sqrt{\frac{(N(0, 0) - Q^*)^2 (N(0, 0) + Q^*)}{2N(0, 0)Q^*}} = 0, \quad (1.21)$$

subject to the condition $Q^* > N(0, 0)$. Thus, the initial time evolution can be followed by the linearization of system (1.15)-(1.17) around $\eta = Q^*$ and $u = 0$. This approach is mostly straightforward, though the details are bit involved, and will be reported elsewhere. Suffices to say that the initial evolution as $\tau \rightarrow -\infty$ (and so $t \rightarrow 0^+$) is asymptotic to

$$\eta(\xi, \tau) = Q^* + \frac{1}{2}(F(e^\tau \xi - \Phi(\tau)) + F(-e^\tau \xi - \Phi(\tau))), \quad u(\xi, \tau) = \frac{1}{2\sqrt{Q^*}}(F(e^\tau \xi - \Phi(\tau)) - F(-e^\tau \xi - \Phi(\tau))), \quad (1.22)$$

for any function $F(\cdot)$ of sufficient regularity, with the function Φ defined by

$$\Phi(\tau) \equiv \sqrt{Q^*} \int_{-\infty}^{\tau} e^{\tau'} \phi_0(\tau') d\tau', \quad (1.23)$$

where the integrand $\phi_0(\tau)$ is

$$\phi_0(\tau) = \frac{N(e^\tau, t(\tau)) - Q^*}{N(e^\tau, t(\tau))V(e^\tau, t(\tau))}. \quad (1.24)$$

Here s_0 is the initial shock speed \dot{x}_s at $t = 0^+$, and the initial conditions on the system's solution (η, u) as $\tau \rightarrow -\infty$ require $F(0) = 0$. Substitution of these expressions evaluated at $\xi = 1$ into the linearized version of the boundary condition (1.18) leads to a functional equation for F , coupled to the evolution (linearized) equation (1.17) for $t(\tau)$. By assuming sufficient regularity for F , the functional equation can be solved approximately by Taylor series. The result to second order is

$$\eta(\xi, \tau) \sim Q^* - F'(0)\Phi_0 e^\tau + \left(-F'(0)\Phi_1 + \frac{1}{2}F''(0)(\xi^2 + \Phi_0^2)\right) e^{2\tau}, \quad (1.25)$$

$$u(\xi, \tau) \sim \frac{F'(0)}{\sqrt{Q^*}} \xi e^\tau - \frac{F''(0)}{\sqrt{Q^*}} \Phi_0 \xi e^{2\tau}, \quad (1.26)$$

where the Taylor coefficients $F'(0)$, $F''(0)$ and those for the asymptotic expansion of $\Phi(\tau) \sim \Phi_0 e^\tau + \Phi_1 e^{2\tau}$ are, respectively,

$$F'(0) \simeq -0.22215\sqrt{g_0 Q}, \quad F''(0) \simeq -0.58487 g_0, \quad (1.27)$$

and

$$\Phi_0 = \frac{\sqrt{Q^*(4Q^* - Q)}}{Q^{3/2}} \simeq 2.33087, \quad \Phi_1 = \frac{\sqrt{g_0 Q^*} (Q + 4Q^*)^2}{3\sqrt{3} Q^3} \simeq -3.6325\sqrt{\frac{g_0}{Q}}, \quad (1.28)$$

with the parametrization (1.12) for the initial data. The asymptotics for the shock location x_s is similarly computed as $\tau \rightarrow -\infty$.

Returning to the original physical variables $\eta(x, t)$, $u(x, t)$, the asymptotic expressions (1.25),(1.26) lead to

$$\eta(x, t) \sim Q^* - F'(0)\Phi_0 s_0 t - \left(F'(0)\Phi_0 s_1 + \left(F'(0)\Phi_1 - \frac{1}{2} F''(0)\Phi_0^2 \right) s_0^2 \right) t^2 + \frac{1}{2} F''(0)x^2, \quad (1.29)$$

$$u(x, t) \sim \frac{x}{\sqrt{Q^*}} (F'(0) - F''(0)\Phi_0 (s_0 t + s_1 t^2)), \quad (1.30)$$

$$x_s(t) \sim s_0 t + s_1 t^2, \quad \text{with } s_0 \simeq 0.400969\sqrt{Q}, \quad s_1 \simeq 0.11703\sqrt{g_0 Q}, \quad (1.31)$$

as $t \rightarrow 0^+$. Note that the asymptotic validity of these expressions is established within the unfolding variable formulation, and hence it is not necessarily maintained after the mapping (1.13), unless this is also consistently expanded with the asymptotics. Nonetheless, the above expression for the η and u field reveal that the layer thickness η jumps, at time $t = 0^+$, to about 87% of the background rest thickness Q , and grows linearly in time while maintaining a local parabolic shape, fixed in time at this order. Similarly, the local velocity u jumps from a discontinuity of amplitude $2\sqrt{Q}$ at $x = 0$ to a continuous linear profile between shocks, with negative slope which evolves linearly in time.

Of course, further progress is most easily achieved by resorting to numerical methods. However, we remark that standard WENO algorithms we have implemented are not able to capture the above details at short times, and one has to resort to the unfolding variable system which can be integrated numerically with spectrally accurate codes. These, once validated on the analytical results, can provide information on the time evolution well past the initial times. Numerical tools are even more challenged by the full case of initial data (1.2) when going beyond the gradient catastrophe, since this case poses additional difficulties due to the singular behavior of $N(x, t_c)$ at the origin. The main details of this and other cases will be reported elsewhere.

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