3 Density Functional Theory Estimation of Isotope Fractionation of Fe, Ni, C
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- 4 Among Species Relevant to Geochemical and Biological Environments
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Abstract

20 This paper reports the values of reduced partition function ratios (as 1000 ln β) for Fe, 21 Ni, Cu, and Zn bound to a number of inorganic and organic ligands. We used Density Functional Techniques to update the existing data and calculate $\ln \beta$ for new ligands. 22 23 This work allows for the mass-dependent isotope fractionation to be predicted for 24 various inorganic (hydrated cation, hydroxide, chloride, sulfate, sulfide, phosphate) and 25 organic (citrate, amino acid) complexes of Fe, Ni, Cu, and Zn. Isotope fractionation 26 among coexisting complexes of these metals was evaluated from the $\ln \beta$ values in a 27 variety of geochemical and biological environments. The results provide a framework for interpretation of isotope fractionation observed in seawater and chemical sediments, 28 29 in the roots and aerial parts of plants, and among the organs and body fluids of 30 mammals.

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1. INTRODUCTION

Stable isotope geochemistry uses the isotopic compositions of naturally occurring elements to shed light on the origins of natural rocks, minerals, and fluids at low and moderate temperatures (< 773 K). Non-traditional stable isotopic systems, e.g., Fe, Ni, Cu, and Zn (Albarède, 2004; Beard and Johnson, 2004; Johnson et al., 2004a,b Fujii et al., 2011a), as opposed to those of the widely used elements, H, C, N, O, and S, give important information on the environment of the modern and ancient Earth and on the origin and evolution of life.

41 Interpretation of the isotope composition of natural samples is elicited from 42 experimental data obtained by modern mass spectrometry, notably multiple-collection 43 inductively coupled plasma mass spectrometry (MC-ICP-MS). Equilibrium isotope 44 fractionation arises during exchange of isotopes among different chemical species. It is 45 possible to compute the electronic state of isotopologues (molecular entities that differ 46 only by the nature of their isotopes). This means that, even for ligand exchange or 47 electron exchange (redox), the isotope fractionation factor of equilibrium reactions can 48 be estimated from the energy differences in the electronic states of the isotopologue 49 reactants and products. Catalysis changes the activation energy of reactions, but does 50 not change the electronic states of the final reactants and products. Kinetic effects can 51 be considered as paths towards equilibrium. If the reaction rates of forward and 52 backward reactions are not too different, the isotope effect in the process may be 53 approximated by equilibrium fractionation. If they are very different from each other, 54 the energy difference between the transitional isotopologues becomes the dominant 55 factor of isotope fractionation (Bigeleisen and Wolfsberg., 1958).

56 This study reports the values of $\ln \beta$, the logarithm of reduced partition function 57 ratios (Bigeleisen and Mayer, 1947), computed by ab initio methods for a variety of 58 chemical species of Fe, Ni, Cu, and Zn relevant to geochemistry and biochemistry. We 59 discuss the computed values of the isotopologue partition function ratio $\ln \beta$ in the 60 context of pH-dependent speciation plots, which enabled us to determine the isotope 61 effect for coexisting species in a multi-species system. As an application to 62 geochemistry, the isotope fractionation of Fe, Ni, Cu, and Zn under the conditions 63 prevailing in seawater was estimated from the $\ln \beta$ values and related to natural data obtained in previous geochemical studies. As an application to biogeochemistry, the 64 65 isotope fractionation for some compounds relevant to soil-plant systems was estimated. 66 In view of the isotope fractionation of Fe, Cu, and Zn observed among mammal organs 67 and body fluids (Albarede et al., 2011; Albarède, 2013; Balter et al., 2013; Moynier et 68 al., 2013a), $\ln \beta$ values of Cu- and Zn-amino acid complexes were also estimated.

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- 70

2. METHODS

71 Orbital geometries and vibrational frequencies of Fe, Ni, Cu, and Zn species were 72 computed using the density functional theory (DFT) implemented by the Gaussian09 73 code (Frisch et al., 2009; Dannington et al., 2009). The DFT method employed here is a hybrid density functional consisting of Becke's three-parameter non-local hybrid 74 75 exchange potential (B3) (Becke, 1993) with Lee-Yang-and Parr (LYP) (Lee et al., 1988) 76 non-local functionals. The 6-311+G(d,p) basis set, which is an all-electron basis set, 77 was used for H, C, N, O, P, S, Cl, Fe, Ni, Cu, and Zn. Molecules were modeled without 78 any forced symmetry. An "ultrafine" numerical integration grid was used and the SCF (self-consistent field) convergence criterion was set to 10^{-8} or 10^{-9} . We tested whether 79

software version and database may have introduced any bias. To this effect, we recalculated with Gaussian09 the Zn and Ni fractionation factors of the same species as those reported by Fujii et al. (2010, 2011a,b) who used Gaussian03 and B3LYP/6-311+G(d,p). The tests show that the results of the previous studies are well reproduced and that changes in ln β are marginal.

The coordination numbers were set to 6 for Fe(II), Fe(III), Ni(II), and Zn(II), to 5 for Cu(II), and to 2 for Cu(I). The validity of optimized structures of model molecules was checked by using spectroscopic data and thermochemical stoichiometry from the literature (Fujii et al., 2006, 2010, 2011a,b, 2013). Calculations were performed for single cluster model molecules. The effect of ligands beyond the first coordination spheres was not considered. The bias associated with this assumption is discussed in Fujii et al. (2010, 2011b).

For the first (3*d*) transition elements, the effect of intramolecular vibrations (Bigeleisen and Mayer, 1947; Urey, 1947) is much stronger than the nuclear field shift effect (Bigeleisen, 1996; Nomura et al., 1996; Fujii et al., 2010, 2011a, 2013), which therefore will not be considered further. The isotope enrichment factor accordingly was evaluated from the reduced partition function ratio (s/s')*f* (Bigeleisen and Mayer, 1947), also noted β , such as,

$$\ln \frac{s}{s'} f = \sum \left[\ln b(u_i') - \ln b(u_i) \right]$$
⁽¹⁾

98 where

$$\ln b(u_i) = -\ln u_i + \frac{u_i}{2} + \ln(1 - e^{-u_i})$$
⁽²⁾

99 and

$$u_i = \frac{hv_i}{kT} \tag{3}$$

In the latter expression, v stands for vibrational frequency, *s* for the symmetry number of the considered compound, *h* for the Plank constant, *k* for the Boltzmann constant, and *T* for the absolute temperature. The subscript *i* denotes the *i*th normal mode of molecular vibration, and primed variables refer to the light isotopologue. The isotope enrichment factor due to molecular vibrations can be evaluated from the frequencies v_i summed over all the different normal modes.

107 The relative enrichment of a specific compound in a multi-species system was 108 estimated by using formation constants and the ln β values of each coexisting species. 109 The standard δ notation for an element M with masses *m*' (light isotope, A') and *m* 110 (heavy isotope, A) is defined as,

$$\delta^{m}M = \left(\begin{array}{c} ([A]/[A'])_{sample} \\ ([A]/[A'])_{reference} \end{array} - 1\right) \times 1000$$
(4)

111

112 The isotope fractionation between two different species X and Y is defined as,

$$\Delta^{m}M = \delta^{m}M_{Y} - \delta^{m}M_{X}$$
⁽⁵⁾

The detailed methods used to evaluate isotope fractionation among coexisting species
was described by Fujii et al. (2010, 2011a, 2013) and are reproduced in *APPENDIX A*.
The absolute ln β values and their relative scale may vary with the calculation
method [e.g., DFT, HF (Hartree-Fock), MPn (Møller Plesset)], basis set, cluster size,
and solvation model (Fujii et al., 2010, 2011a,b; Rustad et al., 2010; Hill et al., 2014).
How different options affect the results and their overall accuracy is discussed in
Appendix B.

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3. RESULTS AND DISCUSSION

122 **3.1. Isotope fractionation among species in modern seawater**

We here reproduce speciation of Cu(II) and Zn(II) inorganic species in seawater as calculated by Zirino and Yamamoto (1972) for 19% chlorinity at 298 K and 1 atm total pressure. The model includes activity coefficients for monovalent, divalent, and neutral species. Concentrations of free HCO_3^- and CO_3^{2-} were estimated as a function of pH. This model was used to estimate isotope fractionation of Fe, Ni, Cu, and Zn among inorganic ligands present in seawater.

We expand the pH range from 7 to 9 in Zirino and Yamamoto's (1972) original Cu(II) speciation model to a range of 6 to 9 (Fig. 1a). The concentrations of HCO_3^- and CO_3^{2-} at pH < 7 were estimated by extrapolating the data reported using Zeebe and Wolf-Gladrow (2001),

$$[HCO_{3}^{-}] = \frac{DIC}{\left(1 + \frac{[H^{+}]}{K_{1}^{*}} + \frac{K_{2}^{*}}{[H^{+}]}\right)}$$
(6)

133 and

$$[CO_{3}^{2-}] = \frac{DIC}{\left(1 + \frac{[H^{+}]}{K_{2}^{*}} + \frac{[H^{+}]^{2}}{K_{1}^{*}K_{2}^{*}}\right)}$$
(7)

where DIC represents the total concentration of dissolved inorganic carbon, $[CO_2] +$ [HCO₃⁻] + $[CO_3^{2-}]$. K_1^* and K_2^* are the equilibrium constants, that is, $K_1^* =$ [HCO₃⁻][H⁺]/[CO₂] and $K_2^* = [CO_3^{2-}][H⁺]/[HCO_3^-]$. These equations were employed in the extrapolation by setting DIC, K_1^* , and K_2^* as parameters. 138 The ln β values of various Cu species were computed by Fujii et al. (2013) for 139 fourfold and fivefold coordinations and agree with the results of Sherman (2013). For example, $\ln \beta$ of Cu(H₂O)₅²⁺ (298 K) is 4.546‰ (Fujii et al., 2013) compared with 140 4.55% for Sherman (2013). The δ^{65} Cu value of each species was calculated as functions 141 of pH (see Appendix A). As shown in Figs. 1a and 1b, at pH ~6, ⁶⁵Cu is enriched in 142 $CuSO_4$ and $CuHCO_3^+$, while ⁶³Cu is enriched in Cu^{2+} and $CuCl^+$. With increasing pH, 143 $Cu(OH)_2$ and $CuCO_3$ become the prevalent species. At a typical pH of seawater, 8.22 144 (Macleod et al., 1994), isotope fractionation favors ⁶³Cu in CuCO₃ and ⁶⁵Cu in Cu(OH)₂. 145 Figure 1c is a summary of δ^{65} Cu for the different Cu species present. At low pH, 146 positive δ^{65} Cu are found in Cu sulfates and carbonates, while δ^{65} Cu are negative in 147 hydrated Cu²⁺ and chlorides (see Figs. 1a and 1c). At high pH, δ^{65} Cu is positive in Cu 148 149 hydroxides and negative in carbonates.

150 The distribution of Zn(II) species in seawater at 298 K and 1 atm as functions of 151 pH was also taken from Zirino and Yamamoto (1972) (Fig. 2a). Electronic structures of hydrated Zn^{2+} , chloride, carbonate, hydroxide, and sulfate were computed, and fivefold 152 153 coordination was found to provide a stable configuration for carbonate species. The 154 optimized structures in Cartesian coordinates are given in Table S1 of the electronic supplement. The ln β values obtained are reported in Table 1. The δ^{66} Zn values for the 155 ⁶⁶Zn/⁶⁴Zn ratio were calculated for each relevant species as a function of pH. A 156 dominant species of free Zn^{2+} shows small isotope fractionation (Fig. 2). ZnSO₄ is 157 enriched in ⁶⁶Zn, whereas Zn chlorides are enriched in ⁶⁴Zn with a Δ^{66} Zn ~0.5‰ being 158 expected between Zn sulfate and chloride (Fig. 2c). With increasing pH, Zn(OH)₂ and 159 $ZnCO_3$ become the dominant species. Small amounts of free Zn^{2+} and $ZnCl^+$ still exist 160

161 at pH = 8.2. In seawater, a fractionation Δ^{66} Zn of ~1‰ is expected between Zn 162 carbonate and chloride (Fig. 2c). Zn hydroxides and sulfates do not play an important 163 role for Zn isotope fractionation for pH ≥ 8.2 (Fig. 2c).

164 A speciation model similar to that of Zirinno and Yamamoto (1972) likewise was 165 applied to Ni(II) and Fe(II) in seawater at 298 K and 1 atm. The speciation diagram of 166 Ni(II) is shown in Fig. 3a. The stability constants were taken from Byrne et al. (1988) (sulfate), Foulliac and Criaud (1984) (carbonates), and Turner et al. (1981) (other 167 168 species). The optimized structures in Cartesian coordinates are given in Table S2 of the 169 electronic supplement. The $\ln \beta$ values obtained are reported in Table 2. The maximum 170 difference in ln β values among the species shown in Fig. 3b is ~0.8‰ (298 K). The values of δ^{60} Ni for the isotope pair (60 Ni/ 58 Ni) are shown for each species Figs. 3b and 171 3c. At low pH, the major Ni species are hydrated Ni²⁺ and Ni chlorides, and only little 172 fractionation with δ^{60} Ni of ~0.1‰ is expected. At higher pH, 60 Ni is enriched in Ni 173 174 carbonates, which would cause +0.5% fractionation compared with that of chlorides.

175 The distribution of Fe(II) species in seawater at 298 K and 1 atm as a function of pH is shown in Fig. 4a. The stability constant for FeHCO₃⁺ was taken from Foulliac and 176 177 Criaud (1984) and the rest from Turner et al. (1981). The optimized structures of Fe(II) 178 complexes are reported in Table S3 of the electronic supplement. For example, the calculation reproduces the hexahydrated structure of $Fe(H_2O)_6^{2+}$ with the $Fe^{2+}-H_2O$ 179 180 distance (2.14-2.18 Å) of 2.12 Å determined by X-ray diffraction (XRD) (Magini et al., 181 1988). The ln β values obtained (Table 3) compare well with *ab initio* studies of Fe 182 complexes (Ottonello and Zuccolini, 2009; Hill et al., 2010; Rustad et al., 2010, Moynier et al., 2013b, and references therein). As shown in Figs. 4b and 4c, ⁵⁴Fe is 183 184 enriched in FeCl⁺. Dependence of ln β on temperature is shown in Figure 5a. In the

185 lower pH region, 56 Fe is enriched in FeSO₄ and FeHCO₃⁺. With increasing pH, FeCO₃ 186 is becoming the prevalent species and is enriched in 56 Fe.

187 Because of the strong hydrolysis taking place in seawater, Fe(III) forms Fe(OH)₃ 188 (Byrne and Kester, 1976). The values of $\ln \beta$ were calculated accordingly for Fe(III) 189 species as well. The optimized structures in Cartesian coordinates are given in Table S4 of the electronic supplement. Again, the hexahydrated structure of $Fe(H_2O)_6^{3+}$ with the 190 191 Fe³⁺-H₂O distance of 2.05 Å determined by XRD (Magini et al., 1988) is well 192 reproduced (2.06 Å). The ln β values obtained are reported in Table 4. Dependence of ln β on temperature is shown in Figure 5b. ln β of hydrated Fe³⁺ and Fe(III) chlorides 193 194 calculated by *ab initio* methods has been reported previously (Fujii et al., 2006; Hill and 195 Schauble, 2009) and the present calculations agree well with literature values. The $\ln \beta$ 196 values of ferric hydroxide complexes are characteristically large. Among the hydroxides, 197 lower order of hydrolytic species shows smaller $\ln \beta$. This suggests that the dissociation of Fe(OH)₃ with increasing acidity enriches ⁵⁴Fe in both Fe(OH)²⁺ and Fe(OH)₂⁺. 198

Under reducing conditions, redox reactions among Fe(II) and Fe(III) species must be considered. For example, the [Fe(II)]/[Fe(III)] ratio of ~4 has been advocated for the Baltic Sea (Kononets et al., 2002). Distribution of Fe(II) and Fe(III) species in seawater at 298 K, 1 atm, and pH = 8.2 was estimated by the method of Zirinno and Yamamoto (1972) for [Fe(II)]/[Fe(III)] = 4. The resulting δ^{56} Fe values are shown in Fig. S1 of the electronic supplement. ⁵⁶Fe is enriched in Fe(OH)₃ and ⁵⁴Fe in free Fe²⁺. The isotope fractionation factor Δ^{56} Fe (Fe(OH)₃ - Fe²⁺) between the two species is ~4.3‰.

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207 **3.2. Euxinic seawater**

The role of sulfides is central to a broad range of geological scenarios. The status of sulfur in ancient oceans in particular is still an outstanding issue (Canfield, 1998). Hydrothermal vent solutions discharging either at mid-ocean ridges (Edmond et al., 1979) or along subduction zones (Mottl et al., 2004) comprise additional environments dominated by sulfides. In a previous study (Fujii et al., 2011b), isotope fractionation among the different Zn sulfide species present in geological fluids between 298 and 573 K was evaluated. Here, we extend this study to Fe, Ni, and Cu.

Speciation of Fe(II) in seawater-like solutions containing hydrogen sulfides has been previously investigated (Rickard and Luther, 2007; Wu et al., 2012). It was pointed out that Fe(II) sulfides are dominant at pH typical of seawater. The speciation diagram of Fe(II) species in the seawater-like matrix at pH = 6-9 was reproduced from Rickard and Luther (2007) (Fig. 6a). At pH = 8.2, the aqueous FeS is prevalent. At lower pH, FeHCO₃⁺ becomes a major Fe(II) species.

221 A characteristic feature of the speciation diagram (Fig. 6a) is the presence of the 222 higher order Fe(II) hydroxide, Fe(OH)₃. In the pH-Eh diagrams, Rickard and Luther (2007) treated FeOH⁺ and Fe(OH)₂ as fougerite, a hydrotalcite mineral of Fe(II) and 223 Fe(III). This may be the reason for the lack of $FeOH^+$ and $Fe(OH)_2$ in the speciation 224 225 diagram. The formation of $Fe(OH)_3^-$ was thermochemically confirmed by Baes and 226 Mesmer (1976), but there are no information on its stereochemical structure. 227 Hydroxylation of Fe(II) and condensation of its hydroxides form the ferrous oxide 228 (Jolivat et al., 2004). $Fe(OH)_2$ has the brucite structure, which is a layer structure with 229 hydroxyl (OH⁻) groups in hexagonal close packing: each Fe(II) is octahedrally 230 coordinated to six OH⁻ groups and these octahedra share edges to form the layers (Zigan and Rothbauer, 1967). $Fe(OH)_2(H_2O)_4$ corresponds to monomeric aqueous $Fe(OH)_2$ and 231

232 is treated as a *trans*-structure (Table S3 and Table 3). By reference to Fe(OH)₂(H₂O)₄, 233 we attempted to treat the structure of $Fe(OH)_3^{-1}$ as $Fe(OH)_3(H_2O)_3^{-1}$ but the species was 234 found to be unstable in our calculation as the three H₂O molecules moved away from its first coordination sphere. Since the dimer $Fe_2(OH)_2(H_2O)_8^{2+}$ of $Fe(OH)^+$ is considered a 235 236 structural unit of the Fe(OH)₂ formation (Jolivat et al., 2004), its hydrolytic species $Fe_2(OH)_6(H_2O)_4^{2-}$ may actually be treated as a dimer of $Fe(OH)_3^{-}$. The electronic states 237 238 of Fe₂ (Irigoras et al., 2003) was referred for the calculation of dimeric species of Fe(II). 239 The results show that Fe(II) possesses a fluttered square plane structure with four OH⁻ ions, and H₂O molecules are bound to the OH⁻ moiety of $Fe_2(OH)_6^{2-}$ via hydrogen bonds. 240 We calculated ln β for Fe₂(OH)₆²⁻ (Table S3 and Table 3), and found that it was close to 241 that of $Fe(OH)_2(H_2O)_4$. We therefore opted for using the value of ln β of $Fe(OH)_2$ for 242 243 Fe(OH)₃.

hydrated Fe(II) mono-hydrogensufide $FeHS(H_2O)_5^+$ would 244 Iron in be sixfold-coordinated. Rickard and Luther (2006) suggested that the structure of 245 246 bis-hydrogensulfide for six-coordination transition metals is similar to that of 247 trans-Mn(HS)₂(H₂O)₄. We therefore adopted the structure trans-Fe(HS)₂(H₂O)₄ in the 248 present calculations. The optimized structures of Fe(II) sulfides are shown in Table S3 249 of the electronic supplement. The $\ln \beta$ values obtained are shown in Table 3. Several 250 authors have reported the stability constant for FeHS⁺ (Dyrssen, 1988; Zhang and 251 Millero, 1994; Luther et al., 1996; Al-Farawati and van den Berg, 1999; Davison et 252 al., 1999). However, the only species that were identified by Dyrssen (1988) and 253 Davison et al. (1999) was the higher order complex Fe(HS)₂. It may be difficult to 254 distinguish the formation of Fe(HS)₂ from polymerized Fe(II) sulfides denoted as

aqueous FeS. We therefore used the structure of trans-Fe(HS)₂(H₂O)₄ to represent monomeric aqueous FeS.

The values of δ^{56} Fe for the 56 Fe/ 54 Fe ratio were calculated for the relevant species as a function of pH (Figs. 6b and 6c). Species other than sulfides show positive δ^{56} Fe at typical seawater pH. At lower pH, a Δ^{56} Fe fractionation of 0.5-1‰ may be expected between these species and sulfides. At pH ≥ 8.2 , Δ^{56} Fe between carbonates and sulfides increases to 1.6‰.

262 The growth of FeS clusters stabilizes FeS as a solid phase. The structure of FeS 263 clusters has been computed by Rickard and Luther (2007). The stable structure of Fe_2S_2 264 is similar to the basic structural component of mackinawite. Polymerization of the hydrated Fe₂S₂ molecules would make the FeS solid phase precipitate. A possible 265 266 structure of the hydrated Fe_2S_2 is $Fe_2S_2(H_2O)_4$ (Rickard and Luther, 2007), in which the 267 coordination number of Fe(II) is four. We successfully reproduced the structure by ab 268 *initio* calculations and the optimized structure of $Fe_2S_2(H_2O)_2$ is shown in Table S3. The 269 corresponding $\ln \beta$ values are shown in Table 3. Fe₂S₂(H₂O)₄ forms via dimerization of 270 Fe(II) mono-hydrogensulfides, according to,

$$2\text{FeHS}(\text{H}_2\text{O})_5^+ \rightleftharpoons \text{Fe}_2\text{S}_2(\text{H}_2\text{O})_4 + 2\text{H}^+ + 6\text{H}_2\text{O}$$
 (8)

271 The ln β values obtained for Fe₂S₂(H₂O)₂ are larger than those of Fe(II) 272 hydrogensulfides, yet still smaller than those of other aqueous Fe(II) species. This 273 suggests that the ⁵⁶Fe/⁵⁴Fe ratio of Fe²⁺ and Fe(II) chlorides, carbonates, and hydroxides 274 show heavier δ^{56} Fe values than that of solid FeS.

275 As for the Fe(II) case, the structures of Ni(II) hydrogensulfides were computed 276 (Table S2 of the electronic supplement). The ln β values obtained are shown in Table 2. The ln β values of Ni(II) hydrogensulfides are smaller than those of the other prevalent aqueous Ni(II) species. ⁶⁰Ni/⁵⁸Ni ratios in ferromanganese crusts suggest that hydrothermal fluids entering the ocean have δ^{60} Ni ~1.5‰ (Gall et al., 2013). As shown in Table 2, ln β is 1.4-1.5‰ (298 K) smaller for Ni(HS)₂ than for NiCO₃ and NiSO₄. The ligand exchange reactions of Ni sulfide-Ni carbonate and Ni sulfide-Ni sulfate possibly create δ^{60} Ni of the same magnitude as that suggested by Gall et al. (2013) for hydrothermal fluids.

284 In a previous study in which Fujii et al. (2013) computed ln ß at 298 K for Cu 285 species, $\ln \beta$ was found to be 1-2‰ smaller for Cu(II) sulfides than for Cu(II) carbonates, hydroxides, and hydrated Cu^{2+} . Again, this suggests that Cu^{2+} and Cu(II)286 287 chlorides, carbonates, and hydroxides are isotopically heavier than sulfides. Since 288 sulfide-bearing systems are reducing, isotope fractionation of Cu caused by the presence of Cu(I) must also be considered. Taking the standard redox potentials of the HS^{-}/SO_{4}^{2-} 289 and Cu²⁺/Cu⁺ couples (Pourbaix, 1974) into account, most of inorganic Cu(II) species 290 291 are reduced to Cu(I) under the redox conditions in which sulfides prevail. The 292 speciation of Cu(I) under hydrothermal conditions at 573 K has been evaluated by 293 Mountain and Seward (1999) and the prevailing species were found to be CuCl, CuCl₂, CuHS, and $Cu(HS)_2^{-}$. The higher order Cu(I) chloride complex $CuCl_3^{2-}$ is also identified, 294 295 but its small stability constant (Mei et al., 2013 and references therein) makes it a minor 296 species. The $\ln \beta$ values of these species were evaluated by *ab initio* methods by Seo et 297 al. (2007) and Sherman (2013) and the results obtained in the present study (Tables S5 298 and 5, and Fig. 7) agree with their values. At high pH and high sulfur concentration, a 299 Cu(I) dimer $Cu_2S(HS)_2^-$ exists (Mountain and Seward, 1999). As shown in Table 5, 300 increasing complexation of chlorides and sulfides results in decreasing $\ln \beta$. We

evaluated the δ^{65} Cu values of Cu(I) species from the speciation model of Mountain and 301 302 Seward (1999) and show the results in Fig. S2 of the electronic supplement as a 303 function of the mole fractions of Cu(I) species. The range of isotopic variation of the Cu(I) species at 573 K is about ± 0.1 %. The ln β values of Cu(II) chlorides and sulfides 304 305 (Fujii et al., 2013) at 573 K are 0.2-0.5‰ higher than those of corresponding Cu(I) species (Table 5). Under hydrothermal conditions, the δ^{65} Cu value of Cu(I) should be 306 307 0.2-0.5% lower than that of Cu(II) with a $\pm 0.1\%$ range of variation among Cu(I) 308 species. Cu(I) chloride may be isotopically light compared with Cu(I) sulfide (Fig. S2), and lighter if $CuCl_3^{2-}$ exists due to its small ln β (Fig. 7 and Table 5). 309

Dekov et al. (2013) reported δ^{65} Cu values of Cu⁰ (native copper) from sedimentary 310 311 layers of 0.41-0.95‰ and suggested that native Cu precipitation in the basaltic basement is a result of low temperature (293 to 358 K) hydrothermal processes under 312 313 anoxic and H₂S poor conditions. These authors proposed a reduction process of Cu(I) chloride $CuCl_2^{-1}$ to Cu^0 to explain the heavy isotope enrichment of the Cu^0 . Since the ln 314 β value of intramolecular vibrations for the uncomplexed Cu⁰ atom is zero, Δ^{65} Cu 315 between Cu^0 and $CuCl_2$ reduces to the negative of ln β for $CuCl_2$. In other words, the 316 reduced Cu^0 should be isotopically lighter than $CuCl_2$. This does not match the positive 317 δ^{65} Cu value for Cu⁰ observed for native copper in the oceanic crust by Dekov et al. 318 (2013). We here suggest that $CuCl_2^-$ was dissociated to CuCl and/or Cu^+ according to 319

$$\operatorname{CuCl}_{2}^{-} + \operatorname{H}_{2}O \rightleftharpoons \operatorname{CuCl}(\operatorname{H}_{2}O) + \operatorname{Cl}^{-}$$
(9)

$$\operatorname{CuCl}_{2}^{-} + 2\operatorname{H}_{2}\operatorname{O} \rightleftharpoons \operatorname{Cu}(\operatorname{H}_{2}\operatorname{O})_{2}^{+} + 2\operatorname{Cl}^{-}$$

$$\tag{10}$$

320 and that the dissociated species were then irreversibly reduced to Cu^0 ,

$$4\operatorname{CuCl}(\operatorname{H}_{2}\operatorname{O}) \to 4\operatorname{Cu}^{0} \downarrow + 4\operatorname{HCl} + \operatorname{O}_{2} + 2\operatorname{H}_{2}\operatorname{O}$$

$$\tag{11}$$

$$4Cu(H_2O)_2^+ \to 4Cu^0 \downarrow + 4H^+ + O_2 + 6H_2O$$
(12)

In these reactions, δ^{65} Cu for Cu⁰ reflects positive Δ^{65} Cu values between CuCl(H₂O) and CuCl₂⁻ and/or between Cu(H₂O)₂⁺ and CuCl₂⁻. From the ln β values shown in Table 5, isotope fractionation of +0.4-0.5‰ ⁶⁵Cu in Cu⁰ at 298-323 K may be expected.

Aggregation of Cu^0 atoms results in the formation of Cu metallic phase. Though it is difficult to simulate metals by using cluster model, intramolecular vibration of a diatomic Cu_2^0 computed gives $\ln \beta \sim 0.8\%$ at 298-323 K. (Table S7). This suggests that the aggregation of Cu^0 atoms also give a positive isotope fractionation. Hence, the isotope fractionation larger than +0.4-0.5‰ ⁶⁵Cu in Cu⁰ (native copper) may be possible from the above reaction scheme.

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331 **3.3. Isotope fractionation of Fe, Ni, Cu and Zn due to inorganic ligands**

332 This section summarizes the common features of isotope fractionation for Fe, Ni, Cu, and Zn estimated by *ab initio* methods. Polyatomic inorganic ligands such as SO_4^{2-} and 333 334 the phosphate ion are important for soil chemistry. Phosphate complexes of Zn and the ln β values of 66 Zn/ 64 Zn were computed by Fujii and Albarède (2012). Structures of 335 336 phosphate complexes of Fe (Tables S3 and S4), Ni (Table 2S), and Cu (Table S6) were 337 computed using the same approach as for Zn phosphates. The ln β values obtained are shown in Tables 2, 3, 4, and 5. The ln β values for a minor species, CuHPO₄⁺, which is 338 339 used only in this section, are separately shown in Table S8. Geochemical application of 340 the results for $\ln \beta$ of phosphates to soil-plant systems will be discussed in Section 3.4.

The ln β values of the divalent states of Fe, Ni, Cu, and Zn bound to hydrated cations, hydroxide, chloride, sulfate, sulfide, carbonate, and simple phosphate complexes are compared in Fig. 8 (note that metal cations are denoted as M²⁺). The magnitude of ln β values varies in the order Fe < Ni > Cu > Zn. This sequence is identical to that of complex stability derived from crystal field theory (Bersuker, 1996). Additional stabilization energy resulting from Cu site deformation (Jahn-Teller effect, Bersuker, 2006) may still be an issue.

The oxygen donor ligands, OH^- , SO_4^{2-} , CO_3^{2-} , HCO_3^- , HPO_4^{2-} , and $H_2PO_4^-$, tend to 348 349 give larger $\ln \beta$ values than those of hydrated cations. The values of $\ln \beta$ for variably deprotonated ligands decrease in the order $CO_3^{2-} > HCO_3^{-}$ and $HPO_4^{2-} > H_2PO_4^{-}$, which 350 reflects that ligand protonation decreases the energy of the bond with M²⁺. Chloride and 351 352 sulfide complexes show smaller $\ln \beta$ values than those of hydrated cations. Higher order 353 complexation of the cations with Cl⁻ or HS⁻ also results in a decrease of $\ln \beta$. A similar 354 decreasing trend was found for the hydrolysis of Fe and Ni. In contrast, Cu hydroxides 355 and carbonates show a trend different from that of the other metals. This may again reflect structural deformations specific to these species. In the first Cu^{2+} coordination 356 sphere, CO3²⁻ occasionally combines with H₂O to produce HCO3⁻ and OH⁻ (Fujii et al., 357 358 2013; Sherman, 2013). Hydrolysis also makes the hydration bond of an axial H₂O 359 molecule at the first coordination sphere weaker (Fujii et al., 2013; Sherman, 2013).

360 The ln β values of Fe(II), Fe(III), Ni(II), Cu(II), and Zn(II) complexes are shown in 361 Fig. 9 as a function of the bond length with ligands. The atomic distances between M 362 and O of the oxygen donor ligands, M and Cl, and M and S of HS are used as bond 363 lengths. A minor Fe(III) sulfide species, FeHS²⁺, was also computed, for which the 364 results are shown in Tables S4 and S9. The lines shown were calculated by the least 365 square method and did not take sulfides into consideration. The $\ln \beta$ values increase with decreasing bond lengths. Since CO_3^{2-} and HCO_3^{-} are treated as bidentate ligands, a 366 367 shorter distance of M-O was shown in Fig. 9. Some of the misfit may be due to the assumption of the bidentate character. The poor correlation of $\ln \beta$ vs bond length for 368 369 Cu(II) may be attributable to the distortion of complexes by the Jahn-Teller effect 370 (Bersuker, 2006). The $\ln \beta$ values for sulfide complexes also deviate from the 371 correlation lines and show smaller values. This can reflect the different bonding energies between O and S donor ligands as explained by the hard and soft acids and 372 373 bases (HSAB) principle (Pearson, 1968a,b).

374 Electronegativity as originally defined by Pauling (1960) is a measure of the power 375 of an atom in a molecule to attract electrons. The term was later extended to groups 376 (Huheey, 1965). We here show that the $\ln \beta$ values correlate with the group electronegativity of ligands (Fig. 10). The source of the values adopted for group 377 378 electronegativities (Allred, 1961; Huheev, 1965, 1966; Bratsch, 1985; Mullay, 1985; 379 Boyd and Boyd, 1992) are listed in Table 7. The $\ln \beta$ values increase with increasing 380 group electronegativity. The correlation between the ln β values of Cu(II) is stronger 381 with electronegativity (Fig. 10d) than with bond length Fig. 9d, The misfit of the $\ln \beta$ 382 values for sulfide complexes is also smaller for electronegativity than for bond length.

383 Overall, isotopologues with stronger chemical bonds, and therefore higher 384 intramolecular vibrational frequencies, have larger ln β (Bigeleisen and Mayer, 1947; 385 Urey, 1947). Both Figs. 9 and 10 demonstrate that the rule works even for transition 386 metal centered polyatomic molecules. Group electronegativities therefore may help 387 predict the direction and strength of isotope fractionation by ligand exchange reactions.
388 Such a property may be particularly useful for isotope fractionation induced by complex
389 metalloproteins.

- 390
- 391 **3.4. The soil-plant system**

 δ^{56} Fe in ferrous ion (Fe²⁺) is 1.3‰ heavier than that in ferrihydrite (Fe₅HO₈·4H₂O) (Beard et al., 1999). The magnitude of this fractionation is expected from the difference in ln β values between Fe(II) and Fe(III) species (see Tables 3 and 4). Beard et al. (1999) assigned isotope fractionation to biogenic effects via Fe-reducing bacteria and suggested electron exchange reactions between Fe and organic matter. Evaluating the ln β values of organometallic compounds of Fe, Cu, Ni, and Zn therefore is important.

The tricarboxylic citric acid is a major organic acid. The citric acid cycle known as the Krebs cycle is used by aerobic organisms to generate energy through the oxidation of acetate into CO_2 (Cowan, 1997). In addition, the citric acid cycle provides precursors of certain amino acids. Haem biosynthesis begins in the mitochondria, with the condensation of succinyl coenzyme A (CoA) produced during the citric acid cycle with glycin (Crichton, 2001).

Citric acid also plays an important role in the transport of trace metals in the
soil-plant system. Citrate is released from roots of vascular plants and act as a biological
chelating agent for the uptake of metals from soil. Isotope fractionation induced by
higher plants has been found for Fe (Guelke and von Blanckenburg, 2007; Kiczka et al.,
2010; Moynier et al., 2013b), Ni (Estrade et al., 2013), Cu (Weinstein et al., 2011;
Jouvin et al., 2012), and Zn (Weiss et al., 2005; Moynier et al., 2009). In a pioneering
work of the isotope fractionation of Zn in the soil-plant system, Weiss et al. (2005)

found that Zn was isotopically lighter in the shoots relative to the roots, with a δ^{66} Zn difference of -0.13 to -0.26. The origin of this isotope fractionation has been explained by the isotopic exchange between Zn(II) phosphates in soil and citrates (or malates) in plants (Fujii and Albarède, 2012).

415 Likewise, *ab initio* calculations of Fe isotope fractionation in higher plants by 416 the Moynier al. (2013b) suggest that roots of strategy-II et plants (Fe(III)-phytosiderophore) are isotopically heavier by about 1‰ (for δ^{56} Fe) than the 417 418 upper parts of the plants. Iron is transported as Fe(III)-citrate in the xylem or 419 Fe(II)-nicotianamine in the phloem. It was suggested that, even in the absence of redox reactions of Fe^{3+}/Fe^{2+} , change in speciation alone can create up to 1.5% isotope 420 421 fractionation (Moynier et al., 2013b). Here, we therefore focus on Fe(III) phosphates and Fe(III) citrates and estimate δ^{56} Fe created by exchange of Fe(III) among these 422 423 species.

The ${}^{56}\text{Fe}/{}^{54}\text{Fe}$ ratio in stems is ~1‰ lighter compared to that of roots and soil 424 425 (Kiczka et al., 2010; Moynier et al., 2013b). Fe(III) citrates are the major Fe species in 426 the xylem of stems in strategy I plants (Kiczka et al., 2010), whereas Fe(III) phosphates 427 and hydroxides are the prevalent species in soils. The prevalent forms were determined 428 using the work of Königsberger et al. (2000) for citrates, while the results of Ciavatta 429 and Iuliano (1995) were used for Fe(III) phosphates. The chemical species of Fe(III) 430 computed are shown in Fig. 11 and Table S4 of the electronic supplement. We assumed 431 sixfold-coordination of Fe(III) for all species. Moynier et al.'s (2013b) ab initio results for Fe citrates, Fe(cit) and Fe(cit)₂³⁻, where (cit) stands for the citrate ion $C_6H_5O_7^{3-}$, 432 tested different basis sets. In the present study of Fe(cit) and $Fe(cit)_2^{3-}$, we used the 433 6-311+G(d,p) basis set. As pointed out by Fujii and Albarède (2012), HPO₄²⁻ contacts 434

435 hydration water around Fe^{3+} , in which H₂O is dissociated to H⁺ and OH⁻. For example, 436 FeHPO₄(H₂O)₅⁺ (model molecule (b-3) of Fig. 11) acts as FeH₂PO₄OH(H₂O)₄⁺.

437 From the stability constants of orthophosphates (Childs, 1970), Fe(III) phosphates 438 (Ciavatta and Iuliano, 1995), citrates (Königsberger et al., 2000), hydroxides (Baes and 439 Mesmer, 1976), and dissociation constants of citric acid (Königsberger et al., 2000), the 440 speciation diagram for a diluted system (activity coefficients of chemical species are 441 treated to be unity) was prepared as functions of pH (Fig. S3 of the electronic 442 supplement). At low pH < 6, various chemical exchange reactions take place among 443 Fe(III) phosphates and citrates. In the neutral pH region, partly hydrolyzed Fe(III) 444 citrates are predominant. At high pH > 8, aqueous $Fe(OH)_3$ becomes a major Fe(III)445 species.

The ln β values for the ⁵⁶Fe/⁵⁴Fe ratio obtained are shown in Table 4. The 446 447 dependence of $\ln \beta$ on temperature is shown in Fig. 12b. The isotope fractionation factors δ^{56} Fe for Fe(III) phosphates, citrates, hydroxides, and hydrated Fe³⁺ ions are 448 shown in Fig. S3 as functions of pH. δ^{56} Fe of Fe(III) citrates is ~1‰ smaller than that of 449 Fe(III) phosphates or hydroxides. Kiczka et al. (2010) observed δ^{56} Fe ~1% in xylem of 450 451 stems, where Fe(III) citrates are major species in strategy I plants. This suggests that the magnitude of Fe isotope fractionation found in soil-stem (Kiczka et al., 2010) is 452 453 possible only by ligand exchange of Fe(III) phosphates-citrates.

Fe(II) nicotianamine is the major Fe species present in the central part of the plant roots (stele) in strategy I plants (Kiczka et al., 2010). The ⁵⁶Fe/⁵⁴Fe ratio in steles was more than 2‰ smaller than that of host soil (Kiczka et al., 2010). Since isotope fractionation is clearly larger than the δ^{56} Fe variability expected from Fe(III) species (Fig. 12b), it may be concluded that the isotope fractionation of Fe found in steles

results from redox reaction between Fe^{3+} and Fe^{2+} . As for the Fe(III) case, the ln β 459 460 values for Fe(II) phosphates and citrates were computed by ab initio methods. Relevant 461 Fe(II) species were chosen from the literature (Ciavatta et al., 1992; Königsberger et al., 462 2000). The optimized structures of Fe(II) complexes and their $\ln \beta$ values are shown in 463 Tables S3 and 3, and Fig. 12a. It is clear that $\ln \beta$ is 3-4‰ smaller for Fe(II) phosphates. 464 citrates, and hydroxides than for Fe(III) species (Table 4). These new results confirm that Fe isotope fractionation in plants of 2‰ or larger is due to Fe^{3+}/Fe^{2+} redox 465 466 reactions.

467 A similar computational work was conducted on Ni(II) and Cu(II). The results are 468 shown in Fig. S4 (Ni) and Fig. S5 (Cu) of the electronic supplement. Since the valence of Ni is fixed at (II), the transport of Ni in the soil-plant system is independent of redox 469 470 reactions. The relevant species of Ni(II) phosphates and citrates were chosen from the 471 literature (Taylor and Diebler, 1976; Hedwig et al., 1980). Their optimized structures 472 are given in Table S2. The $\ln \beta$ values obtained are shown in Table 2 and Fig. 12c. The δ^{60} Ni values of Ni(II) phosphates, citrates, hydroxides, and hydrated Ni²⁺ ions are 473 474 shown in Fig. 16b as a function of pH. Since Ni(II) hydroxides only exist in the high pH region of pH > 9 (Fig. S4), these may not be important in the soil-plant system. At 475 476 neutral pH, the prevalent Ni(II) species are phosphates and citrates, and the range of their δ^{60} Ni values is ~0.6‰. The magnitude of isotope fractionation overlaps with the 477 recent finding (Estrade et al., 2013) of Δ^{60} Ni = 0.2‰ between the roots and leaves of a 478 plant sample. With the decrease of pH, δ^{60} Ni of phosphates and citrates become very 479 similar (Fig. S4). At pH < 6, hydrated Ni²⁺ exists in the system and may have lower 480 ⁶⁰Ni/⁵⁸Ni values than those of phosphates and citrates. 481

482 Potentially significant species of Cu(II) phosphates and citrates were taken from 483 the literature (Petit-Ramel and Khalil, 1974; Ciavatta et al., 1993). The fourfold or 484 fivefold coordinated structures were computed. The optimized structures and the $\ln \beta$ 485 values obtained are shown in Table S6 and Table 6, respectively. The dependence of ln β on temperature is shown in Fig. 12d. As in the case of FeHPO₄(H₂O)₅⁺ (model 486 molecule (b-3) of Fig. 11), $CuH_2(PO_4)_2(H_2O)_2^{2-}$ behaves as $CuH_4(PO_4)_2(OH)_2^{2-}$. The 487 citrate ion (cit) is a tri-negative tridentate anion, where the three carboxylic groups of 488 the citric acid are dissociated. Most of the citrates shown in Table 6 (cit) are coordinated 489 to Cu^{2+} as tridentate anions, and as a pseudo-bidentate anion in the $Cu(cit)_2^{4-}$ complex. 490 This may be due to the Jahn-Teller effect (Bersuker, 2006), which involves the 491 492 distortion of the Cu site. A further dissociation of the alcoholic group of (cit) was also 493 suggested by nuclear magnetic resonance (NMR) analysis (Tananaeva et al. (1990), in the form of $C_6H_4O_7^{4-}$ denoted as $H_1(cit)$. Though a complexation of Cu^{2+} with $H_1(cit)$ 494 was reported (Petit-Ramel and Khalil, 1974), CuH₋₁(cit)²⁻ was found to be unstable and 495 the result suggests a fragmentation of (cit). This suggests that hydroxylation of Cu^{2+} 496 takes place before deprotonation of (cit). We therefore used $Cu(OH)^+$ and $Cu(OH)_2$ as 497 substitutes for $CuH_1(cit)^{2-}$ in the speciation diagram as a function of pH (Fig. S5). The 498 range of δ^{65} Cu for Cu(II) phosphates, citrates, hydroxides, and hydrated Cu²⁺ ions is 499 500 shown in Fig. S5 as a function of pH, where the ln β values of Cu(II) hydroxides and hydrated Cu²⁺ were taken from Fujii et al. (2013). At neutral pH, major Cu(II) species 501 are phosphates and citrates, and a range of ~0.5‰ can be expected for δ^{65} Cu. This range 502 overlaps with observations on higher plants (Weinstein et al., 2011; Jouvin et al., 2012). 503 A reduction of Cu²⁺ to Cu⁺ by a reductase within roots has also been reported by Jouvin 504 et al. (2012). Since the range of $\ln \beta$ values for Cu(I) species (Table 5) is ~2‰ smaller 505

than those of Cu(II) species at 298 K (Table 6), a fractionation of -0.84 to -0.11‰
between roots and nutrient solutions (Jouvin et al., 2012) may be expected.

508 At neutral pH, Zn(II) (Fujii and Albarède, 2012), Ni(II) (Fig. S4), and Cu(II) (Fig. 509 S5) phosphates show higher δ values than those of citrates. The δ values of Zn(II) and 510 Ni(II) phosphates decrease with decreasing pH and become similar to those of citrates at pH ~5. This is not the case for Cu(II): δ^{65} Cu of phosphates becomes smaller than that of 511 citrates by as much as 0.5‰ at pH = 4 (Fig. S5). The H^+ ion released by the adenosine 512 513 triphosphate (ATP) proton pump (Crichton, 2001) makes the pH of extracellular fluid 514 lower than that of the cytosol. If the pH of extracellular area close to the plasma membrane is as low as pH = 4. Cu(II) phosphates would show smaller δ^{65} Cu compared 515 516 to that of Cu (II) citrates. The Cu(II) ligand exchange may have contributed to the Cu 517 isotope fractionation observed between roots and nutrient solutions (Jouvin et al., 2012).

518

519 **3.5. Amino acids**

520 Transition metals are found in hundreds of proteins in the human body. The proteins are 521 involved in a large spectrum of critical biological functions, such as oxygen transport 522 (Fe), electron shuttling (Cu), structural control and protein degradation (Zn). These 523 functions are expressed via complexation of the metals with amino acids under various 524 structural environments, redox and pH conditions of biological fluids. Isotopic 525 variations in metals in organs and body fluids provide an enormous source of untapped 526 information relevant to normal and pathological conditions. In this section, we focus on 527 isotopic fractionation of Zn(II) and Cu(II) between amino acid complexes. Substantial 528 patterns of Zn (Balter et al., 2013; Moynier et al., 2013a) and Cu (Balter et al., 2013) isotope fractionation are observed in some organs such as the liver and the kidney, as 529

well as in blood components. To a large extent, these patterns reflect the binding of Zn
and Cu with different amino acids, variable redox states and electronegativity (Albarède,
2013; Balter et al., 2013; Moynier et al., 2013a).

533 A preliminary assessment of $\ln \beta$ for Zn(II)-amino acid complexes without 534 hydrated water has been done by Moynier et al. (2013a). In the present study, we present a set of ln β values for 66 Zn/ 64 Zn in aqueous Zn(II)-amino acid complexes of 535 536 cysteine (Cys), glutamine (Glu), histidine (His), methionine (Met), and threonine (Thr). 537 Inorganic aqueous Zn(II) complexes usually show sixfold coordinations (Fujii et al., 538 2010, 2011b, 2012), while fourfold complexation is usual for Zn(II) in proteins (Auld, 539 2001, 2009). We calculated both fourfold and sixfold coordinations. Amino acids 540 possess specific dissociable groups. For example, Cys has thiol (-SH), His has 541 imidazole $(C_3H_4N_2)$, and Glu has carboxyl (-COOH) groups. The hydroxyl group (-OH) 542 of Thr tends to bind to the metal ion from the outer coordination sphere, but when it 543 comes closer, it is deprotonated and binds directly to the metal ion (Grenouillet et al., 1973). These functional groups were treated to be dissociated and bind to hydrated Zn^{2+} . 544 545 Met possesses a functional group of thioether (C-S-C). The S donor was treated to be coordinated to hydrated Zn^{2+} . The bonding types of amino acids to cations were 546 547 reviewed by Lippard and Berg (1994).

An objective of this study is to gather information on isotope fractionation of transition metals via their bonding with amino acids in proteins. Besides the functional groups, amino groups (-NH₂) and carboxyl groups (-COO⁻) in the amino acids are possible to contact with metal cations (Lippard and Berg, 1994). In our calculations, redundant contacts of these groups with Zn^{2+} were restricted by protonation to be -NH₃⁺ and -COOH. The protonation sites of amino acids are referred to Dinadyalane at al. (2006). All Zn^{2+} -amino acid complexes are treated as 2+ charged molecules.

The optimized structures of Zn^{2+} -amino acid complexes are shown in Fig. 13a and 555 13b. The ln β values at temperatures ranging from 273 to 373 K are shown in Table 8. 556 557 The ln β values at a typical body temperature of 310 K are shown in Fig. 14a. It has 558 been pointed out that heavy isotopes tend to bind to O-donor ligands, whereas light 559 isotopes are positively fractionated by S-donor ligands (Albarède et al., 2011; Albarède, 560 2013; Balter et al., 2013; Moynier et al., 2013a). This is clearly seen in Fig. 14a for the 561 complexes with identical coordination number. The isotope fractionation correlated with 562 N-donor ligands may be intermediate between O-donor and S-donor systems or even 563 stronger than with O-donor ligands. Besides the donor type, coordination number is 564 important, implying that fourfold complexation gives larger ln β values relative to complexes with sixfold coordination. In β of Zn(His)²⁺ complexes is 0.2 to 0.6% larger 565 than that of $Zn(Cys)^{2+}$. This matches the observation that organs rich in proteins with 566 His residues show larger δ^{66} Zn than organs in which proteins rich in Cys residues 567 568 dominate (Balter et al., 2013; Moynier et al., 2013a).

We further computed large molecules of $Zn(His)_3(Cys)^{2+}$ and $Zn(His)_3(H_2O)^{2+}$, in which the coordination number of Zn(II) is four. The values at 310 K of ln β are 3.43‰, and 3.83‰, respectively ($\Delta^{66}Zn = 0.40\%$). These complexes are core structures of matrix metalloproteinases (MMPs), which are capable of degrading the extracellular matrix proteins (Van Wart, 1990). The displacement of the propeptide cysteine by water is induced by either proteolytic cleavage and/or conformational changes of the propeptide (Auld, 2009). Such a transformation switches the role of Zn from a 576 non-catalytic to a catalytic function. The potential of Δ^{66} Zn as a probe for understanding 577 the mechanism of "cystein switch" is therefore very strong.

578 The optimized structures of Cu^{2+} -amino acid complexes are shown in Fig. 13c. The 579 ln β values at temperatures ranging from 273 to 373 K are shown in Table 9. The ln β 580 values at the body temperature of 310 K typical for mammals are shown in Fig. 14b. 581 Again, the ln β of Cu^{2+} complexes with O and N-donor amino acids is ~1‰ higher than 582 those with S-donor amino acids.

583 Glutathione (L- γ -glutamyl-L-cysteinylglycine) is the most abundant intracellular non-protein thiol and an essential reducing agent of the cell. Glutathione in its reduced 584 585 (GSH) and oxidized (GSSG) forms is responsible for important biological functions 586 such as active transport of amino acids, enzyme activity, formation of complexes with microelements, e.g., Zn(II), and the redox status of the Cu^{2+}/Cu^{+} couple (Shtyrlin et al., 587 2005). The structure of the Zn^{2+} -GSH complex $Zn(GS)^{-}$ is tetrahedral, whereas that of 588 the Cu^{2+} -GSH complex $Cu(GS)H^{0}$ is a distorted planner (Chow et al., 1975). The 589 590 optimized structures are shown in Fig. 13a and 13c. The $\ln \beta$ values at temperatures 591 ranging from 273 to 373 K are listed in Tables 8 and 9. The $\ln\beta$ values at 310 K are 592 shown in Fig. 14. Zn(II) and Cu(II) are bound to O-, N-, and S- donors in GSH, and hence it may be reasonable to anticipate that the ln β values of Zn(GS)⁻ and Cu(GS)H⁰ 593 594 lie among those of amino acid complexes. The ln β value of Zn(GS)⁻ is 0.2‰ larger than that of $Zn(H_2O)_6^{2+}$. An equivalent Zn isotope fractionation is therefore expected 595 upon uptake of Zn(II) by GSH. On the other hand, the ln β value of Cu(GS)⁰ is close to 596 that of $Cu(H_2O)_5^{2+}$, suggesting that Cu isotopes are not fractionated by GSH uptake of 597 Cu(II). A computational study of the Cu^{2+}/Cu^{+} redox reaction in proteins indicates that 598

the fourfold coordination of Cu(II) changes to twofold coordination of Cu(I) (Pavelka and Burda, 2008). This suggests that the ln β values of twofold coordination Cu(I) species shown in Table 5 may be similar to ln β of the Cu(I)-GSH complex. The ln β value of Cu(GS)⁰ is 1.2‰ larger than that of Cu(H₂O)₂⁺ estimated by Fujii et al. (2013) (Fig. 14b). Isotope fractionation of this magnitude is therefore expected for the Cu²⁺/Cu⁺ pair as a result of redox glutathione activity.

605 As a final note, we also report values for Cu-lactate. The observation that, in most 606 cancer cells, glycolysis is remarkably enhanced and redirected from the citric acid cycle 607 towards lactic acid fermentation, even when oxygen is available, is known as the 608 Warburg effect (Pedersen, 2007). Because Cu(II)-lactate complexes are particularly strong (Piispanen, 1995) preliminary estimates of the isotope effect resulting from Cu²⁺ 609 610 complexation by L-lactic acid were calculated. A fivefold coordination of $CuC_{3}H_{5}O_{3}(H_{2}O_{3})^{+}$ suggested from the structure of aquobis-lactate-copper (Prout et al., 611 612 1968) was computed. The ln β is shown in Table 9. Lactate shows the largest ln β value of all the studied amino acid complexes, which is a hint that positive δ^{65} Cu should be 613 614 found in the cytosol of cancer cells. This is an additional and particularly strong 615 indication that stable isotope fractionation of the first transition series elements in biology should be useful biomarkers that would help understand the detailed 616 617 mechanisms of some biological processes.

618

619 **Conclusions**

620 The ln β values calculated for the inorganic species of Fe, Ni, Cu, and Zn are helpful for 621 understanding the isotopic variations found in surface seawater, the deep sea, and 622 hydrothermal systems. The isotope fractionation of Fe, Ni, Cu, and Zn found in higher 623 plants is understandable via isotopic exchange reactions between phosphates and 624 citrates. The ln β values calculated for Zn-amino acid complexes provide insights into a 625 range of isotopic variations observed for organs and body fluids of mammals.

626

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635

636 APPENDIX A

637 The isotope fractionation factor of an element M, δ M for an isotope pair of light isotope 638 A' and heavy isotope A, was estimated as follows. As an example, we consider a system 639 containing three chemical species X, Y, and Z.

Distribution of chemical species X, Y, and Z as functions of pH is calculated from reported stability constants. The mole fractions of χ_X , χ_Y , and $\chi_Z(\Sigma\chi = 1)$ calculated are set as those of isotopologues including the isotope A'. An ambient molar (mol dm⁻³) or molal (mol kg⁻¹) concentration of [M] is then set. The total concentration of A' is [M] $R_{A'}$ for the isotopic abundance $R_{A'}$ of A' in the whole system. The naturally occurring isotopic abundance is set to be $R_{A'}$.

646 The isotopic mass balance is kept at every pH.

$$[M]R_{A'} = [M]R_{A'}\chi_X + [M]R_{A'}\chi_Y + [M]R_{A'}\chi_Z$$
(a-1)

647 The isotope fractionation between the species X and Y is (if ln β is given in ‰ unit, use 648 $10^{-3} \ln \beta$ here),

$$\ln \beta_{\rm Y} - \ln \beta_{\rm X} = \frac{([A]/[A'])_{\gamma}}{([A]/[A'])_{\chi}} - 1$$
(a-2)

649 and hence,

$$[A]_{Y} = (\ln \beta_{Y} - \ln \beta_{X} + 1) \frac{[A']_{Y}}{[A']_{X}} [A]_{X}$$

$$= (\ln \beta_{Y} - \ln \beta_{X} + 1) \frac{\chi_{Y}}{\chi_{X}} [A]_{X}$$
(a-3)

650 Similarly, for the isotope fractionation between the species X and Z

$$[A]_{Z} = (\ln \beta_{Z} - \ln \beta_{X} + 1) \frac{\chi_{Z}}{\chi_{X}} [A]_{X}$$
(a-4)

651 The isotopic mass balance of A is,

$$[M]R_{A} = [A]_{X} + [A]_{Y} + [A]_{Z}$$

$$= \left(\frac{1 + \chi_{Y} \ln \beta_{Y} + \chi_{Z} \ln \beta_{Z} - (\chi_{Y} + \chi_{Z}) \ln \beta_{X}}{\chi_{X}}\right) [A]_{X}$$
(a-4)

652 The concentrations of A of species X, Y, and Z can thus be calculated. The isotope653 fractionation in the species X is

$$\delta M_{X} (\%) = \left(\frac{[A]_{X} / [A']_{X}}{([M]R_{A}) / ([M]R_{A'}]} - 1 \right) \times 1000$$

$$= \left(\frac{\chi_{X}}{1 + \chi_{Y} \ln\beta_{Y} + \chi_{Z} \ln\beta_{Z} - (\chi_{Y} + \chi_{Z}) \ln\beta_{X}} \times \frac{1}{\chi_{X}} - 1 \right) \times 1000$$

$$= \frac{(\chi_{Y} + \chi_{Z}) \ln\beta_{X} - \chi_{Y} \ln\beta_{Y} - \chi_{Z} \ln\beta_{Z}}{1 + [\chi_{Y} \ln\beta_{Y} + \chi_{Z} \ln\beta_{Z} - (\chi_{Y} + \chi_{Z}) \ln\beta_{X}]} \times 1000$$
(a-5)

654 Similarly, δM_Y and δM_Z can be calculated. Since we treat small ln β values, in practice, 655 the number of atoms in unit volume (or weight) is more useful than using molar (or 656 molal) concentration.

657

658 APPENDIX B

Software B3LYP with a basis set 6-311+G(d,p) is a standard DFT method for computing aqueous species of Ni(II), Cu(II), and Zn(II). Using the code with 6-311+G(d,p) basis set or higher is recommended for estimating the accurate electronic structures (de Bruin et al., 1999; Rulíšek and Havlas, 1999). Here, we discuss the accuracy of ln β values reported in this work.

664 Calculating how physicochemical properties vary with hydration of the species is a 665 common strategy for examining the accuracy of theoretical calculations of aqueous species (Fujii et al., 2010, 2011a,b). In the theoretical study on the hydration enthalpy of 666 Fe^{2+} and Fe^{3+} . Li et al. (1996) tested a small cluster model of 6 H₂O molecules as the 667 668 first coordination sphere and a large cluster model of 12 additional H₂O molecules as the second coordination sphere. For Fe^{3+} , the large cluster model brought the calculated 669 data closer to the experimental results. For Fe²⁺, both the small and the large cluster 670 modes reproduced the experimental results. Similarly, the hydration enthalpies of Ni²⁺ 671 and Zn^{2+} were appropriately reproduced by using both the small and the large cluster 672 673 models (Fujii et al, 2010, 2011a).

In a pioneering computational study on hydrated Mg^{2+} (Pye and Rudolph, 1998), 674 675 vibrational frequencies were well reproduced when using the large cluster of $Mg(H_2O)_{18}^{2+}$, while they were underestimated by 8-10% when the small cluster 676 $Mg(H_2O)_6^{2+}$ was used. A similar phenomenon was confirmed for Ni²⁺ and Zn²⁺, though 677 678 the hydration enthalpies were well predicted (Fujii et al, 2010, 2011a,b). The calculation 679 results are shown in Table B1. The totally symmetric stretching mode v_1 of hexaaqua 680 complexes is the fundamental intramolecular vibration mode, and the v_1 frequency is also shown in this table. The calculated v_1 frequencies of Ni(H₂O)₆²⁺ and Zn(H₂O)₆²⁺ 681

are smaller than the literature values determined by Raman spectrometry. Setting 12 H₂O molecules at the second coordination sphere brought the calculation results closer to the literature values. Adding the second hydration sphere increased ln β (298 K) by +0.6‰ for Ni and +0.3‰ for Zn.

686 Continuum solvation models are known to be useful to simulate the 687 physicochemical properties of aqueous complexes (Ginovska et al., 2008). The 688 conductor-like polarizable continuum model (CPCM) of solvation was tested for 689 hydrated ions of Fe, Ni, Cu, and Zn (Table B1). For divalent and trivalent cations, the 690 use of CPCM still increases v_1 frequency, but the effect is smaller than that of a second 691 coordination sphere. The CPCM assumption changed ln β (298 K) by less than 0.3‰.

The accuracy of computed ln β values for hydrated Fe²⁺ and Fe³⁺ has been 692 693 extensively discussed in the literature (Rustad et al., 2010; Saunier et al., 2011, and 694 references therein). It was pointed out that $\ln \beta$ may have been overestimated by some 695 computational studies. Rustad et al. (2010) argues that the absolute values of $\ln \beta$ for hydrated Fe^{2+} and Fe^{3+} are well reproduced by using large $Fe(H_2O)_{18}^{2+}$ clusters with a 696 conductor-like screening model (COSMO) of solvation. Our results for $\ln \beta$ with a small 697 698 cluster is ~0.4‰ (298 K), which is larger than Rustad et al.'s (2010) results for both Fe^{2+} and Fe^{3+} . 699

Isotope fractionation of ⁵⁶Fe/⁵⁴Fe between hydrated Fe³⁺ and Fe²⁺, with Δ^{56} Fe = δ^{56} Fe(Fe³⁺) - δ^{56} Fe(Fe²⁺), has been experimentally determined at 295 K (Johnson et al., 2002; Welch et al., 2003). It is a standard result that Δ^{56} Fe can be approximated by the difference of ln β (⁵⁶Fe/⁵⁴Fe) between hydrated Fe³⁺ and Fe²⁺. At present, two theoretical estimates of ln β (Ottonello and Zuccolini, 2009; Rustad et al., 2010) show

705 values consistent with the experimental results (Johnson et al., 2002; Welch et al., 2003) (Table B2). We computed $Fe(H_2O)_6^{2+}$ and $Fe(H_2O)_6^{3+}$ using B3LYP/6-311G with 706 707 various diffuse and polarization functions. A CPCM calculation including water 708 solvation was also tested. The results are shown in Fig. S6 of the electronic supplement. 709 It is clear that the higher basis set as 6-311+G(d,p) (see Table B2) is required for 710 computing the $\ln \beta$ values of aqueous species of iron. The use of CPCM decreased both ln β and Δ^{56} Fe. The small cluster model has been shown to produce results 711 systematically ~0.4‰ heavier than large clusters (Rustad et al., 2010) (see above). 712 Nevertheless, the calculated Δ^{56} Fe value reproduces the experimental result of Welch et 713 714 al. (2003). When the small cluster model is suspected to fail, it has been suggested to scale frequencies empirically for the effect of the second solvation sphere (Pye and 715 716 Rudolph, 1998). Since we showed that the accuracy of the small cluster model to reproduce relative isotope fractionation is adequate (e.g., Δ^{56} Fe, see Fig. S6 and Table 717 718 B2), scaling factors were not used.

719 Treatment of the second coordination sphere in computational studies turns out to be more difficult when the H₂O molecules at the first coordination sphere are 720 721 substituted by ligands. The large cluster model needs experimental evidence of 722 molecular arrangement in the second coordination sphere, but experiments are fraught 723 with technical difficulties. For the ligand-exchanged complexes, the solvation models of 724 pure water may not be adequate. For example, chloro-complexes exist in chloride 725 solutions, for which the solvent is not identical with pure water. Difficulties of the 726 solvation models are discussed in Ginovska et al. (2008).

We conclude that the absolute values of $\ln \beta$ estimated with the large cluster model are more accurate than estimates based on the small cluster model. Introducing the

729	solvation model may improve accuracy of $\ln \beta$, but its effect as shown in Table B1 is
730	probably not significant, and the scale of relative isotope fractionation still stands.
731	Although the absolute value of $\ln \beta$ estimated with the small cluster model may be
732	improved by addition of a second and more distant coordination spheres, the Δ values
733	estimated from Table B2 are reliable and the isotope fractionation scale presented in this
734	work can be used for predicting isotope fractionation among species in solution.

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1037	
1038	Table 1 Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁶⁶ Zn- ⁶⁴ Zn of
1039	Zn(II) complexes.

Zii(ii) complexes.						
0			Tempera	ature (K)		
Species	273	298	323	373	473	573
$\operatorname{Zn}(\operatorname{H_2O})_6^{2+}$	3.854	3.263	2.797	2.119	1.334	0.915
$ZnCl(H_2O)_5^+$	3.702	3.136	2.689	2.039	1.285	0.882
$ZnCl_2(H_2O)_4$	3.486	2.950	2.528	1.915	1.205	0.826
ZnCl ₃ (H ₂ O) ⁻	3.490	2.952	2.528	1.913	1.202	0.824
$ZnCl_4^{2-}$	2.722	2.293	1.957	1.474	0.921	0.629
$ZnSO_4(H_2O)_5$	4.154	3.527	3.031	2.306	1.460	1.006
$Zn(OH)_2(H_2O)_4$	4.185	3.567	3.075	2.350	1.495	1.032
$ZnHCO_3(H_2O)_3^+$	4.573	3.877	3.326	2.525	1.593	1.095
$ZnCO_3(H_2O)_3$	4.940	4.199	3.612	2.752	1.745	1.202

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1046	Table 2 Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁶⁰ Ni- ⁵⁸ Ni of
1047	Ni(II) complexes.

Tri(II) complexes.						
Spacios			Tempera	ature (K)		
Species	273	298	323	373	473	573
$Ni(H_2O)_6^{2+}$	6.383	5.412	4.644	3.525	2.224	1.528
$\text{NiCl(H}_2\text{O})_5^+$	6.228	5.280	4.531	3.439	2.169	1.490
$NiCl_2(H_2O)_4$	6.009	5.093	4.369	3.315	2.090	1.435
$NiSO_4(H_2O)_5$	6.907	5.868	5.044	3.840	2.433	1.676
$NiOH(H_2O)_5^+$	6.706	5.702	4.904	3.736	2.368	1.631
$Ni(OH)_2(H_2O)_4$	6.551	5.579	4.807	3.671	2.336	1.613
$NiHCO_3(H_2O)_4^+$	6.790	5.757	4.941	3.751	2.368	1.627
$NiCO_3(H_2O)_4$	6.841	5.814	4.999	3.806	2.411	1.661
$NiHS(H_2O)_5^+$	5.609	4.753	4.077	3.092	1.949	1.338
Ni(HS) ₂ (H ₂ O) ₄	5.191	4.401	3.776	2.866	1.809	1.243
$NiH_2PO_4(H_2O)_5^+$	6.806	5.774	4.958	3.767	2.380	1.636
NiHPO ₄ (H ₂ O) ₅	6.910	5.871	5.048	3.843	2.435	1.676
$NiH_2(cit)(H_2O)_3^+$	6.909	5.857	5.026	3.814	2.406	1.653
NiH(cit)(H ₂ O) ₃	6.876	5.833	5.008	3.805	2.404	1.653
Ni(cit)(H ₂ O) ₃	6.824	5.791	4.974	3.781	2.391	1.645
Ni(cit) ₂ ⁴⁻	5.008	4.228	3.616	2.732	1.714	1.174

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Table 3 Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁵⁶Fe-⁵⁴Fe of Fe(II) complexes.

			Tempera	ature (K)		
Species	273	298	323	373	473	573
$\operatorname{Fe}(\operatorname{H_2O})_6^{2+}$	6.013	5.095	4.370	3.314	2.088	1.433
$\operatorname{FeCl(H_2O)}_5^+$	5.746	4.869	4.176	3.166	1.995	1.369
$\operatorname{FeCl}_2(\operatorname{H}_2\operatorname{O})_4$	5.476	4.636	3.974	3.011	1.895	1.300
$\text{FeSO}_4(\text{H}_2\text{O})_5$	6.349	5.397	4.641	3.535	2.241	1.543
$FeOH(H_2O)_5^+$	6.444	5.488	4.726	3.607	2.291	1.579
$\operatorname{Fe(OH)}_2(\operatorname{H}_2O)_4$	6.318	5.394	4.656	3.565	2.275	1.573
${\rm Fe}_{2}({\rm OH})_{6}^{2}$	6.372	5.407	4.643	3.527	2.227	1.530
$\text{FeHCO}_3(\text{H}_2\text{O})_4^+$	6.213	5.267	4.520	3.432	2.167	1.489
$FeCO_3(H_2O)_4$	6.836	5.806	4.991	3.798	2.405	1.656
$FeHS(H_2O)_5^+$	5.210	4.413	3.784	2.868	1.806	1.240
$Fe(HS)_2(H_2O)_4$	4.676	3.960	3.395	2.574	1.621	1.113
$Fe_2S_2(H_2O)_4$	5.718	4.839	4.146	3.139	1.974	1.353
$FeH_2PO_4(H_2O)_5^+$	6.282	5.326	4.572	3.471	2.191	1.505
FeHPO ₄ (H ₂ O) ₅	6.927	5.883	5.056	3.846	2.434	1.675
$FeH_4(PO_4)_2(H_2O)_4$	6.256	5.300	4.546	3.447	2.173	1.492
$\operatorname{FeH}_3(\operatorname{PO}_4)_2(\operatorname{H}_2\operatorname{O})_4$	6.154	5.225	4.491	3.417	2.163	1.489
FeH(cit)(H ₂ O) ₃	6.175	5.236	4.495	3.413	2.156	1.482
Fe(cit)(H ₂ O) ₃	6.191	5.254	4.512	3.430	2.169	1.492
FeH(cit) ₂ ³⁻	5.443	4.603	3.942	2.983	1.877	1.288
$\operatorname{Fe}(\operatorname{cit})_2^{4-}$	4.583	3.867	3.306	2.496	1.565	1.072
Fe(cit) ₂ OH ⁵⁻	7.880	6.695	5.756	4.383	2.777	1.913

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1061	Table 4 Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁵⁶ Fe- ⁵⁴ Fe of
1062	Fe(III) complexes.

			Tempera	ature (K)		
Species	273	298	323	373	473	573
$\operatorname{Fe}(\operatorname{H_2O})_6^{3+}$	9.488	8.070	6.943	5.291	3.354	2.310
$\operatorname{FeCl}(\operatorname{H_2O})_5^{2+}$	8.634	7.333	6.301	4.792	3.030	2.084
$\operatorname{FeCl}_2(\operatorname{H}_2\operatorname{O})_4^+$	8.098	6.873	5.903	4.485	2.834	1.948
$\operatorname{FeSO}_4(\operatorname{H_2O})_5^+$	9.883	8.420	7.256	5.543	3.527	2.435
$FeOH(H_2O)_5^{2+}$	10.559	9.026	7.799	5.982	3.824	2.647
$\operatorname{Fe(OH)}_2(\operatorname{H}_2O)_4^+$	10.838	9.277	8.025	6.166	3.950	2.737
$\operatorname{Fe(OH)}_{3}(\operatorname{H}_{2}O)_{3}$	10.986	9.392	8.114	6.222	3.974	2.749
$\text{FeHCO}_{3}(\text{H}_{2}\text{O})_{4}^{2+}$	9.176	7.795	6.700	5.099	3.228	2.222
$\text{FeCO}_{3}(\text{H}_{2}\text{O})_{4}^{+}$	9.225	7.849	6.755	5.151	3.269	2.253
$FeH_{3}PO_{4}(H_{2}O)_{5}^{3+}$	9.941	8.454	7.273	5.541	3.512	2.418
$FeH_2PO_4(H_2O)_5^{2+}$	10.102	8.601	7.407	5.653	3.592	2.477
$FeHPO_4(H_2O)_5^+$	10.695	9.119	7.862	6.010	3.825	2.640
$FeH_5(PO_4)_2(H_2O)_4^{2+}$	10.147	8.634	7.432	5.668	3.598	2.480
$FeH_4(PO_4)_2(H_2O)_4^+$	10.258	8.733	7.520	5.739	3.647	2.516
$\operatorname{FeH}_{3}(\operatorname{PO}_{4})_{2}(\operatorname{H}_{2}\operatorname{O})_{4}$	10.465	8.915	7.682	5.866	3.730	2.574
$FeH_7(PO_4)_3(H_2O)_3^+$	10.188	8.665	7.457	5.684	3.608	2.487
$\operatorname{FeH}_6(\operatorname{PO}_4)_3(\operatorname{H}_2\operatorname{O})_3$	10.608	9.021	7.761	5.915	3.751	2.584
$Fe(cit)(H_2O)_3$	10.492	8.928	7.685	5.861	3.720	2.565
$Fe(cit)OH(H_2O)_2^{-1}$	10.074	8.582	7.394	5.647	3.591	2.478
$\operatorname{FeH(cit)_2^{2-}}$	10.002	8.491	7.295	5.547	3.510	2.415
$\operatorname{Fe}(\operatorname{cit})_2^{3-}$	8.573	7.262	6.228	4.724	2.980	2.047
Fe(cit) ₂ OH ⁴⁻	8.204	6.970	5.992	4.562	2.890	1.990

1069 Table 5 Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁶⁵Cu-⁶³Cu of Cu(I) complexes.

Spaaias			Temper	rature (K)			- Doforonao
Species	273	298	323	373	473	573	- Reference
$Cu(H_2O)_2^+$	3.368	2.867	2.468	1.882	1.193	0.822	Fujii et al. (2013)
CuCl(H ₂ O)	3.401	2.887	2.480	1.885	1.191	0.818	This study
	3.40	2.89	2.48	1.89	1.19	0.82	Seo et al. (2007)
CuCl ₂	2.775	2.350	2.014	1.526	0.960	0.659	This study
	2.71	2.29	1.97	1.49	0.94	0.64	Seo et al. (2007)
	2.79	2.36	2.03	1.53	0.96	0.66	Sherman (2013)
CuCl ₃ ²⁻	1.012	0.851	0.725	0.545	0.339	0.231	This study
	1.02	0.85	0.73	0.55	0.34	0.23	Seo et al. (2007)
	1.26	1.06	0.91	0.68	0.43	0.29	Sherman (2013)
CuHS(H ₂ O)	3.208	2.722	2.337	1.775	1.121	0.770	This study
	2.96	2.50	2.15	1.63	1.03	0.70	Sherman (2013)
Cu(HS) ₂	2.940	2.489	2.133	1.616	1.017	0.697	This study
	2.90	2.46	2.11	1.60	1.00	0.69	Seo et al. (2007)
	2.72	2.30	1.97	1.49	0.94	0.64	Sherman (2013)
$Cu_2S(HS)_2^2$	2.648	2.239	1.917	1.450	0.911	0.624	This study

1078	Table 6 Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁶⁵ Cu- ⁶³ Cu of
1079	Cu(II) complexes.

Cu(II) complexes:						
Spacios			Tempera	ature (K)		
species	273	298	323	373	473	573
$CuH_2PO_4(H_2O)_4^+$	5.515	4.684	4.026	3.063	1.939	1.334
$CuH_4(PO_4)_2(H_2O)_3$	5.553	4.714	4.050	3.079	1.947	1.339
$CuH_3(PO_4)_2(H_2O)_3$	5.290	4.492	3.861	2.937	1.860	1.280
$CuH_2(PO_4)_2(H_2O)_2^{2-}$	6.360	5.403	4.645	3.535	2.238	1.540
$CuH_2(cit)(H_2O)_2^+$	5.286	4.486	3.852	2.927	1.850	1.272
CuH(cit)(H ₂ O) ₂	5.622	4.772	4.099	3.117	1.972	1.357
$Cu(cit)(H_2O)_2$	6.092	5.177	4.451	3.389	2.147	1.479
$Cu(cit)_2^{4-}$	4.998	4.231	3.626	2.748	1.730	1.188

1084

Group	Electronegativity	Reference
H ₂ O	3.57	Boyd and Boyd (1992)
ОН	3.97	Mullay (1985)
Cl	3.16	Allred (1961)
SO_4	4.60	Huheey (1966)
HS	2.32	Huheey (1965)
HCO ₃	3.63	а
CO_3	4.33	Huheey (1966)
H_3PO_4	3.23	b
H_2PO_4	3.96	b
HPO ₄	5.11	b

1085Table 7 Group electronegativities.

^a Calculated by a method of Bratsch (1985) with electronegativities of H (2.20) (Allred,
1961) and CO₃.

^b Calculated by a method of Bratsch (1985) with electronegativities of H (2.20) (Allred,

1089 1961) and PO₄ (4.50) (Huheey, 1966).

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1094	Table 8 Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁶⁶ Zn- ⁶⁴ Zn of
1095	$Zn(II)$ -amino acid complexes and hydrated Zn^{2+} ion.

	•	Temperature (K)				
		273	298	310	323	373
fourfold	$Zn(H_2O)_4^{2+}$	4.539	3.853	3.577	3.310	2.516
	$Zn(Glu)(H_2O)_2^{2+}$	4.473	3.796	3.524	3.260	2.478
	$Zn(Thr)(H_2O)_3^{2+}$	4.774	4.056	3.767	3.487	2.654
	$Zn(His)(H_2O)_2^{2+}$	4.670	3.959	3.673	3.397	2.578
	$Zn(His)(H_2O)_3^{2+}$	4.635	3.930	3.647	3.373	2.561
	$Zn(Cys)(H_2O)_3^{2+}$	3.912	3.313	3.072	2.840	2.152
	$Zn(Met)(H_2O)_3^{2+}$	4.397	3.733	3.466	3.207	2.438
	Zn(GS) ⁻	4.311	3.655	3.392	3.137	2.381
sixfold	$Zn(H_2O)_6^{2+}$	3.854	3.263	3.026	2.797	2.119
	$Zn(Glu)(H_2O)_4^{2+}$	3.888	3.292	3.053	2.822	2.139
	$Zn(Thr)(H_2O)_5^{2+}$	3.916	3.315	3.075	2.842	2.154
	$Zn(His)(H_2O)_4^{2+}$	3.541	2.996	2.777	2.566	1.943
	$Zn(His)(H_2O)_5^{2+}$	3.724	3.150	2.921	2.699	2.043
	$Zn(Cys)(H_2O)_5^{2+}$	3.196	2.702	2.504	2.313	1.750
	$Zn(Met)(H_2O)_5^{2+}$	3.478	2.947	2.734	2.528	1.918

1100	
1101	Table 9 Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁶⁵ Cu- ⁶³ Cu of
1102	Cu(II)-amino acid complexes, lactate, and hydrated Cu ²⁺ ion.

· · ·	Temperature (K)				
	273	298	310	323	373
$Cu(H_2O)_5^{2+}$	5.355 ^{<i>a</i>}	4.546 ^{<i>a</i>}	4.220	3.905 ^{<i>a</i>}	2.968 ^{<i>a</i>}
$Cu(Glu)(H_2O)_3^{2+}$	5.230	4.436	4.117	3.808	2.891
$Cu(Thr)(H_2O)_4^{2+}$	5.220	4.429	4.110	3.803	2.889
$Cu(His)(H_2O)_3^{2+}$	5.274	4.470	4.148	3.836	2.911
$Cu(His)(H_2O)_4^{2+}$	5.299	4.492	4.168	3.855	2.926
$Cu(Cys)(H_2O)_4^{2+}$	3.981	3.369	3.124	2.888	2.187
$Cu(Met)(H_2O)_4^{2+}$	4.632	3.932	3.650	3.378	2.568
$Cu(GS)H^0$	4.945	4.194	3.892	3.600	2.734
$CuC_{3}H_{5}O_{3}(H_{2}O)_{3}^{+}$	5.530	4.695	4.359	4.034	3.068

^{*a*} Fujii et al. (2013).

Table B1 Vibrational frequencies and $\ln\beta$ values.

Fe(II)

re(II)			
	v_1	ln β at 298 K	Reference
	(cm^{-1})	(‰)	
$Fe(H_2O)_6^{2+}$	325	5.095	This study
$Fe(H_2O)_6^{2+} + CPCM$	338	4.912	This study
Raman	390	-	Kanno and Hiraishi (1982)
	V1	In B at 298 K	Reference
	(cm^{-1})	(‰)	
$Fe(H_2O)_6^{3+}$	407	8.070	This study
$Fe(H_2O)_6^{3+} + CPCM$	467	7.758	This study
Raman	509	-	Sharma (1973)
Raman	506	-	Kanno and Hiraishi (1982)
Ni(II)			
	v_1	ln β at 298 K	Reference
	(cm^{-1})	(‰)	
$Ni(H_2O)_6^{2+}$	347	5.432	Fujii et al. (2011)
$Ni(H_2O)_6^{2+}$	347	5.412	This study
$Ni(H_2O)_6^{2+} + CPCM$	353	5.524	This study
$Ni(H_2O)_{18}^{2+}$	394	6.046	Fujii et al. (2011)
Raman	390	-	Kanno (1977,1978)
Raman	390	-	Edwards and Knowles (1992
Raman	395	-	Bickley et al. (1993)
Cu(I)			
	v_1	ln β at 298 K	Reference
	(cm^{-1})	(‰)	
$Cu(H_2O)_2^+$	370	2.867	Fujii et al. (2013)
$Cu(H_2O)_2^+ + CPCM$	361	2.862	This study
Raman	410	-	Applegarth (2014)
	-		

- 1117 1118

- 1120 1121 Table B 1122 1123 **Cu(II)** Table B1 (continued).

	ν_1	ln β at 298 K	Reference	
	(cm^{-1})	(‰)		
$Cu(H_2O)_5^{2+}$	383 ^a	4.546	Fujii et al. (2013)	
$Cu(H_2O)_5^{2+} + CPCM$	383 ^a	4.547	This study	
$Cu(H_2O)_6^{2+}$	361	4.288	Fujii et al. (2013)	
$Cu(H_2O)_6^{2+} + CPCM$	371	4.297	This study	
Raman	440	-	Hester and Plane (1964)	
Raman	436	-	Davis and Chong (1972)	
Raman	434	-	Kanno (1977,1978)	
Raman	445	-	Castro and Jagodzinski (1991)	
^{<i>a</i>} v_2 (A_1) mode of fivefold coordination (Raman active).				

1125 1126 Zn(II)

	v_1	ln β at 298 K	Reference
	(cm^{-1})	(‰)	
$Zn(H_2O)_6^{2+}$	333	3.263	Fujii et al. (2011)
$Zn(H_2O)_6^{2+} + CPCM$	353	3.280	Fujii et al. (2011)
$Zn(H_2O)_{18}^{2+}$	380	3.576	Fujii et al. (2011)
Raman	390	-	Irish et al. (1963)
Raman	379	-	Yamaguchi et al. (1989)
Raman	385	-	Maeda et al. (1995)
Raman	390	-	Rudolph and Pye (1999)
Raman	389	-	Mink et al. (2003)

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1	1	3	1
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54	$- 1 able BZ \Delta T C 101 T C _{aq} - 1 C _{aq} (275 K)$				
		$\ln \beta (Fe^{3+})$	$\ln \beta (Fe^{2+})$	Δ^{56} Fe	Reference
		(‰)	(‰)		
	6 H ₂ O	8.223	5.194	3.03	This study
	6 H ₂ O (CPCM)	7.905	5.007	2.90	This study
	6 H ₂ O	8.35 ^{<i>a</i>}	5.08 ^{<i>a</i>}	3.27	Ottonello and Zuccolini (2009)
	18 H ₂ O (COSMO)	7.87	4.76	3.11	Rustad et al. (2010)
	Experimental	-	-	2.75 ± 0.15	Johnson et al. (2002)
	Experimental	-	-	2.95 ± 0.53	Welch et al. (2003)

1132 Table B2 Δ^{56} Fe for Fe³⁺_{aq}-Fe²⁺_{aq} (295 K)

1133 ^{*a*} Estimated from ln β at 298K by using $1/T^2$ dependence.

1135 Figure Captions

1136

Figure 1. Mole fractions of Cu(II) species and Cu isotopic variations (65 Cu/ 63 Cu) in seawater as a function of pH at 298 K. (a) Mole fractions of Cu(II) species reproduced from Zirino and Yamamoto (1972). (b) δ^{65} Cu of Cu(II) species relative to the bulk solution. (c) δ^{65} Cu of Cu(II) carbonates, chlorides, hydroxides, sulfates, and hydrated Cu²⁺ ions relative to the bulk solution. Though the metal concentration is not strictly required, total concentration of Cu was set to 1.1×10^{-8} mol kg⁻¹ (Macleod et al., 1994).

Figure 2. Mole fractions of Zn(II) species and Zn isotopic variations (66 Zn/ 64 Zn) in seawater as a function of pH at 298 K. (a) Mole fractions of Zn(II) species reproduced from Zirino and Yamamoto (1972). (b) δ^{66} Zn of Zn(II) species relative to the bulk solution. (c) δ^{66} Zn of Zn(II) carbonates, chlorides, hydroxides, sulfates, and hydrated Zn²⁺ ions relative to the bulk solution. Total concentration of Zn was set to 7.5 × 10⁻⁸ mol kg⁻¹ (Macleod et al., 1994).

1150

Figure 3. Mole fractions of Ni(II) species and Ni isotopic variations (60 Ni/ 58 Ni) in seawater as a function of pH at 298 K. (a) Mole fractions of Ni(II) species were estimated from Zirino and Yamamoto (1972). Stability constants of Ni(II) species were taken from the literature (Turner et al., 1981; Foulliac and Criaud, 1984; Byrne et al., 1988). (b) δ^{60} Ni of Ni(II) species relative to the bulk solution. (c) δ^{60} Ni of Ni(II) carbonates, chlorides, hydroxides, sulfates, and hydrated Ni²⁺ ions relative to the bulk solution. Total concentration of Ni was set to 2.9 × 10⁻⁸ mol kg⁻¹(Macleod et al., 1994).

Figure 4. Mole fractions of Fe(II) species and Fe isotopic variations (56 Fe/ 54 Fe) in seawater as a function of pH at 298 K. (a) Mole fractions of Fe(II) species were estimated from Zirino and Yamamoto (1972). Stability constants of Fe(II) species were taken from the literature (Foulliac and Criaud, 1984; Byrne et al., 1988). (b) δ^{56} Fe of Fe(II) species relative to the bulk solution. (c) δ^{56} Fe of Fe(II) carbonates, chlorides, hydroxides, sulfates, and hydrated Fe²⁺ ions relative to the bulk solution. Total concentration of Fe was set to 3.6×10^{-8} mol kg⁻¹ (Macleod et al., 1994).

1166

Figure 5. Temperature dependence of ln β for aqueous Fe species. The ln β values of Fe(II) (Table 3) and Fe(III) (Table 4) species are shown as a function of T^2 . (a) Fe(II) and (b) Fe(III). Vertical scale factor of (b) is 3/2 of (a). The temperature dependence of ln β for aqueous species of Ni, Cu, and Zn can be found in Fujii et al. (2011a,b, 2013).

1171

Figure 6. Mole fractions of Fe(II) species and Fe isotopic variations (${}^{56}Fe/{}^{54}Fe$) in 1172 1173 seawater-like matrix under sulfide-rich conditions. (a) Mole fractions of Fe(II) species 1174 were reproduced from Fig. 22a of Rickard and Luther (2007), in which total concentrations of Fe and S are reported to be 10⁻⁶ mol kg⁻¹ and 10⁻³ mol kg⁻¹, 1175 respectively. The HS^{-}/SO_{4}^{2-} redox couple was set to be disabled in their calculation. (b) 1176 δ^{56} Fe of Fe(II) species relative to the bulk solution. The ln β values at 298 K in Table 3 1177 were used. (c) δ^{56} Fe of Fe(II) carbonates, chlorides, hydroxides, sulfides, and hydrated 1178 Fe^{2+} ions relative to the bulk solution. 1179

Figure 7. Temperature dependence of $\ln \beta$ for Cu(I) species. The $\ln \beta$ values of hydrated Cu⁺, Cu(I) chlorides, and hydrogensulfides (Table 5) are shown as a function of T^2 . The temperature dependence of $\ln \beta$ for aqueous species of Cu(II) can be found in Fujii et al. (2013).

1185

Figure 8. ln β for Fe(II), Ni(II), Cu(II), and Zn(II) species. The metal cations are denoted as M²⁺. The ln β values at 298 K for hydrated cation, hydroxide, chloride, sulfate, sulfide, carbonate, and simple phosphate complexes are taken from Tables 1, 2, 3, 6, S8, Fujii and Albarède (2012), and Fujii et al., (2013).

1190

1191 Figure 9. $\ln \beta$ vs bond length. The $\ln \beta$ values (298 K) of hydrated complexes with the 1192 ratio M:ligand = 1:1 were taken from Tables 1, 2, 3, 4, 6, S8, Fujii and Albarède (2012), and Fujii et al., (2013), where ligands are OH⁻, Cl⁻, SO_4^{2-} , HS^- , HCO_3^{-} , CO_3^{2-} , and 1193 H₂PO₄⁻. For Fe(III), the ln β value for FeH₃PO₄³⁺ is also shown. For Zn(II) hydroxide 1194 1195 and sulfide, the ln β values for Zn(OH)₂ and Zn(HS)₂ are shown. For reference, the ln β values for hydated M^{2+} (or M^{3+}) are shown. The atomic distances between M and O of 1196 1197 the oxygen donor ligands, M and Cl, and M and S of HS were used as bond lengths. The 1198 atomic distances can be calculated from the Cartesian coordinates given in the 1199 electronic supplement. The correlation lines were drawn for complexes, except for 1200 sulfides, by least squares methods.

1201

1202 Figure 10. ln β vs group electronegativity. The ln β values (298 K) of hydrated 1203 complexes with the ratio M:ligand = 1:1 were taken from Tables 1, 2, 3, 4, 6, S8, S9, Fujii and Albarède (2012), and Fujii et al., (2013), where ligands are OH⁻, Cl⁻, SO₄²⁻, HS⁻, HCO₃⁻, CO₃²⁻, H₂PO₄⁻, and HPO₃²⁻. For Fe(III), the ln β value for FeH₃PO₄³⁺ is also shown. For Zn(II) hydroxide and sulfide, the ln β values for Zn(OH)₂ and Zn(HS)₂ are shown. The correlation lines were drawn for all complexes by least squares methods.

1209

Figure 11. Molecular structures of hydrated Fe³⁺, aqueous Fe(III) hydroxides,
phosphates, and citrates. The structures were drawn using GaussView5 (Gaussian Inc.)
(Dennington et al., 2009).

1213

Figure 12. Temperature dependence of ln β for Fe, Ni, Cu, and Zn phosphates and citrates. The ln β values of hydrated cations, phosphates, and citrates are shown as a function of T^2 . Vertical scale range was fixed at 7‰. (a) Fe(II) (Table 3), (b) Fe(III) (Table 4), (c) Ni(II) (Table 2), (d) Cu(II) (Table 6 and ln β of Cu²⁺ was reproduced from Fujii et al., 2013), and (e) Zn(II) (the ln β values were taken from Fujii and Albarède, 2012).

1220

Figure 13. Molecular structures of hydrated cations and amino acid complexes of Zn(II) and Cu(II). (a) Zn(II), fourfold complexation, (b) Zn(II), sixfold complexation, and (c) Cu(II), fivefold complexation. Abbreviations are Glu (glutamine), Thr (threonine), His (histidine), Cys (cysteine), and Met (methionine). Deprotonated glutathione (GSH) is shown as GS⁻. The structures were drawn using GaussView5 (Gaussian Inc.) (Dennington et al., 2009).

- 1227
- 1228 Figure 14. Zn (66 Zn/ 64 Zn) and Cu (65 Cu/ 63 Cu) isotopic variations in model molecules of
- 1229 Zn^{2+} and Cu^{2+} amino acid complexes at 310 K. (a) Zn(II), (b) Cu(II). ln β of aqueous
- 1230 Cu(I) species, $Cu(H_2O)_2^+$, which is estimated to be 2.664‰ at 310 K from Fujii et al.
- 1231 (2013). Here shown together for comparison.
- 1232
























(c) Fe(III) citrates





(b-4) $\text{FeH}_5(\text{PO}_4)_2(\text{H}_2\text{O})_4^{2+}$







(c-2) Fe(cit)OH(H₂O)₂-

(c-4) Fe(cit)₂³⁻

(b-8) FeH₆(PO₄)₃(H₂O)₃



(b-6) FeH₃(PO₄)₂(H₂O)₄

(b-1) FeH₃PO₄(H₂O)₅³⁺



(a-1) $Fe(H_2O)_6^{3+}$

(b) Fe(III) phosphates







(a-4) Fe(OH)₃(H₂O)₃-











H C O P Fe

















Supporting Information

Density Functional Theory Estimation of Isotope Fractionation of Fe, Ni, Cu, and Zn Among Species Relevant to Geochemical and Biological Environments

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Table S1	Optimized structure Cartesian coordinates of Zn(II) complexes.
Table S2	Optimized structure Cartesian coordinates of Ni(II) complexes.
Table S3	Optimized structure Cartesian coordinates of Fe(II) complexes.
Table S4	Optimized structure Cartesian coordinates of Fe(III) complexes.
Table S5	Optimized structure Cartesian coordinates of Cu(I) complexes.
Table S6	Optimized structure Cartesian coordinates of Cu(II) complexes.
Table S7	Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁶⁵ Cu- ⁶³ Cu of Cu ₂ ⁰ .
Table S8	Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁶⁵ Cu- ⁶³ Cu of Cu(II) mono-hydrogen phosphate.
Table S9	Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁵⁶ Fe- ⁵⁴ Fe of Fe(III) mono-hydrogensulfide.
Fig. S1	Mole fractions of Fe(II) and Fe(III) species and Fe isotopic variations $({}^{56}\text{Fe}/{}^{54}\text{Fe})$ in seawater at pH = 8.2 and 298 K.
Fig. S2	Copper, Cu(I), isotopic variations (65 Cu/ 63 Cu, δ^{65} Cu) under hydrothermal conditions at 573 K.
Fig. S3	Mole fractions of Fe(III) species and Fe isotopic variations (56 Fe/ 54 Fe) in a soil-plant system as a function of pH at 298 K.
Fig. S4	Mole fractions of Ni(II) species and Ni isotopic variations (⁶⁰ Ni/ ⁵⁸ Ni) in a soil-plant system as a function of pH at 298 K.
Fig. S5	Mole fractions of Cu(II) species and Cu isotopic variations (65 Cu/ 63 Cu) in a soil-plant system as a function of pH at 298 K.
Fig. S6	Fe isotopic variation (56 Fe/ 54 Fe) for hydrated Fe ${}^{2+}$ and Fe ${}^{3+}$ species at 295 K.

а

$Zn(H_2O)_6^{2+}$			
Zn	0.000026	0.000014	-0.000010
0	0.612463	-1.838912	-0.878282
0	-1.384498	-1.050191	1.228861
0	1.503404	-0.217831	1.489653
0	1.384546	1.050202	-1.228858
0	-1.503503	0.217749	-1.489545
0	-0.612487	1.838946	0.878196
Н	2.303939	1.251699	-1.006328
Н	1.212923	1.409191	-2.110385
Н	0.270221	-2.717425	-0.662548
Н	1.280310	-1.948331	-1.569258
Н	-1.212869	-1.409192	2.110381
Н	-2.303892	-1.251681	1.006332
Н	-1.682200	-0.393479	-2.217494
Н	-2.125345	0.954279	-1.568664
Н	-1.280341	1.948337	1.569170
Н	-0.270256	2.717469	0.662488
Н	2.125240	-0.954366	1.568780
Н	1.682084	0.393386	2.217617
$ZnCl(H_2O)_5^+$			
Zn	-0.180338	-0.029988	-0.086367
Cl	-2.359692	-0.364869	-0.464125
0	0.017984	-0.069148	2.116516
0	1.913451	0.310455	0.148381
0	-0.469724	2.134093	0.200672
0	0.289817	-0.058811	-2.249425
0	0.072240	-2.246607	0.183937
Н	0.680290	0.642784	-2.785399
Н	-0.547619	-0.297163	-2.674361
Н	2.214903	0.256898	1.064866
Н	2.622024	-0.007593	-0.423181
Н	0.107589	2.896764	0.076394

Table S1. Optimized structure Cartesian coordinates of Zn(II) complexes.

Н	-1.372087	2.397323	-0.027047
Н	0.731065	-2.867416	-0.150595
Н	-0.799596	-2.609107	-0.033948
Н	-0.507402	0.595266	2.581325
Н	-0.218643	-0.926438	2.494922

 $ZnCl_2(H_2O)_4$

Zn	0.000057	0.001479	0.000449
Cl	2.298580	-0.009645	0.007299
Cl	-2.298479	0.012577	-0.006437
0	0.005415	1.832372	1.338329
0	-0.004213	-1.829651	-1.337181
0	-0.012441	-1.336301	1.831427
0	0.011375	1.339144	-1.830575
Н	0.778046	-1.778869	-1.900637
Н	-0.780481	-1.770724	-1.908088
Н	0.763213	-1.908616	1.778482
Н	-0.795300	-1.898317	1.774447
Н	0.792398	1.903756	-1.773208
Н	-0.766224	1.908934	-1.778206
Н	-0.774994	1.780949	1.904335
Н	0.783594	1.773879	1.906700

 $ZnCl_3(H_2O)^-$

Zn	-0.014019	-0.024136	0.084944
Cl	2.270336	0.023341	0.262670
Cl	-1.077383	-1.770294	-0.812798
Cl	-1.027117	2.025090	0.264056
0	-0.021341	-0.025761	2.475465
Н	0.941440	0.078216	2.427447
Н	-0.377791	0.874395	2.419510
ZnCl ₄ ²⁻			
Zn	0.000238	-0.000337	-0.000044
Cl	-0.515644	-0.728532	2.195323
Cl	-1.732854	-0.658214	-1.477573
Cl	0.202003	2.361543	-0.021691

C1	
U	

$ZnSO_4(H_2O)_5$				
Zn	0.083898	-0.105521	-0.003450	
S	0.282612	2.898252	0.474807	
Ο	0.055116	0.097291	2.083435	
Ο	2.166439	0.142821	-0.124202	
Ο	0.008340	-0.072862	-2.241333	
Ο	0.260230	-2.259594	0.014772	
Ο	-0.398684	1.804597	-0.465980	
Ο	-0.068284	4.230640	0.017707	
Ο	-0.247189	2.564428	1.860762	
Ο	1.762175	2.579955	0.413382	
Ο	-2.154948	-0.097850	0.096459	
Н	2.515003	0.091137	-1.021640	
Н	2.158605	1.150490	0.115266	
Н	-0.083026	1.141057	2.138988	
Н	0.880213	-0.081483	2.547761	
Н	1.173226	-2.547075	0.139509	
Н	-0.277815	-2.740912	0.654357	
Н	-2.079154	0.863711	-0.078597	
Н	-2.495338	-0.166184	0.996914	
Н	-0.705841	-0.475505	-2.747270	
Н	-0.169005	0.887882	-2.195682	

2.046075 -0.974202 -0.695982

 $Zn(OH)_2(H_2O)_4$

Zn	-0.063224	0.000870	-0.035787
0	-0.083285	-0.521389	1.848932
0	-0.838479	0.106545	-2.162500
0	-0.323524	-2.373685	0.085427
0	1.500768	0.531654	-1.083524
0	-2.269127	-0.121771	0.477831
0	-0.122743	2.372221	-0.324229
Н	-1.295122	0.953497	-2.232241
Н	0.098715	0.288563	-2.378711
Н	-0.079977	-2.130015	1.008437
Н	0.330925	-2.995810	-0.245339

Н	-1.934474	-0.295256	1.381266
Н	-2.567831	-0.974233	0.138951
Н	-0.047405	2.993992	0.405488
Н	0.785451	2.135836	-0.624283
Н	0.533448	-0.142511	2.478088
Н	2.367847	0.161213	-0.908252

 $ZnHCO_3(H_2O)_3^+$

Zn	0.019328	0.198704	-0.033726
0	2.034268	-0.028613	0.344280
0	1.169426	1.905321	-0.271140
0	3.388661	1.727865	0.131851
0	-0.349392	-1.024681	-1.654930
0	-1.695870	1.374920	-0.214476
0	-0.360675	-0.931459	1.662662
С	2.181023	1.200632	0.066375
Н	0.354322	-1.605697	-1.974836
Н	-0.893856	-0.771390	-2.411535
Н	-2.574418	1.309351	0.180177
Н	-1.465476	2.312999	-0.282336
Н	0.457420	-1.200743	2.104778
Н	-1.081202	-1.493174	1.972429
Н	3.342102	2.665258	-0.105299

 $ZnCO_3(H_2O)_3$

Zn	-0.354832	-0.136112	-0.013998
0	1.710998	-0.254434	0.064240
0	2.742997	1.777514	0.017787
0	0.479534	1.610019	-0.051983
0	-0.089696	-1.346745	1.698573
0	0.081797	-1.430761	-1.642625
0	-2.234571	0.857382	0.008699
С	1.736543	1.106243	0.011089
Н	-0.335515	-1.064674	2.587063
Н	0.875818	-1.181957	1.581755
Н	-0.036564	-1.192638	-2.569129

Н	1.023181	-1.236920	-1.400225
Н	-3.009807	0.794572	-0.560422
Н	-1.871200	1.756124	-0.063340

2+			
$Ni(H_2O)_6$			
Ni	0.000056	0.000036	-0.000025
0	0.635333	-1.815809	-0.808912
0	-1.325485	-1.022179	1.246374
0	1.498116	-0.134704	1.445983
0	1.325691	1.022184	-1.246427
0	-1.498496	0.134410	-1.445539
0	-0.635475	1.815900	0.808641
Н	2.243391	1.240743	-1.034035
Н	1.137818	1.356966	-2.133909
Н	0.299529	-2.691230	-0.572423
Н	1.297263	-1.929057	-1.504630
Н	-1.137561	-1.356958	2.133845
Н	-2.243242	-1.240563	1.034047
Н	-1.660503	-0.494908	-2.161596
Н	-2.124628	0.864679	-1.544054
Н	-1.297307	1.928970	1.504485
Н	-0.299448	2.691354	0.572592
Н	2.125096	-0.864347	1.543752
Н	1.660551	0.494916	2.161677
NiCl(H ₂ O) ₅ ⁺			
Ni	-0.188832	0.000008	-0.000167
Cl	2.099674	0.001078	0.000320
0	-0.281079	-1.307078	1.658158
0	-0.141090	1.741304	1.212221
0	-0.281909	1.307052	-1.658346
0	-0.138826	-1.741463	-1.212295
0	-2.320528	-0.001197	-0.000006
Н	-0.433838	2.227389	-1.402632
Н	0.591749	1.296049	-2.077871
Н	0.592419	-1.294888	2.077990
Н	-0.431569	-2.227606	1.402271
Н	0.791144	-1.908928	-1.430018

Table S2. Optimized structure Cartesian coordinates of Ni(II) complexes.

Н	-0.661221	-1.892194	-2.010071
Н	-2.886052	-0.377906	0.684148
Н	-2.886990	0.374437	-0.683973
Н	0.788572	1.909849	1.430400
Н	-0.664060	1.891352	2.009745

 $NiCl_2(H_2O)_4$

Ni	-0.000157	0.001903	0.000672
Cl	2.339396	-0.010945	0.009962
Cl	-2.339561	0.013070	-0.010490
0	0.015470	1.731777	1.247265
0	0.005631	-1.729384	-1.244906
0	-0.022685	-1.244406	1.731386
0	0.001531	1.247600	-1.730483
Н	0.799522	-1.679874	-1.794065
Н	-0.765946	-1.677141	-1.824361
Н	0.749276	-1.823844	1.684798
Н	-0.816075	-1.793458	1.673711
Н	0.775193	1.824285	-1.679310
Н	-0.790338	1.799700	-1.680303
Н	-0.758070	1.687023	1.824654
Н	0.807359	1.674658	1.798607

 $NiSO_4(H_2O)_5$

	5		
Ni	1.030205	-0.017393	-0.010801
S	-2.038870	0.021551	0.080203
0	1.511958	1.520299	1.415772
0	1.243632	-1.633589	1.405999
0	3.105673	-0.033782	-0.541456
0	0.475420	-1.492837	-1.339329
0	0.527004	1.637461	-1.160623
0	-3.204264	0.037689	0.944878
0	-1.919486	-1.242106	-0.759467
0	-1.908814	1.240736	-0.812237
0	-0.713573	0.032259	0.956047
Н	1.326133	2.294752	0.863065
Н	0.704577	1.411211	1.946856

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 $NiOH(H_2O)_5^+$

2 3			
Ni	-0.011743	-0.033958	-0.019375
0	-1.369819	1.425809	-0.750138
0	-1.173161	-1.524313	-0.966001
0	1.480111	0.001303	-1.264944
0	1.083010	1.714018	0.557346
0	-1.381182	-0.154773	1.600134
0	1.442472	-1.291971	0.898868
Н	1.708687	1.471289	-0.167302
Н	1.590263	1.829766	1.369643
Н	-1.812515	1.304794	-1.599407
Н	-0.953704	2.299031	-0.760473
Н	-0.750790	-2.079633	-1.632936
Н	-1.812083	-2.067577	-0.488270
Н	-2.047389	0.544609	1.619974
Н	-1.099410	-0.310218	2.509944
Н	1.408994	-2.221772	1.149977
Н	1.993666	-1.204737	0.088608
Н	1.432435	0.024780	-2.223559

 $Ni(OH)_2(H_2O)_4$

Ni	0.000055	-0.000065	-0.000005
0	-1.140086	-1.680443	0.775350
0	-1.481103	1.456205	0.706970
0	1.589880	0.173538	1.198347
0	1.481122	-1.456282	-0.707011
0	-1.589716	-0.173738	-1.198358

0	1.140039	1.680536	-0.775165
Н	1.973761	-1.075498	0.069988
Н	1.967987	-1.220885	-1.503978
Н	-1.642771	-1.546490	-0.067214
Н	-0.626228	-2.488863	0.667806
Н	-1.967934	1.220850	1.503969
Н	-1.973706	1.075390	-0.070034
Н	-1.525365	-0.131507	-2.154479
Н	0.625768	2.488728	-0.667875
Н	1.642364	1.546857	0.067662
Н	1.525606	0.131373	2.154476

 $NiHCO_3(H_2O)_4^+$

Ni	-0.110353	-0.099469	-0.028369
0	1.798355	0.177802	0.682150
0	2.988246	1.918155	-0.144407
Ο	0.908065	1.495907	-0.833658
Ο	-1.024864	1.239110	1.339183
0	0.487402	-1.230438	-1.725608
0	-0.613287	-1.768540	1.141127
Ο	-2.022640	-0.193743	-0.896406
С	1.887938	1.182948	-0.092496
Н	-1.420562	-2.291387	1.070209
Н	0.003099	-2.224961	1.726206
Н	-0.713069	2.148383	1.231699
Н	-1.034724	1.052884	2.286538
Н	-2.646653	0.515230	-0.695152
Н	-2.092038	-0.388108	-1.839672
Н	0.926981	-2.086353	-1.646272
Н	1.021527	-0.694506	-2.328648
Н	3.640247	1.557173	0.473854
$N1CO_3(H_2O)_4$			
NT:	0.0(512)	0 100575	0.040100

	5	-			
Ni			-0.265136	-0.189575	-0.042189
0			1.608155	-0.126476	0.604163
0			2.818925	1.741848	0.141828

0	0.663119	1.502040	-0.541112
0	-1.215795	1.358052	1.254975
0	0.739949	-1.537356	-1.435772
0	-0.770329	-1.951614	1.082100
0	-2.211985	-0.018517	-0.934010
С	1.793679	1.100790	0.074914
Н	-0.875923	-2.726314	0.516465
Н	0.034077	-2.102161	1.597337
Н	-0.547901	1.921046	0.791899
Н	-1.016706	1.394577	2.197094
Н	-2.638778	0.574557	-0.298150
Н	-2.201752	0.455944	-1.773846
Н	1.533055	-1.293194	-0.905345
Н	0.902104	-1.200641	-2.324708
	c>↓ +		
N1HS(H ₂	O)5 [°]	0.002209	0.017467
INI S	-0.139930	0.005508	0.01/40/
3	2.100938	-0.13/290	-0.084317
0	-0.210373	1.01626	-1.422409
0	-0.331481	-1.001020	-1.3/32/7
0	-0.635103	-1.500880	1.409436
0	0.1316/3	1.620888	1.402405
0	-2.299151	0.234255	-0.05/488
Н	-0.638819	-2.309823	1.034805
Н	-0.104622	-1.589152	2.201314
Н	0.353/49	1.528705	-2.199114
Н	0.024514	2.429355	-0.98/208
Н	1.089986	1.552249	1.562211
H	-0.306800	1.866//5	2.226622
Н	-2./33286	0.952479	-0.532350
H	-2.90/821	-0.100228	0.611248
H	0.563574	-1./53242	-1./18118
H	-0.961135	-1.688406	-2.102417
Н	2.376599	-1.212273	0.728241
Ni(HS) ₂ ($H_2O)_4$		
Ni	-0.002734	0.018011	0.004964

S	2.410727	-0.061575	0.007929
S	-2.416191	-0.042234	-0.050189
0	0.258801	1.193986	1.818633
0	-0.243885	1.814357	-1.203721
0	0.169682	-1.228121	-1.773016
0	-0.196442	-1.779824	1.217637
Н	-0.199079	-0.857113	-2.583564
Н	1.139500	-1.125367	-1.816684
Н	1.223495	1.030151	1.811833
Н	-0.091995	0.746661	2.598587
Н	0.235685	-2.560803	0.850892
Н	-1.151355	-1.860790	1.030869
Н	-2.670671	0.708932	1.041597
Н	0.034672	2.616858	-0.746065
Н	-1.217894	1.767481	-1.123533
Н	2.648527	1.029832	-0.748863

 $NiH_2PO_4(H_2O)_5^+$

Ni	-0.103023	-0.037123	-0.043001
Р	-0.191908	3.025478	0.397924
0	-0.265913	0.261352	2.003092
0	1.923107	0.481298	-0.290343
0	-0.003812	-0.578044	-2.096143
0	0.419862	-2.067284	0.243047
0	-0.619676	1.843872	-0.485433
0	-0.815778	4.391404	-0.156350
0	1.398595	3.176502	0.013423
0	-2.194292	-0.321476	-0.215109
0	-0.405059	2.845875	1.875927
Н	0.382688	-0.062516	2.637884
Н	-0.364819	1.256858	2.159592
Н	-2.475185	0.601812	-0.320228
Н	-2.716249	-0.693185	0.507089
Н	0.078701	-2.652718	0.929457
Н	0.463169	-2.574038	-0.578544
Н	0.658401	-0.055607	-2.568601

Н	-0.832510	-0.482831	-2.584009
Н	2.662676	0.039345	0.141639
Н	2.052853	1.450597	-0.219392
Н	1.831869	3.965018	0.366116
Н	-1.417802	4.819897	0.465679

NiHPO₄(H₂O)₅

Ni	-0.270102	-0.179779	0.036626
Р	-0.145193	2.816898	0.153754
0	-0.459548	0.247129	2.091354
0	1.688075	0.111157	-0.253432
0	-0.104934	-0.420638	-2.110583
0	0.404166	-2.201530	0.293606
0	-0.966637	1.669755	-0.486934
0	-0.763123	4.218934	-0.349000
0	1.303541	2.804357	-0.535736
0	-2.390656	-0.501068	0.119124
0	-0.099144	2.761008	1.667760
Н	0.163725	-0.158819	2.702051
Н	-0.316496	1.254857	2.120478
Н	-2.563775	0.428577	-0.115507
Н	-2.599616	-0.566323	1.060041
Н	0.053034	-2.824417	-0.353808
Н	1.281041	-1.915594	-0.027472
Н	0.849316	-0.230021	-2.095594
Н	-0.521379	0.415414	-2.369142
Н	2.287425	0.157566	0.497837
Н	1.640885	1.868361	-0.537195
Н	-0.828592	4.845707	0.380116

NiH₂(cit)(H₂O)₃⁺

Ni	1.172985	-0.151442	0.183709
0	-0.369626	0.068338	1.372212
0	-2.437965	-0.546448	1.941342
0	0.526813	1.383695	-1.059016
0	-0.320677	3.314064	-1.590350

0	-1.339481	-3.130598	-1.214519
0	0.295163	-1.775346	-0.727440
0	-3.587127	0.166385	-0.207265
0	2.903835	-0.547916	-0.933330
0	1.842913	-1.382320	1.780260
0	2.416252	1.278921	1.150331
С	-2.185350	0.239454	-0.288800
С	-1.868157	1.710644	-0.656283
С	-1.748847	-0.732053	-1.461150
С	-1.618565	-0.130721	1.137418
С	-0.869516	-1.900786	-1.098186
С	-0.474092	2.087882	-1.099561
Н	-2.127655	2.345423	0.199143
Н	-2.563843	1.993447	-1.455639
Н	-1.194464	-0.189090	-2.229944
Н	-2.681805	-1.070237	-1.911985
Н	3.544866	0.166438	-1.034715
Н	2.870936	-1.039059	-1.762806
Н	1.049167	-1.463582	2.332606
Н	2.192441	-2.269868	1.632507
Н	2.694083	0.960780	2.020013
Н	2.123719	2.191630	1.261554
Н	-3.774925	-0.164811	0.694963
Н	-2.265246	-3.142829	-1.495396
Н	-1.154829	3.803849	-1.627377

NiH(cit)(H₂O)₃

Ni	1.214400	-0.237596	0.053464
0	-0.354280	-0.071713	1.285019
0	-2.486007	-0.179788	1.935564
0	0.521868	0.974943	-1.404894
0	0.249992	2.884328	-0.276195
0	-1.560627	-3.200321	-0.773491
0	0.275456	-2.015752	-0.644284
0	-3.551679	0.168901	-0.345946
0	2.922993	-0.378102	-1.199232

0	1.868966	-1.456023	1.680717
0	2.194135	1.476070	0.845917
С	-2.139725	0.226168	-0.390916
С	-1.720308	1.647331	-0.866509
С	-1.683702	-0.856750	-1.450575
С	-1.622626	-0.037336	1.075036
С	-0.915854	-2.039430	-0.926992
С	-0.214555	1.893406	-0.853207
Н	-2.205988	2.383080	-0.226028
Н	-2.105047	1.782610	-1.882914
Н	-1.027780	-0.383343	-2.178993
Н	-2.599849	-1.182269	-1.945172
Н	3.603907	0.152710	-0.764746
Н	2.573175	0.189346	-1.904894
Н	1.040100	-1.335239	2.178353
Н	1.894898	-2.387351	1.428165
Н	2.114623	1.563717	1.802775
Н	1.635131	2.206591	0.452179
Н	-3.758140	0.080233	0.606230
Н	-2.495829	-3.097153	-0.994469

Ni(cit)(H₂O)₃⁻

Ni	1.247018	-0.120799	0.156658
0	-0.313941	0.077934	1.405240
0	-2.461413	-0.067362	1.987684
0	0.660257	1.190028	-1.305813
0	0.236185	3.069514	-0.178227
0	-1.242938	-3.148416	-0.782270
0	0.550850	-1.828511	-0.640517
0	-3.457290	0.238292	-0.337971
0	2.819944	-0.467343	-1.329422
0	1.785345	-1.615930	1.691021
0	2.112636	1.589440	1.136068
С	-2.031269	0.266770	-0.345354
С	-1.634115	1.702556	-0.815525
С	-1.543283	-0.834013	-1.360957

С	-1.573343	0.060360	1.142313
С	-0.724239	-2.043191	-0.873832
С	-0.147298	2.043690	-0.766081
Н	-2.178774	2.423787	-0.205669
Н	-1.986818	1.803155	-1.848121
Н	-0.950481	-0.346444	-2.135605
Н	-2.455993	-1.229915	-1.803067
Н	2.488049	0.285353	-1.845171
Н	2.314665	-1.230927	-1.655196
Н	1.007745	-1.373431	2.220540
Н	1.431794	-2.282529	1.069418
Н	1.688178	1.582573	2.002712
Н	1.644152	2.315272	0.649025
Н	-3.677976	0.091351	0.602077
$Ni(cit)_2^{4-}$			
Ni	-0.000202	-0.002402	-0.000524
Ο	1.767611	-0.052490	-1.156130
Ο	3.891461	-0.194414	-1.849616
Ο	0.646092	-1.685758	1.104346
Ο	1.707338	-3.235136	2.326307
Ο	2.260714	2.226565	2.947971
Ο	0.940601	1.229555	1.436990
Ο	4.985111	-0.547281	0.352533
Ο	-1.768037	0.044813	1.155035
Ο	-3.893118	0.172971	1.847365
Ο	-0.643747	1.680601	-1.106676
Ο	-1.704402	3.234571	-2.323258
Ο	-2.264341	-2.225371	-2.951545
Ο	-0.944100	-1.234405	-1.436691
Ο	-4.984670	0.550622	-0.351729
С	3.530380	-0.417957	0.507946
С	3.110065	-1.768515	1.136941
С	3.371358	0.819047	1.424865
С	2.985542	-0.200456	-0.953132
С	2.052498	1.446670	1.975008
С	1.680085	-2.231945	1.557469

С	-3.530208	0.419016	-0.507668
С	-3.107035	1.773318	-1.126552
С	-3.373274	-0.811538	-1.433624
С	-2.986300	0.189825	0.951952
С	-2.055266	-1.446230	-1.978200
С	-1.677723	2.231139	-1.554721
Н	3.454288	-2.544581	0.440988
Н	3.733255	-1.887331	2.030223
Н	3.991782	0.617881	2.305130
Н	3.870625	1.646420	0.903907
Н	5.044531	-0.445758	-0.628262
Н	-3.440872	2.544922	-0.420573
Н	-3.736541	1.903915	-2.013635
Н	-3.985388	-0.599009	-2.317134
Н	-3.883052	-1.638900	-0.923044
Н	-5.045061	0.437622	0.627737

2			
$\operatorname{Fe}(\operatorname{H_2O}_6^{2+})$			
Fe	0.000854	0.001698	0.002572
0	2.180247	-0.000258	-0.000286
0	-0.000085	2.169469	0.048148
0	-0.000079	0.019597	2.143082
0	-2.179755	-0.000258	-0.000306
0	-0.000083	-2.168247	-0.048653
0	-0.000067	-0.019935	-2.141410
Н	2.757238	0.758701	-0.163289
Н	2.756999	-0.759457	0.162445
Н	-2.756749	0.758877	-0.162394
Н	-2.756510	-0.759631	0.161516
Н	-0.000132	2.727559	0.838279
Н	-0.000099	2.761464	-0.716868
Н	-0.000096	-2.760761	0.715924
Н	-0.000131	-2.725777	-0.839134
Н	0.778586	0.035119	2.717459
Н	-0.778875	0.035167	2.717278
Н	0.778537	-0.035638	-2.715771
Н	-0.778801	-0.035689	-2.715591
FeCl(H ₂ O) ₅ ⁺			
Fe	-0.047300	-0.000408	0.000859
Cl	2.203867	0.003597	-0.003518
0	-0.457746	-1.508774	1.578056
0	-0.020171	1.916027	1.244654
0	-0.462908	1.507311	-1.578014
0	-0.016908	-1.917274	-1.242924
0	-2.209371	0.000166	0.000466
Н	-0.276160	2.392027	-1.235421
Н	-0.006131	1.443737	-2.426347
Н	-0.001871	-1.446736	2.426955
Н	-0.272815	-2.393633	1.234867
Н	0.902338	-2.083057	-1.496692

Table S3. Optimized structure Cartesian coordinates of Fe(II) complexes.

Н	-0.573264	-2.220548	-1.970658
Н	-2.769369	-0.516006	0.592185
Н	-2.764733	0.517121	-0.594965
Н	0.898299	2.085132	1.498940
Н	-0.578192	2.218574	1.971433

 $\operatorname{FeCl}_2(\operatorname{H}_2\operatorname{O})_4$

Fe	-0.000024	0.001477	0.000504
Cl	2.334141	-0.009280	0.007543
Cl	-2.334144	0.012223	-0.006716
0	0.078543	1.840318	1.349602
0	0.069227	-1.838303	-1.347894
0	-0.085932	-1.346992	1.839492
0	-0.061731	1.350505	-1.839056
Н	0.911186	-1.784062	-1.818346
Η	-0.631022	-1.791656	-2.010036
Η	0.613001	-2.010870	1.798032
Η	-0.928624	-1.815339	1.778651
Н	0.643723	2.007186	-1.793583
Н	-0.899928	1.827488	-1.783544
Н	-0.625452	1.800068	2.008185
Н	0.917582	1.778202	1.824302

FeSO₄(H₂O)₅

	÷		
Fe	0.963581	-0.036794	-0.006732
S	-2.103307	0.091907	0.053528
0	1.649029	1.782860	1.255139
0	1.477238	-2.007631	1.096753
0	3.131664	-0.118186	-0.265691
0	0.328662	-1.668384	-1.340920
0	0.473868	1.753323	-1.191188
0	-3.230431	0.100233	0.967668
0	-2.062834	-1.114766	-0.863859
0	-1.957490	1.365442	-0.756542
0	-0.742512	-0.001777	0.881951
Н	1.256671	2.404394	0.618548
Н	1.084063	1.821831	2.037428

0.509710	-1.731155	-2.284017
-0.706114	-1.528039	-1.232110
1.039658	-2.539019	0.410149
0.899802	-2.057485	1.869350
0.663152	1.882966	-2.125814
-0.569766	1.692259	-1.091180
3.503611	0.619263	0.237352
3 441460	-0 926858	0.164570
	0.509710 -0.706114 1.039658 0.899802 0.663152 -0.569766 3.503611 3.441460	0.509710-1.731155-0.706114-1.5280391.039658-2.5390190.899802-2.0574850.6631521.882966-0.5697661.6922593.5036110.6192633.441460-0.926858

 $FeOH(H_2O)_5^+$

- 2 - 5			
Fe	0.000222	-0.082203	-0.234095
0	-1.385300	1.499494	-0.807072
0	-1.362396	-1.640465	-1.038600
0	1.626413	-0.015355	-1.187092
0	1.015261	1.962395	0.444104
0	-1.216390	-0.232987	1.565299
0	1.340695	-1.470531	1.051196
Н	1.733272	1.684574	-0.157804
Н	1.421939	2.305812	1.247746
Н	-1.982650	1.568073	-1.560725
Н	-0.932503	2.347947	-0.695090
Н	-1.276481	-2.048690	-1.908185
Н	-1.834126	-2.265937	-0.475686
Н	-1.923493	0.392835	1.765067
Н	-0.752909	-0.432276	2.388484
Н	1.345997	-2.420502	1.215583
Н	2.062374	-1.290899	0.421844
Н	1.847915	-0.134837	-2.112884

 $Fe(OH)_2(H_2O)_4$

Fe	-0.000260	0.000061	-0.000029
0	-1.058841	-1.938205	0.755668
0	-1.582364	1.723203	0.679660
0	1.579209	0.145930	1.133767
0	1.581590	-1.721695	-0.682316
0	-1.579980	-0.147432	-1.133066

0	1.061003	1.937944	-0.753685
Н	2.030643	-1.233419	0.047770
Н	2.137739	-1.643064	-1.463873
Н	-1.641739	-1.754874	-0.007114
Н	-0.462848	-2.639917	0.465932
Н	-2.138700	1.647132	1.461216
Н	-2.031509	1.233371	-0.049300
Н	-1.587178	-0.072456	-2.089430
Н	0.464940	2.640072	-0.465133
Н	1.641982	1.753578	0.010360
Н	1.585987	0.069477	2.090023
Fe ₂ (OH) ₆ ²	!		
Fe	-0.038997	0.017342	0.054048
Fe	2.456853	0.099663	-1.979911
0	-0.033815	0.176443	2.041927
0	1.859052	-0.792624	-0.199426
0	0.558526	0.908855	-1.726623
0	2.451545	-0.062361	-3.967606
0	4.394981	0.327263	-1.569708
0	-1.977758	-0.204885	-0.356247
Н	2.457183	-0.667884	0.542061
Н	-0.041131	0.783283	-2.466730
Н	3.206491	-0.589848	-4.243923
Н	4.794989	0.889356	-2.239515
Н	-2.378752	-0.768444	0.311756
Н	-0.787289	0.706048	2.318358
FeHCO ₃ (I	$(H_2O)_4^+$		
Fe	-0.162404	-0.065415	-0.013894
0	1.922838	0.208264	0.788573
0	3.030830	1.980540	-0.080097
0	0.950870	1.498195	-0.702496
0	-1.579472	0.507537	1.460397
0	-0.151883	-1.655909	-1.417747

-0.151883-1.655909-1.4177470.080700-1.6203411.631585-1.5006840.759945-1.673705

С	1.962653	1.206530	0.018041
Н	-0.131287	-2.557387	1.716685
Н	1.006939	-1.498062	1.888816
Н	-2.052502	1.343501	1.547868
Н	-1.477784	0.117200	2.338406
Н	-2.444842	0.931213	-1.771508
Н	-1.028722	1.541848	-1.996280
Н	0.474766	-2.384036	-1.504625
Н	-0.507536	-1.454083	-2.293197
Н	3.716299	1.648794	0.517592

FeCO₃(H₂O)₄

Fe	-0.422145	-0.278536	0.004348
0	1.613924	-0.237547	0.160285
0	2.680734	1.765842	-0.004654
0	0.418263	1.577865	-0.161918
0	-1.684261	0.995589	1.365073
0	0.242421	-1.943393	-1.359961
0	-0.078193	-1.831778	1.590645
0	-1.708692	0.646220	-1.589203
С	1.671744	1.101790	-0.002484
Н	-0.063810	-2.747624	1.287163
Н	0.853811	-1.538319	1.587301
Н	-1.007703	1.685249	1.197250
Н	-1.699613	0.835286	2.315188
Н	-2.546142	1.023755	-1.293974
Н	-1.069054	1.385236	-1.583246
Н	1.143671	-1.593028	-1.197602
Н	0.082435	-1.891365	-2.308839
FeHS(H ₂ O) ₅ ⁺			
Fe	-0.048342	-0.011667	0.026668
S	2.256730	-0.248663	-0.175267
0	-0.277421	1.731376	-1.371323

-0.048342	-0.011667	0.026668
2.256730	-0.248663	-0.175267
-0.277421	1.731376	-1.371323
-0.351898	-1.793478	-1.417257
-0.719907	-1.646013	1.378368
0.228017	1.787126	1.438293

0	-2.197966	0.334713	0.061537
Н	-0.664687	-2.494870	0.918691
Н	-0.445888	-1.794775	2.291278
Н	0.096778	1.747232	-2.260844
Н	0.024774	2.535731	-0.928118
Н	1.175475	1.797060	1.638744
Н	-0.239491	2.093368	2.225168
Н	-2.646113	0.999552	-0.474209
Н	-2.854296	-0.185647	0.538556
Н	0.556783	-1.952978	-1.715328
Н	-0.934422	-1.945198	-2.171950
Н	2.448199	-1.137923	0.821848

Fe(HS)₂(H₂O)₄

Fe	0.000245	-0.000211	-0.000454
S	1.760517	0.157825	1.643969
S	-1.759883	-0.157490	-1.644850
0	-1.022401	-1.849206	0.908895
0	-1.416849	1.644483	0.883469
0	1.416550	-1.645240	-0.883810
0	1.022224	1.850069	-0.908245
Н	2.124316	-1.478983	-0.238019
Н	1.795934	-1.557778	-1.765622
Н	-1.493686	-2.146263	0.115657
Н	-0.360082	-2.525631	1.094704
Н	-1.796964	1.555959	1.764851
Н	-2.123870	1.477756	0.236940
Н	-1.108031	0.214862	-2.764208
Н	0.359572	2.526724	-1.092050
Н	1.493790	2.145437	-0.114503
Н	1.108292	-0.212606	2.763726
$Fe_2S_2(H_2O)_4$			
Fe	0.257577	-0.256141	0.070816
Fe	2.106584	0.270263	1.574775
S	2.405917	0.011903	-0.681692
S	-0.041641	0.001504	2.327142

0	2.422189	2.080682	2.795251
0	3.886707	-0.936719	2.067388
0	-0.584911	0.941502	-1.579213
0	-0.984481	-2.079540	0.002154
Н	1.512396	2.156280	3.124635
Н	2.647380	2.916885	2.371088
Н	3.796261	-1.730477	2.606928
Н	4.090593	-1.220622	1.162090
Н	-1.101490	1.733422	-1.391538
Н	0.261755	1.227430	-1.957768
Н	-1.120977	-2.158333	0.959746
Н	-0.599937	-2.908721	-0.304516
FeH ₂ PO ₄	$(H_2O)_5^+$		
Fe	-0.169024	0.000675	-0.054911
Р	-0.078159	3.064937	0.479448
0	-0.027447	0.206389	2.106745
0	1.991386	0.549931	-0.349914

0	-0.02/44/	0.206389	2.106/45
0	1.991386	0.549931	-0.349914
0	0.019383	-0.676676	-2.138068
0	0.263956	-2.081088	0.306873
0	-0.673480	1.894527	-0.326046
0	-0.871327	4.416306	0.144781
0	1.359830	3.288377	-0.273824
0	-2.379637	-0.281081	-0.366905
0	0.077681	2.828678	1.953640
Η	-0.681469	-0.109927	2.739820
Η	0.069553	1.192405	2.267295
Η	-2.757963	0.608661	-0.398467
Η	-3.038326	-0.866154	0.024531
Н	0.600778	-2.461582	1.126701
Η	0.456563	-2.688962	-0.417590
Η	0.760117	-0.313964	-2.640755
Η	-0.758435	-0.645856	-2.709146
Н	2.693253	0.275497	0.252001
Н	2.067300	1.521268	-0.437713
Η	1.828140	4.095993	-0.024308
Н	-1.286776	4.816134	0.919529

$FeHPO_4(H_2O)_5, [FeH_2PO_4OH(H_2O)_3] \cdot H_2O$

Fe	-0.337893	0.106012	0.255407
Р	-0.288527	3.103394	0.749932
0	0.382426	0.318253	2.269365
0	1.194712	0.496035	-0.874913
0	0.990382	-1.887125	-2.075182
0	-0.279674	-2.077685	0.245497
0	-1.270686	1.969472	0.388930
0	-1.093669	4.500524	0.791672
0	0.683020	3.318504	-0.516594
0	-2.243436	-0.034789	-0.826690
0	0.458566	2.878967	2.048964
Н	0.023453	-0.041693	3.085257
Н	0.473864	1.325858	2.374244
Н	-2.508251	0.885766	-0.643756
Н	-2.992621	-0.620147	-0.679629
Н	0.285337	-2.531779	0.879601
Н	0.106236	-2.253837	-0.652546
Н	1.164158	-0.933061	-1.856347
Н	0.710840	-1.930400	-2.993824
Н	2.065360	0.489554	-0.462480
Н	0.964558	2.447403	-0.871942
Н	-0.850398	5.009270	1.573061

$FeH_4(PO_4)_2(H_2O)_4$

Fe	0.003758	0.001170	0.002735
Р	0.148926	-3.183234	0.365685
Р	-0.152621	3.185187	-0.354165
0	-0.262402	-0.507203	2.111910
0	2.174293	-0.478191	-0.087391
0	-0.335720	-2.007413	-0.462456
0	1.754535	-3.266766	-0.015094
0	-0.400645	-4.582644	-0.241782
0	-0.051502	-3.144668	1.858907
0	0.281579	0.513455	-2.109511

0	-2.168676	0.476961	0.058120
0	0.329886	2.009043	0.474752
0	-1.759544	3.267339	0.020244
0	0.392186	4.585070	0.256255
0	0.054360	3.147321	-1.846629
Н	2.292714	-1.442190	-0.183607
Н	2.529310	-0.072712	-0.887196
Н	-0.171876	-1.484697	2.273239
Н	0.256227	-0.022135	2.761271
Н	2.222598	-3.954127	0.474266
Н	-1.006577	-5.008477	0.375506
Н	-2.294019	1.437532	0.176396
Н	-2.562259	0.044966	0.824981
Н	0.186176	1.492138	-2.263418
Η	-0.242653	0.032950	-2.757850
Η	-2.226132	3.953880	-0.471691
Н	1.008077	5.007443	-0.353478

 $FeH_3(PO_4)_2(H_2O)_4^-$

Fe	0.087042	-0.000140	-0.054055
Р	0.406349	-2.991807	0.382882
Р	-0.363186	3.163468	-0.556261
0	-0.562529	-0.474527	2.078171
0	2.277269	-0.387927	-0.036967
0	-0.246937	-1.866404	-0.513873
0	1.924739	-2.940016	0.296215
0	-0.065638	-4.432752	-0.262146
0	-0.226721	-2.912980	1.792762
0	0.366034	0.422777	-2.299946
0	-2.323633	0.451562	0.176184
0	0.264322	2.032157	0.255851
0	-1.844289	3.515559	0.069827
0	0.443988	4.528513	-0.131271
0	-0.519591	2.992612	-2.031869
Н	2.296637	-1.396358	0.105014
Н	2.527718	-0.255113	-0.957701

-0.457048	-1.544748	2.064726
0.010722	-0.142578	2.774943
-0.975158	-4.602197	0.005111
-2.490420	1.402936	0.211830
-2.138937	0.174290	1.090234
-0.029543	1.309552	-2.462070
-0.206600	-0.246254	-2.690914
1.347936	4.301436	0.114366
-1.768026	3.826070	0.979712
	-0.457048 0.010722 -0.975158 -2.490420 -2.138937 -0.029543 -0.206600 1.347936 -1.768026	-0.457048-1.5447480.010722-0.142578-0.975158-4.602197-2.4904201.402936-2.1389370.174290-0.0295431.309552-0.206600-0.2462541.3479364.301436-1.7680263.826070

FeH(cit)(H₂O)₃

Fe	1.269537	-0.258146	0.089848
0	-0.386824	-0.182106	1.240771
0	-2.518824	-0.111273	1.909262
0	0.534416	1.014547	-1.400870
0	0.209259	2.872675	-0.201385
0	-1.689373	-3.159765	-0.627599
0	0.250903	-2.132176	-0.690410
0	-3.559045	0.114980	-0.414104
0	2.953450	-0.241390	-1.317467
0	1.881110	-1.607955	1.742501
0	2.228948	1.551273	0.891109
С	-2.145677	0.195572	-0.429450
С	-1.728640	1.627423	-0.886517
С	-1.648450	-0.872339	-1.480307
С	-1.653989	-0.057317	1.043146
С	-0.950433	-2.080386	-0.912428
С	-0.224575	1.897134	-0.830655
Н	-2.241829	2.351693	-0.254308
Η	-2.086407	1.761926	-1.912576
Н	-0.938048	-0.404490	-2.157620
Н	-2.538272	-1.171183	-2.037002
Н	3.741068	0.195783	-0.971294
Н	2.527084	0.405113	-1.906219
Н	1.074042	-1.462191	2.264017
Н	1.856477	-2.535802	1.477551

Н	2.245627	1.683813	1.845070
Н	1.620035	2.251459	0.513184
Н	-3.790751	0.108509	0.535203
Н	-2.623171	-2.984115	-0.807488

Fe(cit)(H₂O)₃⁻

Fe	1.292308	-0.117141	0.133833
0	-0.366594	-0.011088	1.409565
0	-2.533632	-0.041979	1.959919
0	0.681633	1.276247	-1.311981
0	0.187697	3.030769	-0.026408
0	-1.213299	-3.126627	-0.643346
0	0.598517	-1.818134	-0.673719
0	-3.464261	0.215427	-0.399934
0	2.843891	-0.450795	-1.594710
0	1.741090	-1.610188	1.890151
0	2.146624	1.629865	1.141391
С	-2.039555	0.250325	-0.365065
С	-1.638276	1.691371	-0.827430
С	-1.515905	-0.850185	-1.365016
С	-1.620556	0.039279	1.135404
С	-0.690492	-2.039067	-0.837722
С	-0.160449	2.055025	-0.720119
Н	-2.214575	2.403057	-0.236279
Н	-1.954598	1.782753	-1.872200
Н	-0.918353	-0.360249	-2.134882
Н	-2.414671	-1.264337	-1.819373
Н	2.425289	0.297225	-2.048537
Н	2.299673	-1.217992	-1.835073
Н	0.878897	-1.321616	2.243176
Н	1.505078	-2.345154	1.301762
Н	1.876046	1.607376	2.066473
Н	1.586072	2.345757	0.725077
Н	-3.709948	0.101338	0.538999

FeH(cit)₂³⁻

Fe	-0.251682	-0.259662	0.489097
0	1.796340	-0.094579	-1.251012
0	3.366514	1.523342	-1.230135
0	0.916022	-1.671448	1.373704
0	2.413983	-2.978514	2.398633
0	2.262284	2.049444	3.080993
0	1.144062	1.185469	1.326899
0	4.987350	-0.260132	-0.024364
0	-2.010599	0.404679	1.309215
0	-4.206721	0.590357	1.669126
0	-0.384499	0.993084	-1.292642
0	-1.203343	1.743370	-3.234605
0	-2.687825	-2.809497	-2.040803
0	-1.388514	-1.665698	-0.618336
0	-4.878041	1.144768	-0.687851
С	3.576140	-0.247243	0.371826
С	3.194624	-1.735513	0.522173
С	3.470331	0.574791	1.683091
С	2.867146	0.500080	-0.780333
С	2.157427	1.325735	2.071097
С	2.084187	-2.150193	1.528633
С	-3.513588	0.641325	-0.620059
С	-2.682042	1.730088	-1.356820
С	-3.553549	-0.739196	-1.327658
С	-3.202488	0.525371	0.917844
С	-2.424016	-1.816067	-1.327068
С	-1.329347	1.443249	-2.038039
Н	2.917359	-2.132326	-0.458587
Н	4.104069	-2.242284	0.848363
Н	3.740928	-0.077332	2.519224
Н	4.239162	1.351726	1.635525
Н	5.104321	0.597938	-0.458232
Н	-2.509087	2.552637	-0.652530
Н	-3.338220	2.106469	-2.142788
Н	-3.797821	-0.551010	-2.378649
Н	-4.426910	-1.256057	-0.915512
Н	-5.158858	1.033341	0.245686
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Н	1.046708	0.522849	-1.594775
Fe(cit) ⁴⁻			
Fe	-0.000039	-0.000247	-0.000155
0	1.885867	-0.050911	-1.118802
0	3.990389	-0.353559	-1.830187
0	0.790434	-1.745435	1.097688
0	1.847505	-3.334686	2.282121
0	2.318697	2.082475	3.057594
0	1.037826	1.177295	1.455078
0	5.102813	-0.539206	0.387058
0	-1.885822	0.049366	1.118675
0	-3.990362	0.351354	1.830245
0	-0.790116	1.745372	-1.097413
0	-1.847134	3.333712	-2.283101
0	-2.319215	-2.082474	-3.057992
0	-1.038184	-1.1771	-1.455713
0	-5.102719	0.539612	-0.386765
С	3.645747	-0.436959	0.540271
С	3.248895	-1.781132	1.203685
С	3.46567	0.81847	1.428891
С	3.096209	-0.265166	-0.927156
С	2.138441	1.374733	2.029283
С	1.815859	-2.29926	1.556174
С	-3.645691	0.437042	-0.540121
С	-3.248544	1.781689	-1.202391
С	-3.466004	-0.817638	-1.429878
С	-3.096155	0.263812	0.927159
С	-2.138831	-1.374533	-2.029847
С	-1.815549	2.298958	-1.5562
Н	3.669505	-2.564131	0.56033
Н	3.824522	-1.838557	2.13506
Н	4.139544	0.675932	2.279885
Н	3.886434	1.658806	0.860139
Н	5.156305	-0.522278	-0.598597

Н	-3.667572	2.564297	-0.557498
Н	-3.825395	1.840693	-2.132888
Н	-4.139158	-0.673733	-2.281232
Н	-3.887923	-1.658168	-0.862275
Н	-5.156247	0.521548	0.598873
Fe(cit) ₂ OH ⁵⁻			
Fe	-0.757539	-1.111181	0.176199
0	1.175641	-1.535013	-0.513485
Ο	3.180802	-1.695194	-1.488592
Ο	3.427629	-3.243721	2.170168
0	5.487182	-2.553376	2.790323
Ο	1.432463	-0.035502	3.635196
Ο	0.10703	-0.703887	1.971413
Ο	3.83392	0.564958	-0.281846
Ο	-2.626277	-0.312838	0.925144
Ο	-4.36954	1.028559	1.360913
Ο	-0.522447	0.996243	-0.426297
Ο	-0.505576	3.044333	-1.335137
Ο	-2.543259	-1.136838	-3.736912
Ο	-1.538156	-1.282147	-1.740861
Ο	-4.945206	1.581245	-1.015482
Ο	-1.225097	-2.861719	0.663112
С	3.13496	-0.4544	0.540327
С	4.32734	-1.141404	1.285645
С	2.188456	0.406575	1.40077
С	2.419982	-1.331347	-0.556966
С	1.19631	-0.188016	2.423358
С	4.369645	-2.435033	2.173926
С	-3.644644	0.914415	-0.915722
С	-2.598202	1.976777	-1.342963
С	-3.752264	-0.306292	-1.862465
С	-3.510792	0.503188	0.59135
С	-2.493582	-0.942926	-2.499517
С	-1.071158	1.973953	-1.000143
Н	5.04816	-1.366199	0.490519

Н	4.788238	-0.344953	1.880266
Н	2.831589	1.092945	1.961745
Н	1.591249	0.991714	0.696413
Н	4.008071	0.058973	-1.092948
Н	-2.961169	2.925938	-0.937775
Н	-2.658481	2.079549	-2.433956
Н	-4.392525	0.00964	-2.689158
Н	-4.276082	-1.109579	-1.329568
Н	-5.159721	1.653081	-0.056333
Н	-0.675581	-3.108034	1.417057

3+			
$Fe(H_2O)_6$			
Fe	0.000665	0.001616	0.002405
0	2.059449	0.000436	0.001405
0	0.002334	2.059899	0.000385
0	0.001336	0.002406	2.060211
0	-2.058763	0.000551	0.001372
0	0.002385	-2.058314	0.000325
0	0.001293	0.002383	-2.058008
Н	2.640669	0.785328	0.000424
Н	2.639610	-0.785380	0.001320
Н	-2.639949	0.785451	0.002084
Н	-2.639061	-0.785136	0.000320
Н	0.001459	2.641222	0.785229
Н	0.002234	2.639994	-0.785507
Н	0.000746	-2.639681	0.785057
Н	0.001514	-2.638393	-0.785502
Н	0.786184	0.002297	2.641483
Н	-0.784561	0.000632	2.640411
Н	0.785999	0.001344	-2.639294
Н	-0.784543	0.002346	-2.638121
$\operatorname{FeCl}(H_2O)_{5}^{2+}$			
Fe	0.08745	0.00003	-0.00019
Cl	2.23693	0.00121	0.00136
0	-0.25426	-1.39619	1.58008
0	-0.08868	1.56805	1.38586
0	-0.25618	1.39628	-1.57975
0	-0.08593	-1.56810	-1.38639
0	-2.07999	-0.00116	-0.00012
Н	-0.06264	2.34120	-1.47626
Н	-0.06762	1.17643	-2.50568
Н	-0.06214	-1.17601	2.50507
Н	-0.06075	-2.34125	1.47640
Н	0.67175	-2.02243	-1.78849

Table S4. Optimized structure Cartesian coordinates of Fe(III) complexes.

Н	-0.89916	-1.96425	-1.73549
Н	-2.64810	-0.52184	0.58720
Н	-2.64866	0.51930	-0.58711
Н	0.66817	2.02336	1.78841
Н	-0.90264	1.96265	1.73499

 $\operatorname{FeCl}_2(\operatorname{H}_2\operatorname{O})_4^+$

Fe	-0.000194	0.000959	0.000898
Cl	2.238809	-0.009528	0.010218
Cl	-2.239100	0.011578	-0.008603
0	0.002701	1.624522	1.396117
0	-0.000750	-1.621850	-1.395337
0	-0.014419	-1.395095	1.623324
0	0.012496	1.397188	-1.621920
Н	0.788502	-1.870555	-1.894401
Н	-0.787365	-1.863683	-1.901837
Н	0.769824	-1.901307	1.873362
Н	-0.806109	-1.893138	1.866156
Н	0.803865	1.895396	-1.865556
Н	-0.772016	1.903274	-1.871170
Н	-0.785828	1.875021	1.895374
Н	0.790130	1.868184	1.900512

 $\text{FeSO}_4(\text{H}_2\text{O})_5^+$

	5		
Fe	0.809247	-0.058921	0.075939
S	-2.332699	0.056166	-0.287204
0	1.687785	1.408490	1.411783
0	1.232885	-1.434528	1.675553
0	2.910070	-0.199370	-0.520830
0	0.341153	-1.459374	-1.341406
0	0.437738	1.606630	-1.094606
0	-3.437375	0.026549	0.629788
0	-2.146595	-1.143741	-1.159063
0	-2.077758	1.337130	-1.007457
0	-0.905692	-0.048861	0.642478
Н	1.345133	2.301869	1.260913

Н	1.660125	1.248163	2.365443
Н	0.704780	-1.666031	-2.210158
Н	-0.693759	-1.474112	-1.393475
Н	1.913263	-2.118453	1.723163
Н	0.419712	-1.793949	2.061029
Н	0.842384	1.874342	-1.928460
Н	-0.586428	1.662322	-1.177875
Н	3.544248	0.428462	-0.149223
Н	3.381500	-0.768396	-1.141952

 $\text{FeOH(H}_2\text{O})_5^{2+}$

0.150273	-0.018595	-0.267140
-1.364489	1.411791	-0.815210
-1.136793	-1.576563	-0.867882
1.088840	-0.039500	-1.750483
1.268477	1.537826	0.611615
-1.003796	0.007618	1.552843
1.275204	-1.439309	0.897696
2.017672	1.967959	0.171437
1.149996	1.949788	1.480385
-2.046227	1.203829	-1.472169
-1.159289	2.353061	-0.922408
-1.063384	-2.024192	-1.724791
-1.878616	-1.968460	-0.383494
-1.794377	0.543096	1.714272
-0.818575	-0.511766	2.348925
1.273046	-2.385358	0.686468
2.162804	-1.234606	1.229140
1.607076	-0.050170	-2.567113
	0.150273 -1.364489 -1.136793 1.088840 1.268477 -1.003796 1.275204 2.017672 1.149996 -2.046227 -1.159289 -1.063384 -1.878616 -1.794377 -0.818575 1.273046 2.162804 1.607076	$\begin{array}{cccccccccccccccccccccccccccccccccccc$

 $\operatorname{Fe(OH)}_2(\operatorname{H}_2O)_4^+$

Fe	0.005429	0.018578	-0.006113
0	-1.111523	-1.731989	0.699426
0	-1.496417	1.357976	0.805143
0	1.051701	0.017089	1.498451
0	1.531302	-1.257942	-0.903065

0	-1.065723	0.054330	-1.493574
0	1.153855	1.737572	-0.644173
Н	2.345329	-1.254441	-0.381827
Η	1.773459	-1.242147	-1.837174
Η	-1.718717	-2.154288	0.079131
Η	-0.748146	-2.414094	1.276975
Н	-1.723797	1.479871	1.734428
Н	-2.302994	1.389699	0.273643
Н	-1.220666	-0.329424	-2.359097
Н	0.768124	2.407644	-1.222107
Н	1.636553	2.180834	0.066090
Н	1.121904	-0.259562	2.414296

 $Fe(OH)_3(H_2O)_3$

Fe	-0.098495	0.172435	0.099263
0	0.220292	0.380546	1.953476
0	2.158162	-0.224161	0.406825
0	-0.101343	-2.139330	0.094166
0	0.534807	-0.217417	-2.089375
0	0.092354	1.877492	-0.700115
0	-1.907240	-0.347253	-0.104780
Н	-0.518415	0.417089	2.567145
Н	-1.068674	-2.142091	0.191948
Н	0.277853	-2.504865	0.900591
Н	2.706605	0.373776	-0.112148
Н	2.102034	0.131035	1.310076
Н	0.121608	2.665004	-0.150530
Н	-0.106890	-0.756280	-2.564225
Н	0.363849	0.713725	-2.310305
Н	-2.566883	0.297989	-0.372221

FeHCO₃(H₂O)₄²⁺

Fe	-0.131045	-0.099925	-0.000627
0	1.762920	0.155316	0.713852
0	2.930076	1.956064	-0.060055
0	0.853814	1.491978	-0.714309

0	-1.275980	0.762911	1.478961
0	0.168719	-1.504302	-1.477679
0	-0.030440	-1.754095	1.329780
0	-1.600841	0.676194	-1.327560
С	1.888916	1.209849	-0.008019
Н	-0.661027	-2.462000	1.535529
Н	0.777979	-1.907644	1.845518
Н	-1.786491	1.582178	1.377945
Н	-1.228994	0.548760	2.424077
Н	-2.507989	0.402973	-1.535653
Н	-1.397701	1.477228	-1.838683
Н	0.673094	-2.327429	-1.380020
Н	0.018499	-1.350769	-2.424162
Н	3.655271	1.651049	0.515518

 $\operatorname{FeCO}_{3}(\operatorname{H_{2}O})_{4}^{+}$

Fe	-0.130866	-0.087524	0.000878
0	1.683426	0.043237	0.625923
0	2.877204	1.885231	-0.001922
0	0.711528	1.523472	-0.627294
0	-1.377693	0.791295	1.464008
0	0.174244	-1.581193	-1.462793
0	-0.158576	-1.698073	1.520388
0	-1.621365	0.528626	-1.518248
С	1.882801	1.232933	-0.001253
Н	-0.519426	-2.594057	1.502573
Н	0.727018	-1.739854	1.913874
Н	-1.385033	1.754809	1.554786
Н	-1.404721	0.405495	2.350546
Н	-2.586890	0.555579	-1.499434
Н	-1.306547	1.357296	-1.912041
Н	1.054840	-1.971643	-1.556436
Н	-0.192555	-1.450386	-2.348180
FeHS(H ₂ O) ₅ ²⁺			
Fe	0.082822	-0.024258	0.013251

S	2.382431	-0.225020	-0.138489
0	0.023681	1.579188	-1.407883
0	-0.401761	-1.391133	-1.556433
0	-0.370906	-1.549542	1.431171
0	0.044788	1.409920	1.595590
0	-2.133847	0.309399	0.023542
Н	-0.458146	-2.493394	1.231399
Н	-0.296104	-1.462274	2.393017
Н	0.104001	1.462210	-2.366381
Н	0.271644	2.495184	-1.212073
Н	0.800914	1.790046	2.066411
Н	-0.757546	1.822318	1.947848
Н	-2.599780	0.976838	-0.499789
Н	-2.801014	-0.164737	0.539476
Н	0.209970	-1.911132	-2.097984
Н	-1.299304	-1.571366	-1.872341
Н	2.604480	-1.237301	0.734519

FeH₃PO₄(H₂O)₅³⁺

2 0		
-0.152437	0.027702	0.088825
-0.115691	3.358669	0.478663
-0.045277	-0.393104	2.145923
1.877005	0.416061	0.028404
-0.199052	-0.065573	-2.007714
0.273274	-2.020443	-0.078643
-0.484837	1.869401	0.255877
-0.641367	4.310383	-0.629634
1.456826	3.363838	0.336700
-2.170861	-0.438507	0.117250
-0.630763	3.690050	1.908681
-0.813420	-0.489773	2.734254
0.736865	-0.283787	2.712691
-2.895017	0.205738	0.201200
-2.568743	-1.325316	0.066606
0.411227	-2.638724	0.659343
0.376293	-2.519875	-0.906879
	-0.152437 -0.115691 -0.045277 1.877005 -0.199052 0.273274 -0.484837 -0.641367 1.456826 -2.170861 -0.630763 -0.813420 0.736865 -2.895017 -2.568743 0.411227 0.376293	-0.152437 0.027702 -0.115691 3.358669 -0.045277 -0.393104 1.877005 0.416061 -0.199052 -0.065573 0.273274 -2.020443 -0.484837 1.869401 -0.641367 4.310383 1.456826 3.363838 -2.170861 -0.438507 -0.630763 3.690050 -0.813420 -0.489773 0.736865 -0.283787 -2.895017 0.205738 -2.568743 -1.325316 0.411227 -2.638724 0.376293 -2.519875

Н	0.528641	0.179460	-2.603993
Н	-1.004353	-0.128459	-2.548974
Н	2.623967	-0.202007	-0.051192
Н	2.226001	1.325831	0.091391
Н	1.945354	4.199462	0.214821
Н	-1.516231	4.737807	-0.636816
Н	-0.574371	4.566699	2.330600

 $FeH_2PO_4(H_2O)_5^{2+}$

Fe	-0.179799	0.101945	-0.050992
Р	-0.010111	3.181871	0.563853
0	-0.200945	0.348816	1.961763
0	1.876551	0.497350	-0.309357
0	-0.081798	-0.520299	-2.088900
0	0.272426	-1.963784	0.272275
0	-0.488642	1.901565	-0.268270
0	-0.726911	4.466792	0.018265
0	1.492027	3.291805	0.004355
0	-2.224049	-0.394106	-0.259998
0	-0.149543	2.881802	2.020746
Н	-0.360563	-0.203083	2.739225
Н	-0.197019	1.344632	2.221165
Н	-2.897302	0.303622	-0.268906
Н	-2.663593	-1.227680	-0.035432
Н	0.494224	-2.445586	1.082558
Н	0.391609	-2.572361	-0.472671
Н	0.671372	-0.260247	-2.641264
Н	-0.864479	-0.550291	-2.659726
Н	2.632560	-0.021426	0.001370
Н	2.127086	1.444576	-0.312772
Н	1.954286	4.140546	0.092675
Н	-1.401495	4.888031	0.573758

 $FeHPO_4(H_2O)_5^+$

Fe	-0.032490	-0.034990	0.119867
Р	0.003018	3.019679	0.328898
0	-0.098746	0.166963	1.974439

0	1.753019	0.460423	-0.251411
0	0.005466	-0.661324	-2.013718
0	0.088314	-2.209653	0.252187
0	-0.719650	1.812242	-0.285370
0	-0.654738	4.406802	-0.008077
0	1.451328	3.154026	-0.260601
0	-2.187515	-0.450490	-0.295551
0	-0.002913	2.905413	1.906191
Н	-0.611462	-0.314477	2.632005
Н	-0.006845	1.961545	2.212846
Н	-2.709511	0.361854	-0.247822
Н	-2.723187	-1.176094	0.047126
Н	0.525183	-2.665894	0.982481
Н	0.311134	-2.669830	-0.568364
Н	0.796497	-0.291339	-2.430626
Н	-0.748064	-0.485241	-2.591626
Н	2.469116	0.166544	0.324400
Н	1.841674	2.252517	-0.398859
Н	-1.238020	4.785265	0.662626

 $FeH_5(PO_4)_2(H_2O)_4^{2+}$

Fe	-0.014976	-0.109037	-0.273615
Р	0.189740	-3.108630	0.665674
Р	-0.255092	3.275053	-0.354083
0	-0.037619	-0.156523	1.780976
0	2.060128	-0.460257	-0.315265
0	-0.288049	-1.948263	-0.309026
0	1.728933	-3.221624	0.193807
0	-0.453006	-4.472783	0.203849
0	-0.025980	-2.714501	2.088841
0	-0.044440	-0.110115	-2.398181
0	-2.097872	0.284457	-0.300783
0	0.284623	1.868214	-0.313942
0	-1.759849	3.242157	0.145520
0	0.424805	4.311549	0.613699
0	-0.156450	3.751728	-1.854529

Н	2.358088	-1.388825	-0.251629
Н	2.788564	0.100229	-0.610923
Н	-0.061093	-1.103163	2.147945
Н	0.446631	0.416878	2.389007
Н	2.215591	-4.018310	0.454932
Н	-1.115280	-4.868071	0.790433
Н	-2.497485	1.092919	0.054327
Н	-2.751718	-0.427527	-0.265703
Н	-0.064897	0.630861	-3.018004
Н	-0.141293	-0.934234	-2.895869
Н	-2.153420	4.040338	0.532845
Н	1.339918	4.603004	0.480430
Н	-0.434412	4.644566	-2.112419

 $FeH_4(PO_4)_2(H_2O)_4^+$

Fe	-0.000024	0.000116	-0.000016
Р	0.171448	-3.126351	0.430507
Р	-0.171106	3.126422	-0.430636
0	-0.107395	-0.424297	2.037871
0	2.091766	-0.418252	0.020062
0	-0.31494	-1.8558	-0.33115
0	1.716206	-3.202239	-0.072325
0	-0.479218	-4.425653	-0.223055
0	0.007676	-3.033087	1.914368
0	0.106401	0.424251	-2.037929
0	-2.091883	0.41816	-0.019881
0	0.315725	1.855872	0.330759
0	-1.715672	3.202189	0.072944
0	0.480058	4.425558	0.222752
0	-0.00814	3.033306	-1.914576
Н	2.342934	-1.343065	-0.16573
Н	2.736029	0.178286	-0.379839
Н	-0.055393	-1.398938	2.271256
Н	0.306359	0.119655	2.718571
Н	2.198981	-4.000153	0.183802
Н	-1.136739	-4.865229	0.332553

Н	-2.342745	1.34302	0.166159
Н	-2.735548	-0.178246	0.381226
Н	0.054663	1.398773	-2.271624
Н	-0.306985	-0.119942	-2.718651
Н	-2.198767	3.999765	-0.183646
Н	1.13631	4.865879	-0.333768

 $FeH_3(PO_4)_2(H_2O)_4$

	2 .		
Fe	-0.249926	-0.020554	0.072560
Р	0.208167	-3.172794	0.435980
Р	-0.212896	3.026506	-0.371405
0	-0.136553	-0.472443	2.151315
0	2.023286	-0.432943	0.154760
0	-0.395501	-1.963269	-0.291515
0	1.699608	-3.278093	-0.221384
0	-0.479602	-4.532289	-0.097808
0	0.226865	-3.108142	1.934932
0	0.251491	0.328977	-2.009186
0	-2.046755	0.473124	-0.082747
0	0.355875	1.793867	0.406321
0	-1.740205	3.193103	0.063807
0	0.457130	4.350856	0.244860
0	0.025270	2.921028	-1.857639
Н	2.264699	-1.283293	-0.246387
Н	2.438756	0.265319	-0.364506
Н	0.039619	-1.436039	2.323163
Н	0.521402	0.056306	2.616777
Н	2.225499	-4.003205	0.138370
Н	-1.096876	-4.896632	0.547959
Н	-2.175719	2.310139	0.060494
Н	-2.739543	-0.189017	-0.171089
Н	0.168270	1.322355	-2.168761
Н	-0.314544	-0.127644	-2.641016
Н	0.880948	4.875011	-0.444204

FeH₇(PO₄)₃(H₂O)₃⁺ Fe -0.478399 -0.254569 -0.546881

Р	1.964221	-2.298021	-0.545740
Р	-3.343060	-1.911226	-0.547349
Р	-1.265793	2.675749	0.352334
0	-1.294137	-0.331327	1.519390
0	0.252230	-0.154817	-2.510849
0	1.236383	0.574655	0.389261
0	0.457082	-1.926657	-0.398210
0	2.169858	-3.822441	-0.134158
0	2.127753	-2.234616	-2.179530
0	-2.198006	-1.093590	-1.090255
0	-4.769289	-1.400899	-1.015071
0	-3.163465	-3.408558	-1.042728
0	-1.079514	1.565762	-0.715686
0	-1.803004	2.152096	1.656770
0	0.218286	3.307068	0.445984
0	-2.125754	3.879565	-0.243001
0	-3.462752	-1.895142	1.016041
0	2.923453	-1.407337	0.162571
Н	0.936281	-0.796649	-2.781207
Н	0.320376	0.661241	-3.020360
Н	1.426732	1.524534	0.400155
Н	2.063291	0.014874	0.424176
Н	-0.616489	-0.599085	2.156309
Н	-1.546572	0.645461	1.738425
Н	1.554692	-4.453770	-0.529996
Н	3.036110	-2.366385	-2.486075
Н	-4.923689	-1.357478	-1.968757
Н	-3.804980	-4.056191	-0.717915
Н	0.298492	4.110297	0.978163
Н	-3.041564	3.906581	0.063990
Η	-2.749926	-1.364910	1.447187

 $\operatorname{FeH_6(PO_4)_3(H_2O)_3}$

Fe	-0.465389	0.020141	-0.137766
Р	1.498694	-2.680298	-0.186206
Р	-2.672477	-2.293855	-0.567857

Р	-1.233906	3.213215	0.222161
0	-0.430778	0.428122	1.986870
Ο	-0.329376	-0.404101	-2.211580
Ο	1.429858	1.084738	-0.329333
Ο	0.646977	-1.453040	0.280919
Ο	0.665894	-3.982244	0.168955
Ο	1.352418	-2.504007	-1.844278
Ο	-2.166830	-0.922941	-0.069912
Ο	-4.227283	-2.396339	-0.167933
Ο	-2.658776	-2.066662	-2.202593
Ο	-1.245426	1.785066	-0.334412
Ο	-1.490411	2.946528	1.831105
Ο	0.372508	3.602339	0.179704
Ο	-2.075952	4.302854	-0.306657
Ο	-1.970181	-3.549599	-0.194790
Ο	2.918897	-2.744730	0.209853
Н	0.357595	-1.099945	-2.321335
Н	-1.161998	-0.803080	-2.523328
Н	1.382745	2.045408	-0.165733
Н	2.142523	0.700492	0.195725
Н	-0.560229	-0.323313	2.576149
Н	-0.961413	1.191055	2.284465
Н	-0.315906	-3.889576	0.058242
Н	2.098217	-2.908879	-2.303481
Н	-4.694751	-1.555165	-0.245251
Н	-2.766313	-2.896105	-2.685830
Н	0.511403	4.535721	-0.021280
Н	-1.881380	3.707562	2.276388
Fe(cit)(H ₂ O) ₃			
Fe	1.069944	-0.211234	-0.014539
Ο	-0.403464	-0.034044	1.233597
0	-2.503289	-0.195237	1.961374
0	0.454523	1.141178	-1.230282
0	0.112230	3.014865	-0.077092
0	-1.377440	-3.085918	-0.844896

0	0.394372	-1.732680	-0.887835
0	-3.641073	0.266126	-0.326712
0	2.961756	-0.662048	-1.002685
0	1.874141	-1.282278	1.762557
0	2.225964	1.495852	0.793658
С	-2.228295	0.278751	-0.375663
С	-1.817913	1.738108	-0.772403
С	-1.781932	-0.788116	-1.455354
С	-1.706524	-0.021151	1.061616
С	-0.923325	-1.981075	-1.019535
С	-0.336268	2.044149	-0.677190
Н	-2.359025	2.433919	-0.131701
Н	-2.154075	1.899360	-1.801525
Н	-1.252375	-0.279480	-2.261399
Н	-2.711304	-1.199910	-1.843397
Н	3.491918	0.112459	-1.229202
Н	2.810196	-1.172094	-1.809397
Н	1.057739	-1.328693	2.286100
Н	2.160637	-2.190799	1.606801
Н	2.397245	1.539327	1.741432
Н	1.668258	2.277748	0.549677
Н	-3.874977	0.074175	0.597953

 $Fe(cit)OH(H_2O)_2^{-}$

Fe	1.170506	-0.293435	0.423406
0	-0.519085	0.100923	1.378565
0	-2.691790	-0.001837	1.864051
0	0.618392	1.083837	-1.083351
0	0.174703	3.046201	-0.114284
0	-1.359221	-3.145603	-0.948094
0	0.370551	-1.816293	-0.497519
0	-3.573193	0.313614	-0.513718
Ο	2.682234	-0.530438	-1.318104
0	2.232037	-1.088435	1.731081
0	2.117545	1.593716	1.123433
С	-2.149820	0.306800	-0.445443

С	-1.679202	1.726486	-0.885843
С	-1.638942	-0.810099	-1.426737
С	-1.769017	0.097639	1.058717
С	-0.867157	-2.030479	-0.901095
С	-0.195848	2.009723	-0.683103
Н	-2.255436	2.468669	-0.332936
Н	-1.925145	1.835855	-1.948079
Н	-0.997260	-0.343314	-2.176109
Н	-2.533660	-1.192522	-1.914617
Н	2.253815	0.207831	-1.782023
Н	2.243885	-1.327652	-1.648389
Н	1.950800	-1.974337	1.981833
Н	2.084544	1.590144	2.085995
Н	1.553040	2.332193	0.786363
Н	-3.850267	0.182876	0.413156
FeH(cit) ^{2–}			
Fe	-0.218898	-0.110580	0.359954
Ο	1.765294	-0.219397	-1.317881
Ο	3.447214	1.274253	-1.425941
Ο	0.749466	-1.448266	1.390620
Ο	2.070164	-2.834026	2.526101
Ο	1.987615	1.998290	2.937107
Ο	1.059861	1.250036	1.040021
0	4.916741	-0.439415	0.070600
Ο	-1.826659	0.411239	1.316816
О	-3.997531	0.424852	1.793905
О	-0.400664	1.042806	-1.308121
О	-1.264829	2.017151	-3.110172
Ο	-2.158796	-2.441476	-2.404158
Ο	-1.139277	-1.512506	-0.655569
Ο	-4.837047	0.969823	-0.535494
С	3.500071	-0.327546	0.360882
С	3.036965	-1.778129	0.622795
С	3.350259	0.618516	1.580876
С	2.890386	0.340642	-0.885088

С	2.015473	1.332714	1.908357
С	1.877683	-2.040900	1.607749
С	-3.444747	0.595628	-0.526274
С	-2.720322	1.772875	-1.237166
С	-3.385599	-0.754601	-1.289865
С	-3.064621	0.453246	0.984543
С	-2.120987	-1.619011	-1.492386
С	-1.381167	1.587722	-1.965810
Н	2.779244	-2.253664	-0.327387
Н	3.903357	-2.297714	1.031346
Н	3.665649	0.074780	2.474845
Н	4.072008	1.429663	1.442601
Н	5.135596	0.364220	-0.422206
Н	-2.578001	2.578301	-0.507709
Н	-3.422181	2.138538	-1.986234
Н	-3.800634	-0.571836	-2.283857
Н	-4.099578	-1.417433	-0.787340
Н	-5.105362	0.825075	0.392247
Н	1.185662	0.419494	-1.812378
$\operatorname{Fe}(\operatorname{cit})_2^{3-}$			
Fe	-0.075745	-0.333239	-0.031024
0	0.130439	-0.579650	2.004182
0	0.890147	-0.937747	4.068913
0	1.623894	0.745941	-0.072907
0	3.713199	1.522550	-0.100047
0	2.564426	-3.578061	-0.596984
0	1.010615	-1.999128	-0.349121
0	3.275573	-1.436167	3.502015
0	-0.281731	-0.086576	-2.066190
0	-1.037043	0.291166	-4.129037
0	-1.775371	-1.412679	0.014701
0	-3.858393	-2.206235	0.029889
0	-2.728795	2.898000	0.539535
0	-1.162285	1.332985	0.283350
0	-3.429190	0.759651	-3.565104

С	2.495300	-1.143471	2.311104
С	3.188147	0.101179	1.706183
С	2.606918	-2.433980	1.464329
С	1.042469	-0.859945	2.826253
С	2.024079	-2.664260	0.048242
С	2.810382	0.823556	0.389335
С	-2.648704	0.468014	-2.374071
С	-3.331905	-0.785554	-1.776795
С	-2.773290	1.753320	-1.521084
С	-1.192611	0.198915	-2.887771
С	-2.184255	1.988534	-0.108199
С	-2.957791	-1.501675	-0.455706
Н	3.140956	0.882033	2.475327
Н	4.247486	-0.154922	1.615063
Н	3.675153	-2.655522	1.388729
Н	2.182885	-3.240323	2.075582
Н	2.587650	-1.340721	4.197296
Н	-3.267892	-1.565863	-2.545299
Н	-4.394977	-0.542090	-1.695621
Н	-3.843998	1.960043	-1.438145
Н	-2.364427	2.567485	-2.132267
Н	-2.738069	0.678408	-4.259013
Fe(cit) ₂ OH ⁴⁻			
Fe	-0.510640	-0.870021	0.111101
0	1.357832	-1.168550	-0.743071
0	3.388151	-1.689664	-1.522295
0	2.364879	-3.125576	1.805892
0	4.287977	-3.225984	2.983388
0	1.658834	0.882178	3.322883
0	0.327813	-0.030346	1.780861
0	4.456797	0.249770	-0.185741
0	-2.455853	-0.424627	0.934820
0	-4.607290	0.052754	1.317548
0	-0.547886	1.114498	-0.758798
0	-0.876746	3.053824	-1.832626

0	-2.749472	-1.660948	-3.532956
0	-1.566256	-1.369032	-1.656293
0	-5.068820	1.342295	-0.766305
0	-0.508387	-2.601897	0.803804
С	3.345466	-0.438069	0.503366
С	4.045690	-1.365498	1.543446
С	2.524836	0.725971	1.096566
С	2.614206	-1.190976	-0.665003
С	1.433979	0.482921	2.160991
С	3.476268	-2.688105	2.164449
С	-3.710574	0.795232	-0.773428
С	-2.807398	2.041844	-0.926236
С	-3.698028	-0.175561	-1.978129
С	-3.554842	0.069087	0.609630
С	-2.547384	-1.132747	-2.411484
С	-1.276945	2.038570	-1.210177
Н	4.975279	-1.684992	1.061188
Н	4.352411	-0.732479	2.383831
Н	3.254351	1.399801	1.554794
Н	2.047256	1.246284	0.260941
Н	4.584593	-0.349775	-0.945007
Н	-2.927916	2.633736	-0.009127
Н	-3.258703	2.636085	-1.726200
Н	-3.933955	0.426535	-2.862866
Н	-4.563748	-0.834885	-1.840675
Н	-5.382448	0.985693	0.097391
Н	0.378036	-2.814858	1.147242

CuCl(H ₂ O)			
Cu	0.036942	0.074314	0.043155
Cl	-2.057940	-0.063980	0.074830
0	2.004689	0.205762	0.009429
Н	2.502159	-0.106062	0.774133
Н	2.473097	-0.085036	-0.781548
$CuCl_2^-$			
Cu	-0.000055	0.043433	0.043431
Cl	2.169412	-0.018947	0.082887
Cl	-2.169357	0.105513	0.003682
CuCl ₃ ²⁻			
Cu	-0.005126	0.506655	0.034208
Cl	2.106385	-0.653066	0.075348
Cl	-2.066920	-0.735974	0.128658
Cl	-0.054341	2.912384	-0.098214
CuHS(H ₂ O)			
Cu	-0.011636	-0.248242	-0.030432
S	2.118534	-0.441174	-0.041410
Ο	-1.999404	-0.132648	-0.070277
Н	2.342706	0.584117	0.810873
Н	-2.504543	-0.362548	0.717514
Н	-2.425131	0.628496	-0.481195
$Cu(HS)_2^-$			
Cu	0.000067	-0.279988	-0.000206
S	2.204089	-0.30569	-0.150469
S	-2.205099	-0.295192	0.127476
Н	2.446249	0.510525	0.900414
Н	-2.420237	0.760894	-0.689656
$Cu_2S(HS)_2^{2-}$			
Cu	0.133977	-0.122383	2.110868
Cu	-1.786975	-0.104253	-0.745506
S	0.277476	-0.331589	-0.061234

Table S5. Optimized structure Cartesian coordinates of Cu(I) complexes.

S	0.192536	0.089698	4.364498
S	-3.855727	0.102047	-1.642872
Н	-4.058614	1.398392	-1.311408
Н	-0.652741	-0.926301	4.657057

CuH ₂ PO ₄	$(H_2O)_4^+$		
Cu	-0.410131	0.052746	0.548220
Р	-0.084032	3.057843	0.729853
0	0.221424	0.346939	2.444874
0	1.468939	0.388902	-0.646826
0	-0.137745	-1.930471	0.822551
0	-0.903862	1.852526	0.192761
0	-0.924592	4.401601	0.547857
0	1.066116	3.223744	-0.403029
0	-1.653150	-0.413237	-1.054587
0	0.456009	2.863712	2.118343
Н	-0.370659	0.132506	3.177067
Н	0.413673	1.349287	2.502253
Н	-1.819901	0.438400	-1.487374
Н	-2.513067	-0.836960	-0.934075
Н	0.490533	-2.210974	1.501282
Н	-0.045072	-2.517917	0.060867
Н	2.343796	-0.013082	-0.614255
Н	1.595401	1.332399	-0.848165
Н	1.598948	4.028082	-0.340006
Н	-1.347796	4.724939	1.354077

Table S6. Optimized structure Cartesian coordinates of Cu(II) complexes.

CuHPO₄(H₂O)₄

Cu	-0.609161	-0.021907	0.630228
Р	-0.052556	2.883407	0.526304
0	0.256383	0.048185	2.332598
0	1.470221	0.161312	-0.581758
0	-0.079656	-2.023616	0.692530
0	-1.201017	1.811462	0.470717
0	-0.736930	4.340408	0.546184
0	0.931164	2.792667	-0.596546
0	-1.753077	-0.191718	-1.102412
0	0.588091	2.766441	2.001119
Н	-0.375042	0.071311	3.061054

0.637112	1.816222	2.277505
-2.012773	0.746492	-1.125878
-2.549743	-0.730557	-1.157156
0.407836	-1.877492	1.526101
0.584376	-2.123106	-0.004263
2.235072	0.161952	0.005023
1.373492	1.119191	-0.826213
-1.267907	4.492248	1.336555
	0.637112 -2.012773 -2.549743 0.407836 0.584376 2.235072 1.373492 -1.267907	0.6371121.816222-2.0127730.746492-2.549743-0.7305570.407836-1.8774920.584376-2.1231062.2350720.1619521.3734921.119191-1.2679074.492248

 $CuH_4(PO_4)_2(H_2O)_3$

Cu	-0.273859	0.104002	-0.253640
Р	0.504466	2.982905	0.800121
Р	0.602414	-2.783471	0.143895
0	-0.489817	-0.272879	1.918198
0	1.622372	0.259201	-1.018969
0	-0.390882	2.006601	0.033434
0	0.571624	-2.780402	1.642152
0	0.256045	4.440303	0.816929
0	0.558979	2.339280	2.315992
0	2.035209	2.653603	0.241518
0	-0.283416	-1.787349	-0.612701
0	2.105183	-2.452297	-0.441444
0	0.339396	-4.262572	-0.454818
0	-2.349055	0.272643	-0.623185
Н	2.134887	-0.559696	-0.904740
Н	2.099680	1.029065	-0.656543
Н	-0.081403	-1.137389	2.142085
Н	-0.075032	0.435887	2.432373
Н	-2.565838	1.137232	-0.247424
Н	-2.866330	-0.391897	-0.153521
Н	-0.115915	-4.821051	0.186098
Н	2.794965	-2.989856	-0.033604
Н	0.802631	2.996805	2.978233
Н	2.535042	3.467607	0.107901

 $CuH_3(PO_4)_2(H_2O)_3^-$

Cu	-0.263817	0.254663	-0.314626
Р	0.665632	3.071986	0.803683
Р	0.411803	-2.709366	0.043567
0	-0.740807	-0.141032	1.822244
0	1.664120	0.298258	-1.068340
0	-0.233759	2.189651	-0.040147
0	0.428349	-2.376808	1.540720
0	0.592952	4.557419	0.792640
0	0.508505	2.466645	2.330285
0	2.232208	2.612001	0.439490
0	-0.402024	-1.593129	-0.749634
0	1.984456	-2.282940	-0.515023
0	0.114106	-4.102374	-0.405346
0	-2.500440	-0.034709	-0.394230
Н	2.015361	-0.610207	-0.893520
Н	2.179547	0.961021	-0.580690
Н	-0.258720	-1.044283	1.884484
Н	-0.287788	0.519110	2.362533
Н	-2.526850	-0.026440	0.574812
Н	-2.244619	-0.951625	-0.613081
Н	2.642799	-2.827901	-0.072637
Н	0.856825	3.078140	2.988005
Н	2.749422	3.399247	0.237052

 $CuH_2(PO_4)_2(H_2O)_2^{2-}$

Cu	0.030985	-0.000223	-0.115202
Р	1.226935	-2.907117	-0.048313
Р	-1.168548	2.903396	-0.192081
0	1.945516	-0.045406	0.198185
0	-0.020981	-2.001970	-0.208990
0	1.059252	-4.386253	-0.143258
0	1.926588	-2.494434	1.381324
0	2.336136	-2.379254	-1.140754
0	0.080268	2.001257	-0.021194
0	-1.003888	4.383267	-0.104559
0	-2.280837	2.379124	0.899172

0	-1.862640	2.481951	-1.621386
0	-1.883823	0.040292	-0.428433
Н	2.177584	0.884644	0.280713
Н	2.067797	-1.521522	1.328070
Н	2.441804	-1.417365	-0.957342
Н	-2.112172	-0.891180	-0.503837
Н	-2.383528	1.416071	0.720975
Н	-2.006595	1.509824	-1.562774

$CuH_2(cit)(H_2O)_2^+$

Cu	1.232197	-0.207584	0.128561
0	-0.156048	0.035530	1.372068
0	-2.238743	-0.081455	2.117049
0	0.537981	1.454367	-1.095552
0	-0.358696	3.414719	-1.434891
0	-1.366748	-3.173514	-0.855924
0	0.405005	-1.928217	-0.589383
0	-3.434574	-0.031866	-0.125843
0	2.886464	-0.472149	-1.013311
0	2.415844	0.597985	1.629346
С	-2.050617	0.188680	-0.238995
С	-1.858875	1.667408	-0.685035
С	-1.552044	-0.784124	-1.369254
С	-1.441885	0.004859	1.203612
С	-0.784330	-1.995481	-0.906966
С	-0.470892	2.143945	-1.067250
Η	-2.244244	2.315636	0.109850
Η	-2.516446	1.830139	-1.547927
Η	-0.885494	-0.269246	-2.060599
Η	-2.448363	-1.072008	-1.919618
Η	3.143831	0.252708	-1.599276
Η	2.969117	-1.298456	-1.507855
Η	1.902336	0.662579	2.447348
Η	3.315941	0.330744	1.851024
Η	-3.626808	-0.061127	0.832255
Н	-2.300783	-3.132486	-1.108849

CuH(cit)(H₂O)₂

Cu	1.173822	-0.227277	0.090587
0	-0.218395	0.021607	1.356417
0	-2.318049	0.127551	2.064606
0	0.510312	1.038894	-1.606061
0	0.298816	2.845992	-0.306419
0	-1.448925	-3.252238	-0.626384
0	0.334728	-2.016437	-0.387865
0	-3.462808	0.131184	-0.210460
0	2.744783	-0.350840	-1.217404
0	2.129201	1.429783	0.896408
С	-2.056028	0.229213	-0.304370
С	-1.693327	1.633381	-0.887773
С	-1.607349	-0.908257	-1.300858
С	-1.496097	0.102729	1.160476
С	-0.843980	-2.072638	-0.731273
С	-0.188218	1.901721	-0.956220
Н	-2.167730	2.390452	-0.264131
Н	-2.129865	1.688998	-1.889395
Н	-0.944325	-0.470835	-2.049056
Н	-2.523610	-1.247324	-1.785394
Н	3.527061	0.084752	-0.854484
Н	2.326231	0.302145	-1.816922
Н	1.882810	1.501328	1.828145
Н	1.593914	2.155437	0.433743
Н	-3.654950	0.199661	0.744947
Н	-2.375011	-3.182515	-0.894108

$Cu(cit)(H_2O)_2^-$

Cu	1.192643	-0.241528	0.243424
0	-0.278277	0.021000	1.452377
0	-2.418747	0.164468	2.017834
0	0.600705	1.337709	-1.767604
0	0.383383	2.866966	-0.143598

0	-1.170138	-3.201642	-0.877502	
0	0.564916	-1.877338	-0.470847	
0	-3.401186	0.232489	-0.329081	
0	2.639305	-0.213249	-1.259252	
0	2.036128	1.386053	1.170240	
С	-1.974834	0.267998	-0.331568	
С	-1.586200	1.695002	-0.860384	
С	-1.489003	-0.859806	-1.310249	
С	-1.532304	0.122163	1.165540	
С	-0.690096	-2.079946	-0.821766	
С	-0.088711	2.013362	-0.940778	
Н	-2.072545	2.428512	-0.216723	
Н	-2.023255	1.765633	-1.860787	
Н	-0.878662	-0.393687	-2.086661	
Н	-2.398595	-1.248433	-1.764205	
Н	2.119297	0.516929	-1.691247	
Н	2.366938	-1.021594	-1.716181	
Н	1.530708	1.387029	1.995352	
Н	1.583718	2.121764	0.637576	
Н	-3.632426	0.233297	0.619067	
$Cu(cit)_2^{4-}$				
Cu	0.002093	-0.283947	-0.309252	
0	1.778911	-1.029903	-0.883764	
0	3.939473	-1.325145	-1.338886	
0	2.850677	-3.771515	1.298160	
0	4.667681	-3.938347	2.633581	
0	1.189856	-0.992232	3.545949	
0	0.204227	-0.777376	1.558809	
0	4.493279	0.339280	0.633697	
0	-1.006580	1.402687	0.065572	
0	-2.527140	2.878884	0.759781	
0	-2.384071	3.049091	-2.910327	
0	-4.535655	3.350302	-3.537692	
0	-2.615438	-0.905673	-3.368456	
Ο	-0.848266	-0.597192	-2.044888	

0	-4.315758	0.904849	0.500451
С	3.485874	-0.714263	0.903788
С	4.298201	-1.834127	1.616867
С	2.470478	0.030060	1.791421
С	2.999120	-1.087072	-0.544982
С	1.213798	-0.666605	2.350689
С	3.856287	-3.317599	1.873242
С	-3.370764	1.139184	-0.616885
С	-4.176064	2.022943	-1.612781
С	-3.068913	-0.292768	-1.096570
С	-2.183611	1.877182	0.094349
С	-2.114920	-0.584776	-2.279826
С	-3.610398	2.866923	-2.808390
Н	5.214128	-1.939394	1.024785
Н	4.611974	-1.421877	2.581912
Н	3.035801	0.393773	2.654466
Н	2.136718	0.904010	1.223281
Н	4.867737	0.023835	-0.205742
Н	-4.674599	2.771633	-0.987244
Н	-4.974443	1.390157	-2.015018
Н	-4.034819	-0.729355	-1.366927
Н	-2.697523	-0.836521	-0.221219
Н	-4.177345	1.700882	1.038049

$OI CU_2$.							
Service		Temperature (K)					
Species	273	298	323	373	473	573	
$\operatorname{Cu}_2^{0 a}$	1.026	0.864	0.738	0.556	0.348	0.238	
	$(\alpha, \beta, \alpha, \beta)$						

Table S7 Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁶⁵Cu-⁶³Cu of Cu₂⁰.

^{*a*} Atomic distance of Cu⁰-Cu⁰ was calculated to be 2.28Å.

Species			Tempera	ature (K)		
	273	298	323	373	473	573
CuHPO ₄ (H ₂ O) ₄	5.827	4.958	4.268	3.254	2.064	1.422

Table S8 Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁶⁵Cu-⁶³Cu of Cu(II) mono-hydrogen phosphate.

Species			Tempera	ature (K)		
	273	298	323	373	473	573
$\text{FeHS(H}_2\text{O})_5^{2+}$	6.502	5.507	4.722	3.580	2.255	1.547

Table S9 Logarithm of the reduced partition function, $\ln \beta$ (‰), for the pair ⁵⁶Fe-⁵⁴Fe of Fe(III) mono-hydrogensulfide.



Figure S1. Mole fractions of Fe(II) and Fe(III) species and Fe isotopic variations $({}^{56}\text{Fe}/{}^{54}\text{Fe})$ in seawater at pH = 8.2 and 298 K. Total concentration of Fe was set to 3.6×10^{-8} mol kg⁻¹. The ratio of concentrations was set to [Fe(II)]/[Fe(III)] = 4. Mole fractions of Fe(II) species were taken from Fig. 4. Mole fractions of Fe(III) hydroxides were estimated by using stability constants (Byrne and Kester, 1976).



Figure S2. Copper, Cu(I), isotopic variations (65 Cu/ 63 Cu, δ^{65} Cu) under hydrothermal conditions at 573 K. (a) The conditions for Broadlands fluid. (b) The conditions of high salinity at low pH. (c) The conditions of high sulfur concentration at high pH. Estimated mole fractions are from Mountain and Seward (1999), where the detailed conditions of pH, concentrations of Cl and S, and H₂ fugacity can be found.



Figure S3. Mole fractions of Fe(III) species and Fe isotopic variations (⁵⁶Fe/⁵⁴Fe) in a soil-plant system as a function of pH at 298 K. (a) Dissociation constants of citric acid and stability constants of orthophosphates and Fe(III) species were taken from the literature (Childs, 1970; Baes and Mesmer, 1976; Ciavatta and Iuliano, 1995; Königsberger et al., 2000). (b) δ^{56} Fe of Fe(III) phosphates, citrates, hydroxides, and hydrated Fe³⁺ ions relative to the bulk solution. Total concentrations of Fe, P, and citrate were set to 0.005, 0.05, and 0.01 mol kg⁻¹, respectively. Activity coefficients of related species were set to unity.



Figure S4. Mole fractions of Ni(II) species and Ni isotopic variations (60 Ni/ 58 Ni) in a soil-plant system as a function of pH at 298 K. (a) Dissociation constants of citric acid and stability constants of orthophosphates and Ni(II) species were taken from the literature (Childs, 1970; Baes and Mesmer, 1976; Taylor and Diebler, 1976; Hedwig et al., 1980). (b) δ^{60} Ni of Ni(II) phosphates, citrates, hydroxides, and hydrated Ni²⁺ ions relative to the bulk solution. Total concentrations of Ni, P, and citrate were set to 0.005, 0.1, and 0.01 mol kg⁻¹, respectively. Activity coefficients of related species were set to unity.


Figure S5. Mole fractions of Cu(II) species and Cu isotopic variations (65 Cu/ 63 Cu) in a soil-plant system as a function of pH at 298 K. (a) Dissociation constants of citric acid and stability constants of orthophosphates and Cu(II) species were taken from the literature (Childs, 1970; Baes and Mesmer, 1976; Petit-Ramel and Khalil, 1974; Ciavatta et al., 1993). (b) δ^{65} Cu of Cu(II) phosphates, citrates, hydroxides, and hydrated Cu²⁺ ions relative to the bulk solution. Total concentrations of Ni, P, and citrate were set to be 0.005, 0.05, and 0.01 mol kg⁻¹, respectively. Activity coefficients of related species were set to unity.



Figure S6. Fe isotopic variation (⁵⁶Fe/⁵⁴Fe) for hydrated Fe²⁺ and Fe³⁺ species at 295 K. DFT calculations of B3LYP/6-311G for various diffuse and polarization functions were tested with or without CPCM. (a) ln β for hydrated Fe²⁺. (b) ln β for hydrated Fe³⁺. (c) Isotope fractionation Δ^{56} Fe between hydrated Fe³⁺ and Fe²⁺. Literature data (see Table B2) of theoretical values (Ottonello and Zuccolini, 2009; Rustad et al., 2010) (dotted lines) and experimental values (Johnson et al., 2002; Welch et al., 2003) (bold lines) are shown together.