

Constitutive model for soft rocks considering structural healing and decay

Mamoru Kikumoto^{1*}, Nguyen Pham Quang Vu², Hideaki Yasuhara³, and Kiyoshi Kishida⁴

¹Department of Civil Engineering, Yokohama National University, Japan,
kikumoto@ynu.ac.jp

²Department of Civil Engineering, Yokohama National University, Japan, nguyen-vu-
hr@ynu.jp

³Department of Civil and Environmental Engineering, Ehime University, Japan,
yasuhara.hideaki.me@ehime-u.ac.jp

⁴Department of Urban Management, Kyoto University, Japan, kishida.kiyoshi.3r@kyoto-
u.ac.jp

*Corresponding author

Abstract

The behavior of soft rock depends on the contact area between mineral particles and the tensile strength of the interparticle cementation, which are usually referred to as structures. We investigated the effects of structural decay and healing on the behavior of soft rock through monotonic and slide-hold-slide triaxial tests under the drained condition with constant effective confining pressure. We developed a constitutive model for soft rocks incorporating structural healing and decay in the context of the extended critical state theory. The model was validated via laboratory tests and captured the behavior of soft rock, including the healing and decay phenomena.

Keywords: soft rock, rock structure, slide-hold-slide triaxial test, healing, decay, critical state theory

1. Introduction

In soft rock, mineral particles consist of aggregates of microcrystals formed by ionic, atomic, or molecular bonding, and these mineral particles are usually cemented or adhered mutually at the interparticle contact interfaces [1]. Thus, the stiffness and strength of the soft rock will depend primarily on the contact area between mineral particles and the tensile strength of interparticle cementation, which are usually referred to as rock structures.

The effects of the decay of the structures of soft rocks have been studied extensively. For example, the effects of the rock structures on the stress–strain characteristics have been investigated via laboratory experiments such as oedometer tests on Culebra shale [2], one-dimensional compression tests on chalk [3] and tuff [4, 5], and monotonic triaxial compression tests on calcarenite and tuff [6]. Leroueil and Vaughan [7] and Kavvas [8] discussed the effects of structures on the strength and stiffness of natural soils and weak rocks, and pointed out the similarities in behavior between natural soils and soft rocks.

Shao and Henry [9] have developed an elastoplastic model for porous rocks by extending a model for sands [10], and they have predicted the behavior of porous chalk. Gens and Nova [11], Kavvas et al. [12], Adachi and Oka [13], and Lagioia and Nova [14], among others, have also proposed constitutive models for various types of weak rocks such as mudstones, claystones, marls, shales, tuffs, weak limestones, and weak sandstones, and validated their simulation through comparison with a series of laboratory tests. The common features of the constitutive models for soft rocks are: (a) the models are formulated by extending the original models for unstructured geomaterials; (b) the structure of the weak rock is assumed to be destroyed due to the breakage of the interparticle cementation during loading.

Meanwhile, Dieterich and Kilgore [15] indicated that the contact area of the solid interface increases over a period of time, and that frictional resistance arises from the development of the contact area. It is reasonable to expect that a similar mechanism exists in the contacts between the mineral particles of soft rock at a microscopic level. Thus, we presume that the structure of soft rocks will recover to some extent after the loading process.

This leads to an increase in the stiffness and strength of soft rocks during the hold phase. Though a number of models [9, 11, 12, 13, and 14] have considered the effects of structural decay, the healing effect of the structure over time has not been considered.

Thus, it is crucial to consider both the effects of structural healing as well as structural decay in constitutive models, especially when estimating the long-term behavior of soft rocks. Therefore, the objective of the current study was to consider the effects of both structural healing and decay on the behavior of soft rocks. We first conducted triaxial tests on soft sedimentary rock with repeated slide-hold-slide (SHS) processes to observe the effects of structural healing and decay on the strength and stiffness of soft rock. After the slide-hold-slide processes, we investigated the effect of time on the structural recovery of soft sedimentary rock. We then developed a constitutive model that considered the effects of both structural healing and decay of soft rocks. In our model, the critical state theory was extended to consider the effect of the rock structure. Moreover, the subloading surface concept [16] was incorporated into the model to appropriately consider the combined effects of density and structure. The healing and decay of the structure was modeled using a newly introduced state variable and evolution law. The model was finally validated via monotonic and slide-hold-slide triaxial tests under drained condition.

2. Slide-Hold-Slide Triaxial Tests

We conducted consolidated, drained triaxial compression tests with multiple SHS processes on soft sedimentary rock to observe the effects of structural healing and decay on the stress-strain characteristics. The tests were performed under three types of effective confining pressures, where several holding time periods were applied. The experimental results were used to develop and validate a constitutive model for describing the fluctuation of the rock structure.

2.1 Test overview

The SHS triaxial tests were conducted on saturated cylindrical specimens of a pumice lapilli tuff, a type of soft volcanic sedimentary rock. The rock specimen was a porous vesicular rock primarily consisting of pyroclastic materials such as pumice. Photomicrographs of the specimen are shown in Figure 1. The modal composition of the specimen is measured by a method of point-counting and the sample consists of 85.0 % natural glassy mineral (39.0 % of Celadonite, 32.75 % of Zeolite and 12.75 % of Montmorillonite) and 15.0 % of crystal and lithic spall (7.0 % of Plagioclase, 5.75 % of Quartz and 2.25 % of Lithic). To measure the volumetric behavior of the specimens accurately, achieving the saturation of the specimens with water is essential. For this, water-soluble carbon dioxide (CO₂) gas is first percolated through the sample, followed by flushing with de-aired water. Backpressure is applied to improve the saturation during the tests.

The experiments were conducted using the apparatus shown in Figure 2 under consolidated, drained conditions. The temperature was kept constant at 20 °C throughout the test to ensure that thermal effect on the specimen were negligible. First, an isotropic consolidation path was applied until the predetermined effective confining stress σ_r' of 300 and 700 kPa was reached. The specimen was then sheared under drained condition with a constant effective confining pressure. The constant axial strain rate of 0.01 %/min was applied precisely by a screw jack until the post-peak phase, where the stress state approached the residual state. The holding process was then applied by maintaining a constant axial strain under various holding time periods from 60–241200 s. The holding process was always followed by the re-shearing process.

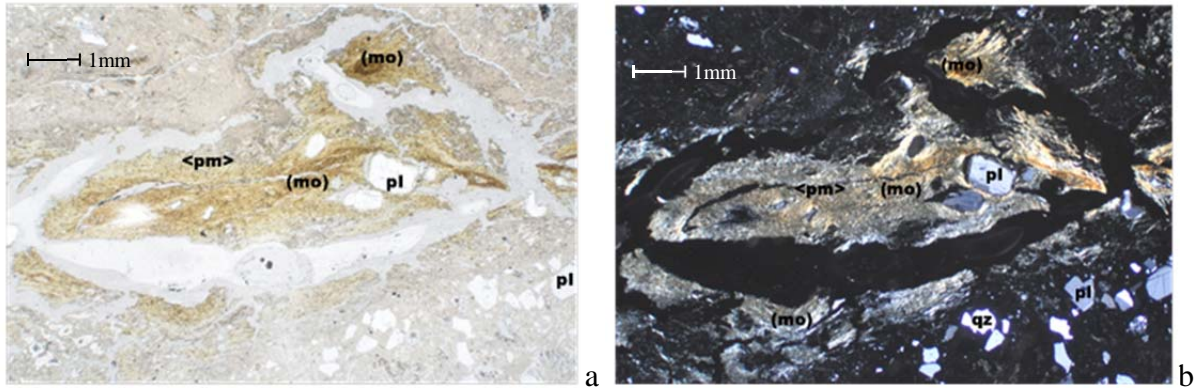


Figure 1: Photomicrographs of pumice lapilli tuff (magnification $\times 15$; qz: quartz, pl: plagioclase, pm: pumice, mo: Montmorillonite with Fe) (a) original image; (b) image focusing on the Fe montmorillonite.

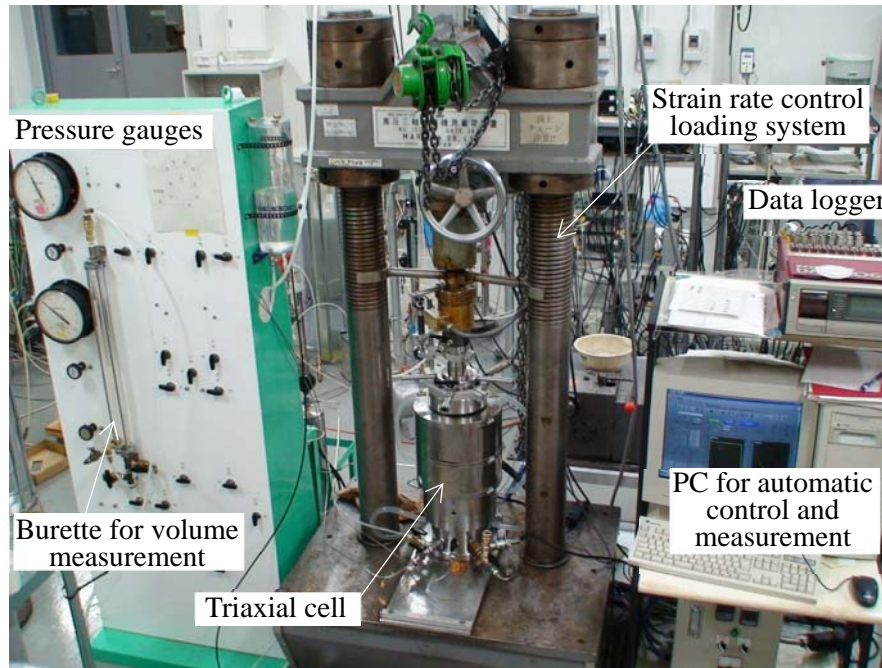


Figure 2: Triaxial testing apparatus.

2.2 SHS triaxial test results

Examples of the SHS triaxial test results are shown in Figures 3 and 4 with applied effective confining stresses σ_r of 300 kPa and 700 kPa, respectively. From the stress-strain relationship in the initial stage of shearing, a relatively high stiffness was exhibited, and an apparent peak stress was observed in the beginning stage of shearing. After the peak strength,

strain softening with positive dilatancy was observed and the stress finally reached the residual strength. The specimen exhibited the typical stress–strain behavior of soft sedimentary rocks, and these results appeared consistent with previous experimental observations (e.g., Adachi and Ogawa [17]). Enlarged views of the relationship between the axial strain ε_a and the deviator stress $q \left(= \sqrt{\frac{3}{2} \mathbf{s} : \mathbf{s}} = \sqrt{\frac{3}{2} \left\{ \boldsymbol{\sigma} : \boldsymbol{\sigma} - \frac{1}{3} (\text{tr} \boldsymbol{\sigma})^2 \right\}} \right)$ during the SHS process are shown in Figures 3(b) and 4(b). During the holding process in which the axial strain is held constant, stress relaxation with a reduction of the deviator stress could be observed. In the ensuing re-sliding process, the deviator stress increased with a high stiffness, reaching a peak value and then returning to the residual value. The magnitude of the strength recovery depended on the duration of the holding process, as the higher peak strength was particularly seen after a longer holding period. According to the experimental results at different confining pressures (300 and 700 kPa), strength recoveries could be observed after some of the longer holding periods. Such strength recoveries are considered to be a result of the healing of the rock structure. An example of the specimens after the SHS triaxial shearing is shown in Figure 5, which clearly shows a shear band formation. We expected the interparticle cementation to be destroyed due to the shear band formation during the first shearing process; then the rock structure recovered with the increase in the real area of the interparticle contact surfaces.

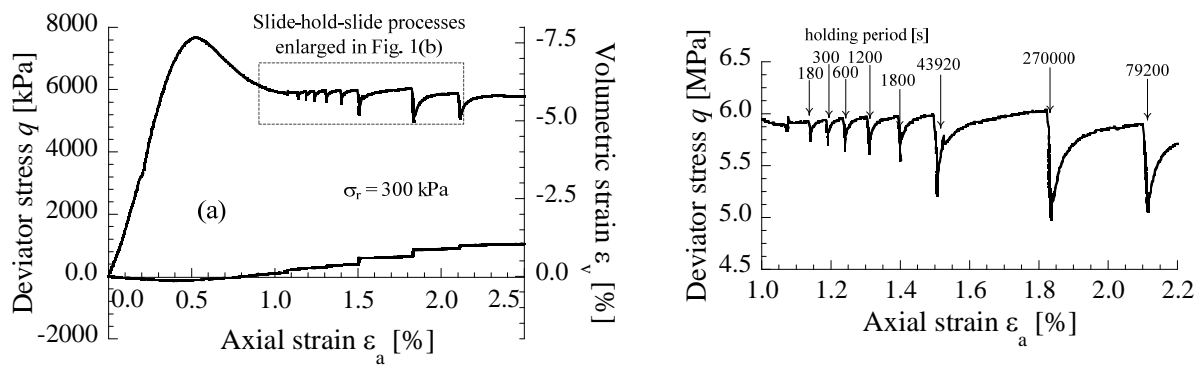


Figure 3: Stress–strain relationship in the SHS triaxial test ($\sigma_r = 300$ kPa): (a) Overview; (b)

Enlarged view.

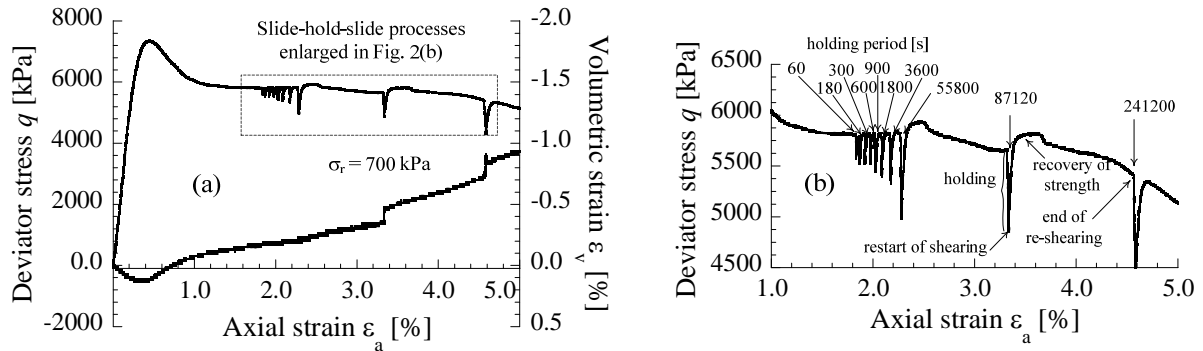


Figure 4: Stress–strain relationship in the SHS triaxial test ($\sigma_r = 700$ kPa): (a) Overall view; (b) Enlarged view.



Figure 5: Photo of specimen after the SHS triaxial CD tests (effective confining stress σ_r of 700 kPa).

3. Elastoplastic model for soft rocks considering structural healing and decay

Pellegrino [6] conducted triaxial compression tests on soft rocks such as calcarenite and tuff, and indicated that the soft rocks show a typical brittle behavior (approximately linear stress–strain relationship at small strains and brittle failure at large strains) at low stresses; while it shows typically ductile behavior at higher stresses (nonlinear elastoplastic behavior and

ductile failure). Novello [18] compared the behaviors of soils, soft rocks, and hard rocks in triaxial tests and argued that the brittle-ductile transition due to the increase in confining pressure in rocks was similar to the transition from overconsolidated to normally consolidated behavior in soils. Such a similarity in the behaviors of soft rocks and soils under different levels of confining pressure has also been confirmed by Hicher et al. [19]. In this regard, the critical state framework [10, 20] could be broadly applied to various geomaterials. Several extended versions of critical state models have been proposed for describing the behaviors of structured geomaterials such as structured soils [20] or soft sedimentary rocks [7]. However, the effect of healing of the rock structure has not been considered. Herein, we formulate a novel elastoplastic model based on the critical state theory [10, 20] for predicting the long-term behavior of soft rocks, including the effects of healing and decay phenomena on the rock structure.

First, we assumed an additive decomposition of the total strain rate tensor as

$$\dot{\boldsymbol{\varepsilon}} = \dot{\boldsymbol{\varepsilon}}^e + \dot{\boldsymbol{\varepsilon}}^p \quad (1)$$

where $\dot{\boldsymbol{\varepsilon}}^e$ and $\dot{\boldsymbol{\varepsilon}}^p$ are elastic and plastic strain rate tensors, respectively.

3.1 Elastic stress–strain relationship

For the elastic stress–strain relationship, we assumed a conventional, nonlinear elastic bulk modulus K given by

$$K = \frac{v_0}{\kappa} p' \quad (2)$$

where v_0 is the initial specific volume, κ is the swelling index that represents the slope of the elastic volumetric relationship in the semi-logarithmic $\ln p'$ – v plane, and p' is the mean effective stress given by $p' = \frac{\text{tr} \boldsymbol{\sigma}'}{3}$, where $\boldsymbol{\sigma}'$ is the Cauchy effective stress tensor. We assumed that Poisson's ratio ν_e was constant. Thus, the rate form of the elastic relationship was given by

$$\dot{\boldsymbol{\sigma}}' = \mathbf{D}^e : \dot{\boldsymbol{\varepsilon}}^e = \mathbf{D}^e : (\dot{\boldsymbol{\varepsilon}} - \dot{\boldsymbol{\varepsilon}}^p) \quad (3)$$

where \mathbf{D}^e is the elastic stiffness tensor:

$$\mathbf{D}^e = K \mathbf{1} \otimes \mathbf{1} + 2G \left(\mathbf{I} - \frac{1}{3} \mathbf{1} \otimes \mathbf{1} \right) \quad \#(4)$$

175 where G is the shear modulus:

$$G = \frac{3K(1 - 2\nu_e)}{2(1 + \nu_e)}. \quad \#(5)$$

176 3.2 Yield function for soft rocks considering structural healing and decay

177 The critical state is the ultimate condition approached by all states of rock when the rock is
 178 sheared. The critical state line (*CSL*) is chosen to be linear in a semi-logarithmic compression
 179 plane, which is the specific volume v ($= 1+e$) versus the logarithm of the mean effective
 180 stress $\ln p'$. Similar to the *CSL*, the limiting isotropic compression line (*LICL*) is a reference
 181 line over the *CSL* in the v - $\ln p'$ plane, which any state of rock approaches under isotropic
 182 compression. A state boundary surface, which defines the upper limit of the specific volume
 183 in stress-specific volume space above which no state of soft rock can exist, has been utilized
 184 in the formulation of the critical state model [18]. This surface contains *CSL* and *LICL* in the
 185 space of v , $\ln p'$, and $\zeta(\eta)$, which is a function of stress ratio η , (Figure 6). The specific
 186 volume v_{sbs} on the state boundary surface, which defines the least dense state of rock at stress
 187 (p, η) , is given by considering the combined effects of compression and dilation:

$$v_{\text{sbs}} = N - \lambda \ln \frac{p'}{p_a} + (\Gamma - N)\zeta(\eta) \quad \#(6)$$

188 where η ($= q/p'$) is the stress ratio, q is the deviator stress, p_a ($= 98$ kPa) denotes atmospheric
 189 pressure, λ is the compression index, and $\zeta(\eta)$ is a monotonic increasing function of stress
 190 ratio η satisfying $\zeta(0) = 0$ on *LICL* and $\zeta(M) = 1$ on *CSL*. Here, N and Γ represent specific
 191 volumes on *LICL* ($\eta = 0$) and *CSL* ($\eta = M$) at $p' = p_a$, respectively. Different functions of $\zeta(\eta)$
 192 have been used for different versions of critical state models. In the current model, $\zeta(\eta)$ is
 193 defined in accordance with the modified Cam clay [20]:

$$\zeta(\eta) = \frac{\ln \left\{ 1 + \left(\frac{\eta}{M} \right)^2 \right\}}{\ln 2} \quad \#(7)$$

194 where M is the critical state stress ratio ($= \eta_{\text{cs}}$).

$$S^d(\Psi) = b\Psi^2 \#(10)$$

where b is a constitutive parameter controlling the rate of decay of the structure. The second term of Eq. (9) describes the time-dependent healing of the structure, and $S^h(\Psi)$ is a function of Ψ representing the healing rate of the rock structure. Though this function may be dependent on the confining pressure and other factors, a simplified expression of the healing rate is given in this study as

$$S^h(\Psi) = \frac{1}{v_0} \frac{\Psi_{\max} - \Psi}{t_{\text{ref}}} \#(11)$$

as shown in Figure 8 (a), where Ψ_{\max} is a parameter defining the maximum value of Ψ , and t_{ref} is a parameter having a dimension of time, which describes the convergence rate of Ψ to Ψ_{\max} . We could explicitly describe the variation of the state parameter Ψ due to time-dependent healing by integrating Eq. (9) if we assumed that initially the rock had no structure ($t = 0, \Psi = 0$) and that no plastic deformation occurs ($\varepsilon^p = 0$)

$$\Psi = \Psi_{\max} \left(1 - \exp \frac{t}{t_{\text{ref}}} \right). \#(12)$$

From this equation, Ψ increases with time until it approaches its maximum value Ψ_{\max} , as shown in Figure 8 (b).

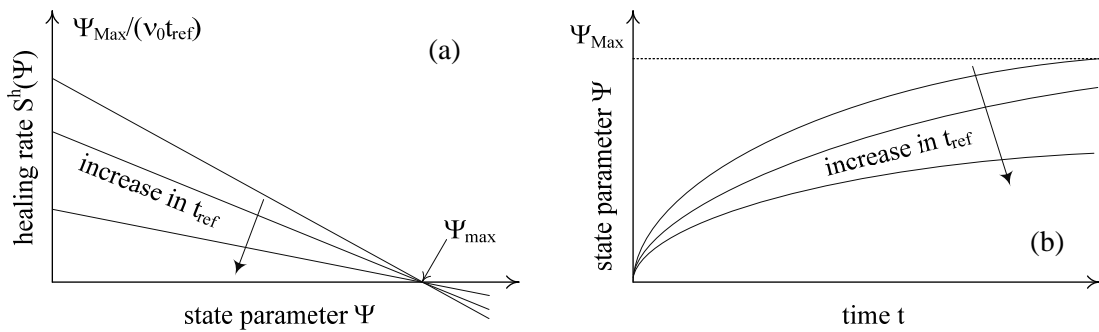


Figure 8: Modeling of the time-dependent healing of the rock structure via state parameter

Ψ : (a) rate of healing, (b) schematic figure showing an image of structural healing.

The subloading surface concept [16] was further introduced to describe the elastoplastic deformation of soft rocks whose states lie under the state boundary surface. As all states of rock are located on or below the state boundary surface in Figure 7, the state boundary surface defined the loosest, upper limit of the specific volume of rocks. We thus defined state parameter Ω as the specific volume difference between the current state and the least dense state under the same stress (p' , η) on the state boundary surface, as illustrated in Figure 9. Using state parameter Ω , the combination of the specific volume and mean effective stress was incorporated in this model to describe the changing strength and stiffness. According to this concept, the irreversible deformation below the state boundary surface and gradual approach to the state boundary surface with loading were properly modeled. Using state variable Ω , we could represent an arbitrary specific volume v as

$$v = v_{\text{sbs}}^{\text{structure}} - \Omega = N - \lambda \ln \frac{p'}{p_a} + (\Gamma - N)\zeta(\eta) + \Psi - \Omega. \#(13)$$

State parameter Ω always refers to the volumetric distance from the current state to the least dense state of soft rock (specific volume on the state boundary surface) under the current stress condition p' and q as well as the current state parameter Ψ for the rock structure.

During plastic flow, Ω decreased with the development of plastic deformation and converged to zero. The evolution of Ω could therefore be represented by

$$\frac{\dot{\Omega}}{v_0} = -Q(\Omega) \|\dot{\epsilon}^p\| \#(14)$$

where $\dot{\epsilon}^p$ is the plastic strain rate tensor and $Q(\Omega)$ is a function of Ω given by

$$Q(\Omega) = \omega \Omega^2 \#(15)$$

where ω is a parameter controlling the effect of density.

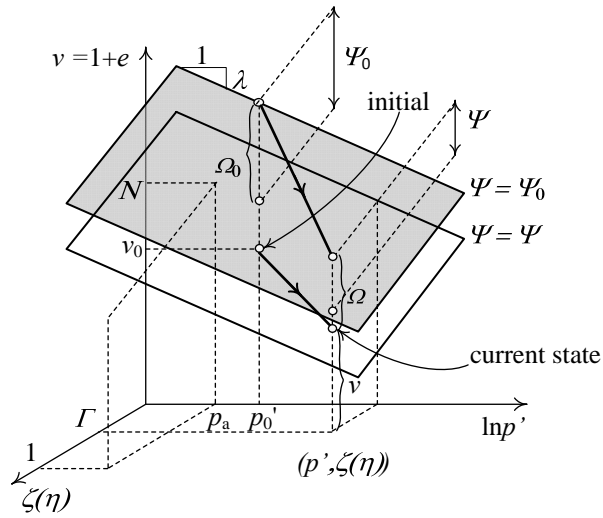


Figure 9: Modeling of the volumetric behavior of soft rocks considering the structural change.

From the current specific volume v (Eq. (13)), the initial specific volume v_0 can be obtained by substituting the initial states $v = v_0, \Psi = \Psi_0, \Omega = \Omega_0, p' = p'_0$, and $q = 0$ in Eq. (13) as

$$v_0 = N - \lambda \ln \frac{p_0}{p_a} + \Psi_0 - \Omega_0. \#(16)$$

The total volumetric strain (where compression is assumed to be positive) generated from the initial state to the current state is given by

$$\varepsilon_v = -\frac{dv}{v_0} = \frac{v_0 - v}{v_0}. \#(17)$$

By substituting Eqs. (13) and (16) in Eq. (17), we obtained

$$\varepsilon_v = \frac{1}{v_0} \left\{ \lambda \ln \frac{p'}{p_0} + (N - \Gamma) \zeta(\eta) - (\Psi - \Psi_0) + (\Omega - \Omega_0) \right\} \# (18)$$

by taking the trace on both sides of Eq. (3), the elastic volumetric strain could be obtained:

$$\varepsilon_v^e = \frac{\kappa}{v_0} \ln \frac{p'}{p_0} \quad \#(19)$$

The plastic volumetric strain could be determined by taking the difference between the total

279 volumetric strain given by Eq. (18) and the elastic volumetric strain given by Eq. (19):

$$\varepsilon_v^p = \frac{1}{v_0} \left\{ (\lambda - \kappa) \ln \frac{p'}{p_0} + (N - \Gamma) \zeta(\eta) - (\Psi - \Psi_0) + (\Omega - \Omega_0) \right\} \quad \#(20)$$

280 From Eq. (20), the yield function f for soft rocks, which considered the effect of the structure,
281 could be written as follows:

$$f = \frac{1}{v_0} \left\{ (\lambda - \kappa) \ln \frac{p'}{p_0} + (N - \Gamma) \zeta(\eta) - (\Psi - \Psi_0) + (\Omega - \Omega_0) \right\} - \varepsilon_v^p \quad \#(21)$$

282 Assuming associated flow in the proposed model, we obtained the plastic strain rate tensor:

$$\dot{\boldsymbol{\varepsilon}}^p = \langle \dot{\lambda} \rangle \frac{\partial f}{\partial \boldsymbol{\sigma}'} \quad \#(22)$$

283 where $\dot{\lambda}$ is the rate of the plastic multiplier. The loading criterion was thus given by $\dot{\lambda} > 0$.
284 Since an unlimited distortional strain was exhibited at the critical state without any change in
285 the stress or volume, $\text{tr} \left(\frac{\partial f}{\partial \boldsymbol{\sigma}'} \right)$ became zero when η equaled M . Thus, $(N - \Gamma)$ was equal to
286 $(\lambda - \kappa)/\ln 2$ in the case where Eq. (7) was applied, and the yield function could be given as
287 follows:

$$f = \frac{\lambda - \kappa}{v_0} \left[\ln \frac{p'}{p_0} + \ln \left\{ 1 + \left(\frac{\eta}{M} \right)^2 \right\} \right] - \frac{\Psi - \Psi_0}{v_0} + \frac{\Omega - \Omega_0}{v_0} - \varepsilon_v^p \quad \#(23)$$

288 3.3 Elastoplastic stress–strain relationship

289 In the purely elastic regime, the rate of the plastic multiplier $\langle \dot{\lambda} \rangle$ remains zero. Meanwhile,
290 during elastoplastic deformation, the stress remains on the yield surface, and the yield
291 function f remains equal to zero. The time derivative of the yield function \dot{f} consequently
292 vanishes whenever the rate of the plastic multiplier $\langle \dot{\lambda} \rangle$ is positive. Therefore, we could write
293 a consistency condition that has validity for either elastic or elastoplastic deformation as

$$0 = \langle \dot{\lambda} \rangle \dot{f}. \quad \#(24)$$

294 During plastic flow, we applied the consistency condition to the time derivative of the yield
295 function $\dot{f}(\boldsymbol{\sigma}', \varepsilon_v^p, \Psi, \Omega)$ calculated from Eq. (23) as follows

$$\dot{f} = \frac{\partial f}{\partial \boldsymbol{\sigma}'} : \dot{\boldsymbol{\sigma}}' + \frac{\partial f}{\partial \Psi} \dot{\Psi} + \frac{\partial f}{\partial \Omega} \dot{\Omega} + \frac{\partial f}{\partial \varepsilon_v^p} \dot{\varepsilon}_v^p \quad (25)$$

$$\dot{f} = \frac{\lambda - \kappa}{v_0} \left[\frac{1}{p'} \frac{\partial p'}{\partial \boldsymbol{\sigma}'} + \frac{2\eta}{M^2 + \eta^2} \frac{\partial \eta}{\partial \boldsymbol{\sigma}'} \right] : \dot{\boldsymbol{\sigma}}' - \frac{\dot{\Psi}}{v_0} + \frac{\dot{\Omega}}{v_0} - \dot{\varepsilon}_v^p = 0 \quad (26)$$

Substituting Eqs. (3), (9), (14), (15), and (22) in Eq. (26), we obtained the plastic multiplier:

$$\langle \dot{\Lambda} \rangle = \left\langle \frac{\frac{\partial f}{\partial \boldsymbol{\sigma}'} : \mathbf{D}^e : \dot{\boldsymbol{\varepsilon}} - S^h(\Psi)}{\text{tr} \left(\frac{\partial f}{\partial \boldsymbol{\sigma}'} \right) + \frac{\partial f}{\partial \boldsymbol{\sigma}'} : \mathbf{D}^e : \frac{\partial f}{\partial \boldsymbol{\sigma}'} + \{Q(\Omega) - S^d(\Psi)\} \left\| \frac{\partial f}{\partial \boldsymbol{\sigma}'} \right\|}} \right\rangle. \quad (27)$$

We consequently obtained the rate form of the elastoplastic stress–strain relationship from Eqs. (3), (21), and (26):

$$\dot{\boldsymbol{\sigma}}' = \mathbf{D}^e : \dot{\boldsymbol{\varepsilon}} - \left\langle \frac{\frac{\partial f}{\partial \boldsymbol{\sigma}'} : \mathbf{D}^e : \dot{\boldsymbol{\varepsilon}} - S^h(\Psi)}{\text{tr} \left(\frac{\partial f}{\partial \boldsymbol{\sigma}'} \right) + \frac{\partial f}{\partial \boldsymbol{\sigma}'} : \mathbf{D}^e : \frac{\partial f}{\partial \boldsymbol{\sigma}'} + \{Q(\Omega) - S^d(\Psi)\} \left\| \frac{\partial f}{\partial \boldsymbol{\sigma}'} \right\|}} \right\rangle \mathbf{D}^e : \frac{\partial f}{\partial \boldsymbol{\sigma}'} \quad (28)$$

When the rate of the plastic multiplier $\dot{\Lambda}$ is positive, the rate form of the elastoplastic stress–strain relationship can be expressed as

$$\dot{\boldsymbol{\sigma}}' = \mathbf{D}^{ep} : \dot{\boldsymbol{\varepsilon}} - \mathbf{D}^t \quad (29)$$

where \mathbf{D}^{ep} and \mathbf{D}^t are defined as follows:

$$\mathbf{D}^{ep} = \mathbf{D}^e - \frac{\mathbf{D}^e : \frac{\partial f}{\partial \boldsymbol{\sigma}'} \otimes \frac{\partial f}{\partial \boldsymbol{\sigma}'} : \mathbf{D}^e}{\text{tr} \left(\frac{\partial f}{\partial \boldsymbol{\sigma}'} \right) + \frac{\partial f}{\partial \boldsymbol{\sigma}'} : \mathbf{D}^e : \frac{\partial f}{\partial \boldsymbol{\sigma}'} + \{Q(\Omega) - S^d(\Psi)\} \left\| \frac{\partial f}{\partial \boldsymbol{\sigma}'} \right\|}} \quad (30)$$

$$\mathbf{D}^t = \frac{\mathbf{D}^e : \frac{\partial f}{\partial \boldsymbol{\sigma}'} S^h(\Psi)}{\text{tr} \left(\frac{\partial f}{\partial \boldsymbol{\sigma}'} \right) + \frac{\partial f}{\partial \boldsymbol{\sigma}'} : \mathbf{D}^e : \frac{\partial f}{\partial \boldsymbol{\sigma}'} + \{Q(\Omega) - S^d(\Psi)\} \left\| \frac{\partial f}{\partial \boldsymbol{\sigma}'} \right\|}} \quad (31)$$

3.4 Determination of the parameters involved in the proposed model.

The first set of constitutive parameters (λ , κ , N , M , ν , a) can be readily obtained from the results of elementary tests. The results of isotropic consolidation tests on soft rock plotted in the $(e - \ln p')$ plane can be used to determine the slope λ of the LICL, κ from the slope of the

swelling line unloading part, and N from the specific volume on the LICL under atmospheric pressure at the destructured state. The parameter “ a ” can be obtained by fitting the compression curve under the reloading path. The slope M of the CSL in the q - p' plane, and Poisson’s ratio ν , are then calibrated from the result of the monotonic triaxial CD or $\overline{\text{CU}}$ tests.

The second set of constitutive parameters (b , t_{ref} , Ψ_0 , Ψ_{max}) that control the rate of the decay and healing of the rock structure can be deduced by fitting the simulations to SHS triaxial tests. Ψ_0 and Ψ_{max} are the initial “structure” and the maximum “structure” assumed in the material, respectively. Their values can be obtained by fitting the stress–strain relationship for the SHS tests, so that the material can gain its maximum structure. $1/t_{ref}$ describes the rate of the increase in Ψ to Ψ_{max} when the rock gains its structure during the hold phase. “ b ” describes the rate of decrease in Ψ under plastic deformation.

4. Simulation results

4.1 Monotonic triaxial shearing and decay of the structure

The proposed model was compared with the monotonic triaxial shearing test under drained conditions with a constant effective confining pressure (Adachi and Oka [13]) to validate the modeling of structural decay due to deformation. The set of material parameters shown in Table 1 was used for all simulations with different confining pressures. The initial conditions are summarized in Table 2. The axial strain rate applied during the drained triaxial shearing was 3.33 %/h. As shown in Figure 10, the proposed model can capture the tendency of strain hardening and strain softening as well as the tendency of dilatancy in soft rock.

Table 1: Constitutive parameters for soft rock (pumice lapilli tuff) for the monotonic triaxial

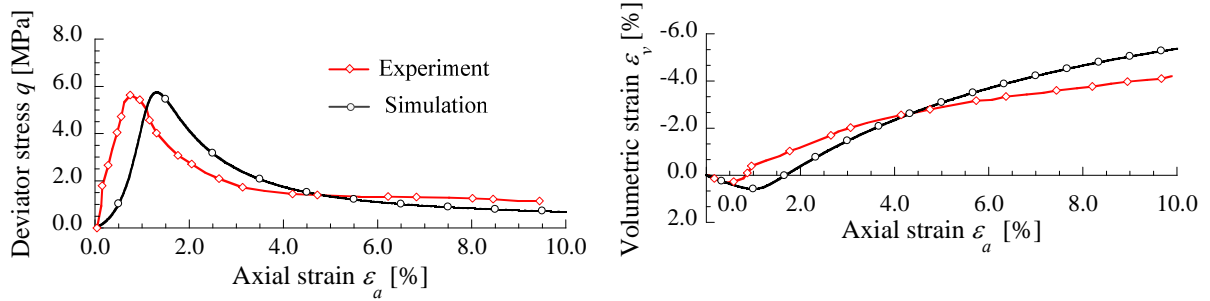
CD test.

λ	Compression index	0.053
κ	Swelling index	0.004

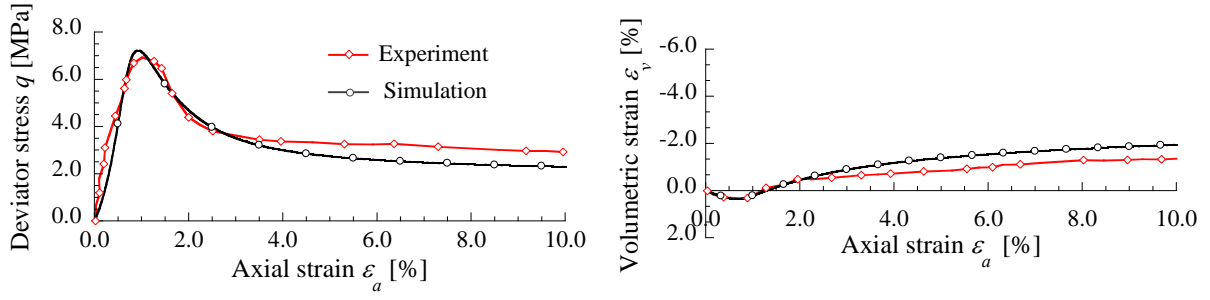
M	Critical state stress ratio	1.7
N	Void ratio on LICL at atmospheric pressure p' (= 98 kPa)	0.83
ν	Poisson's ratio	0.12
ω	Rate of evolution of Ω	1×10^4
b	Rate of decay of rock structure Ψ due to plastic strain	100
$1/t_{\text{ref}}$	Rate of healing of rock structure Ψ [1/h]	0.10
Ψ_{max}	Upper limit of rock structure Ψ	0.90

Table 2: Initial conditions for the monotonic triaxial CD test with different confining pressures.

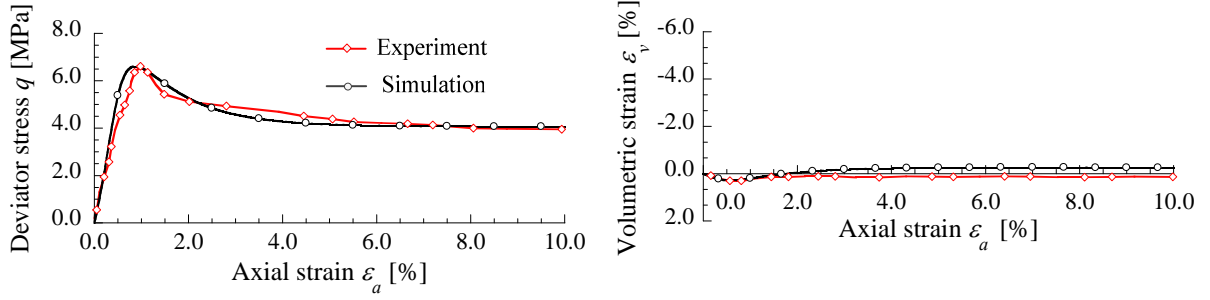
Effective confining stress σ'_0 (kPa)	Initial void ratio e_0	Initial state parameter Ψ_0
98	0.72	0.105
490	0.72	0.130
980	0.72	0.125
1960	0.72	0.130



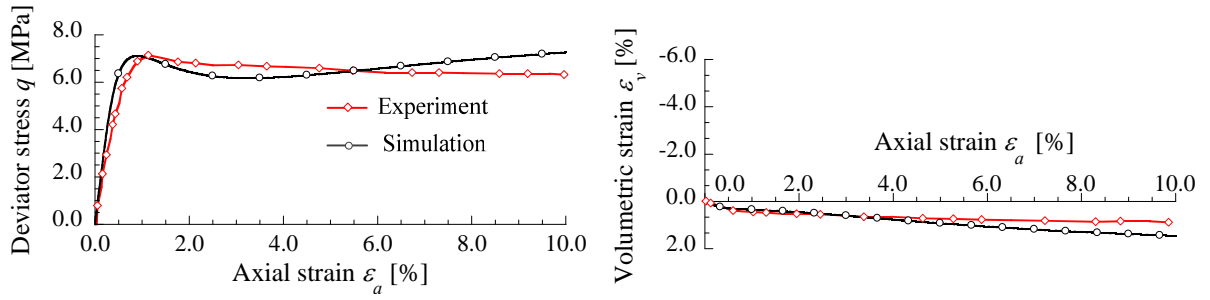
(a) Effective confining pressure $\sigma_r = 98$ kPa



(b) Effective confining pressure $\sigma_r = 490$ kPa



(c) Effective confining pressure $\sigma_r = 980$ kPa



(d) Effective confining pressure $\sigma_r = 1960$ kPa

Figure 10: Comparison between experimental and numerical simulation results on consolidated, drained monotonic triaxial compression test for various confining pressures.

4.2 Slide-hold-slide triaxial shearing and structural healing and decay

The characteristics of the proposed model are explained using examples of the simulation results of the SHS triaxial test shown in Figure 11 and Figure 12. SHS shearing was simulated under the drained condition with a constant effective confining pressure of 700 kPa. Four holding processes were applied with a stepwise increase in the holding time periods.

In the model, the state parameter Ψ increased due to the effect of the holding time (Figure 11 (c) and Figure 12 (c)) in the evolution law in Eq. (9). The increase in Ψ led to the upward movement of the state boundary surface in the plane of mean effective stress and specific volume. Consequently, the state parameter Ω , which is the specific volume difference between the current state and the least dense state under the same state on the state boundary surface, also increased (Figure 11 (d) and Figure 12 (d)). As the holding time increased, Ψ increased. After a certain long holding time, Ψ gradually reached its maximum value Ψ_{\max} , which was assumed as the fully structured state of the rock. Decay in the rock strength was observed in the re-sliding process. First, Ψ decreased (Figure 11 (c)) because of the larger decay effect of plastic strain over the healing effect of holding time in the evolution of the state variable in Eq. (9). The decrease in Ψ moved the state boundary surface downward in the direction of specific volume, leading to a decrease in the state variable Ω (Figure 11 (d)). This resulted in a decrease in the stiffness. In addition, the deviator stress increased with a rather high stiffness, reaching a peak value before returning to the residual value. The magnitude of the strength recovery depended on the duration of the holding process, as higher peak strengths were typically observed after a longer holding time.

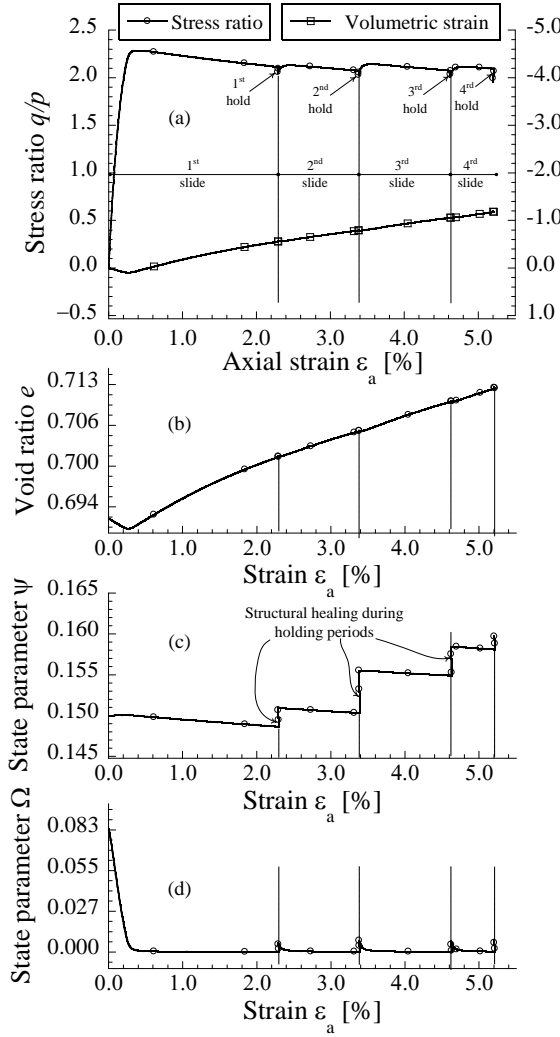


Figure 11: Simulation results of the
SHS triaxial CD test
(variations in relation to axial strain)

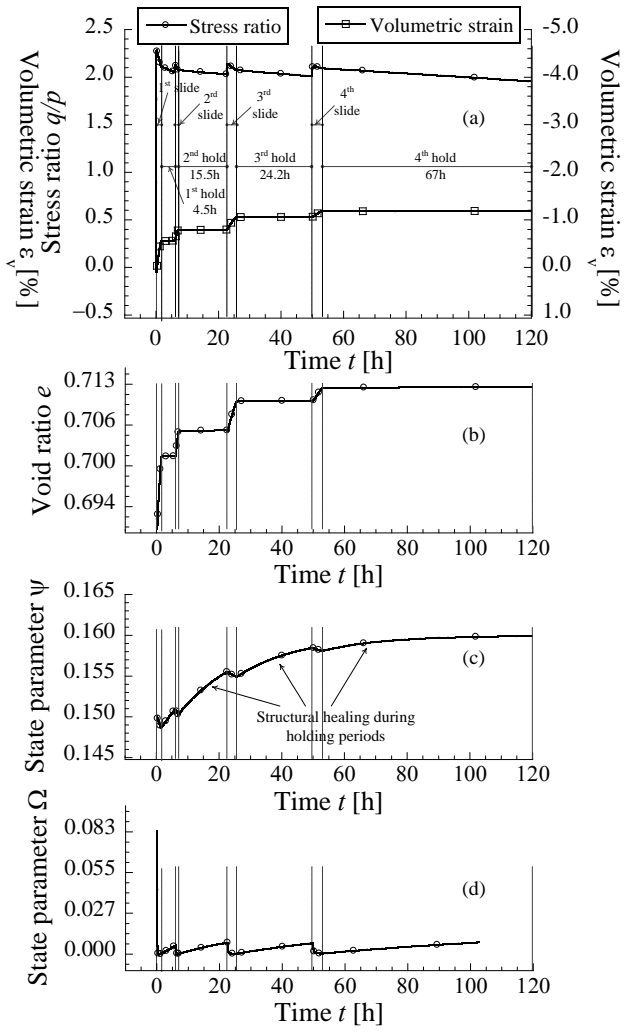


Figure 12: Simulation results of the
SHS triaxial CD test
(variations in relation to time)

Next, consolidated drained triaxial tests on a pumice lapilli tuff with multiple SHS processes under constant effective confining pressures of 300 and 700 kPa were simulated using the proposed model. The constitutive parameters used for the simulation are listed in Table 3, while the initial conditions are given in Table 4.

From Figures 13, 14, 15, and 16, the proposed model accurately predicted the stress–time–strain relationship of soft rock with SHS processes under two different effective confining pressures. The healing and decay phenomena of the soft rock structure and their effect on the strength and dilatancy characteristics were properly captured by the proposed

model. The characteristic of the proposed model was confirmed: longer holding times result in greater recovery of the rock strength.

Table 3: Constitutive parameters for soft rock (pumice lapilli tuff) for the SHS triaxial CD

test

λ	Compression index	0.0360
κ	Swelling index	0.0012
M	Critical state stress ratio	1.9
N	Void ratio on LICL at atmospheric pressure p' (= 98 kPa)	0.70
ν	Poisson's ratio	0.30
ω	Rate of evolution of Ω	3×10^5
b	Rate of decay of Ψ due to plastic strain	2.0
$1/t_{\text{ref}}$	Rate of healing of rock structure Ψ [1/h]	0.05
Ψ_{max}	Upper limit of rock structure Ψ	0.16

Table 4: Initial conditions for the SHS triaxial CD test

Effective confining stress	Initial void	Initial state
σ'_0 (kPa)	ratio e_0	parameter Ψ_0
300	0.692	0.15
700	0.692	0.15

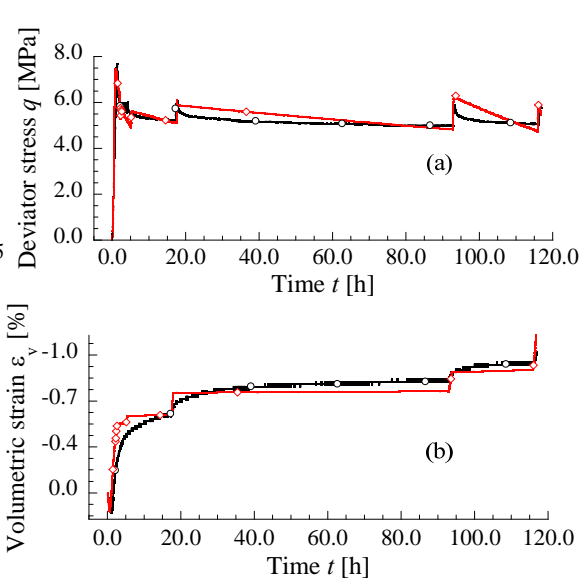
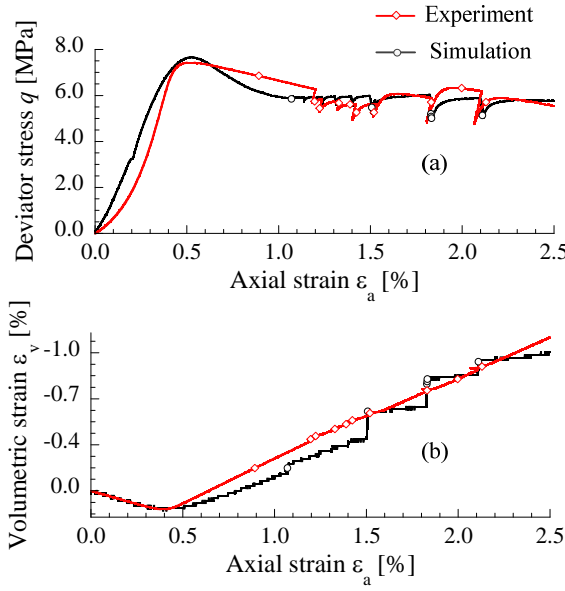


Figure 13: Comparison of the experimental and numerical simulation results of the SHS triaxial CD test
 $\sigma_r = 300$ kPa
 (variation in relation with axial strain)

Figure 14: Comparison of the experimental and numerical simulation results of the SHS triaxial CD test
 $\sigma_r = 300$ kPa
 (variation in relation with time)

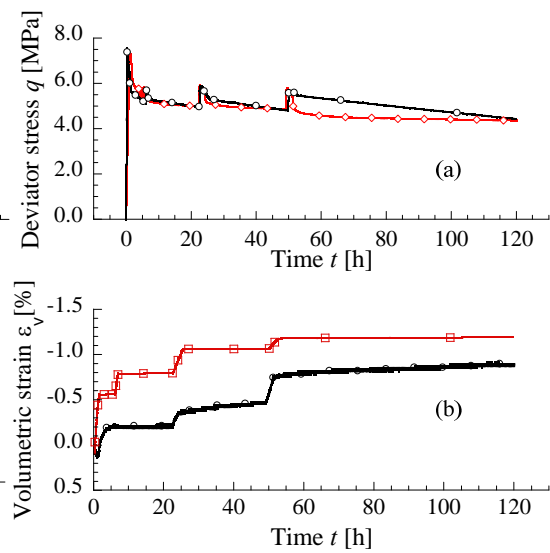
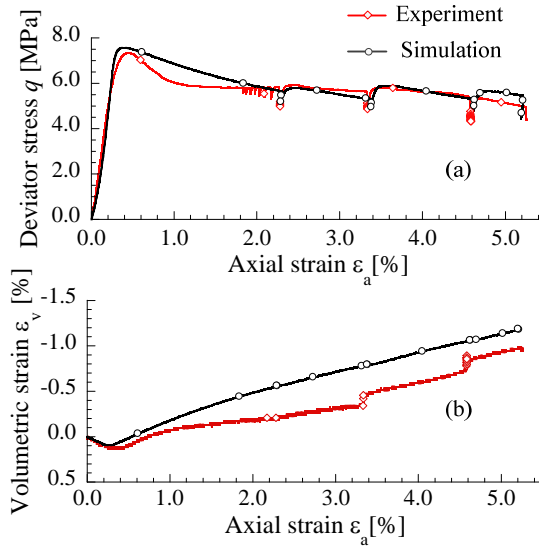


Figure 15: Comparison of the experimental and numerical simulation results of the SHS triaxial CD test
 $\sigma_r = 700$ kPa
 (variation in relation with axial strain)

Figure 16: Comparison of the experimental and numerical simulation results of the SHS triaxial CD test
 $\sigma_r = 700$ kPa
 (variation in relation with time)

4.4 Parametric study via monotonic shear in consolidated, drained triaxial compression test

A parametric study of the proposed model was presented using numerical simulations of a monotonic triaxial compression test in the consolidated, drained condition. The same material parameters and initial conditions as those listed in Table 3 and Table 4 were applied in this simulation. Analysis was performed on the effect of material parameter b accounting for the decaying rate with different values of b ranging from 2 to 16 and the effects of a wide range of strain rates (from $2.5 \times 10^{-3} \%$ /h to $8.0 \times 10^{-2} \%$ /h).

Figure 17 shows that different decaying rates of rock friction for different rock types can be controlled by the material parameter b in the proposed model. As the value of b or the strain rate increases, both decaying of rock friction and a lower negative dilatancy during the softening process were observed. Similarly, different healing rates for various rock types can also be controlled through the material parameters t_{ref} . Meanwhile, as shown in Figure 18, the proposed model can consider the effect of different shearing rates on the behavior of the same rock. These parameters should be determined by analyzing the experimental results of SHS tests for each type of rock.

Regarding the strength, as observed in Figure 17 (a), the rock reached its critical state stress ratio, in which $(q/p)_{\text{cs}} = M$. The rate of structural decay could be studied by observing the variation of the state variable Ψ in Figure 17 (c) and 18 (c) when either b or the strain rate was changed. In Figure 18 (c), Ψ reaches a limiting value, in which $d\Psi$ is zero. In this limiting state, the healing and decay given by Eq. (9) likely had the same amount of effect on the rock friction. The state variable Ω , which is the distance from the current void ratio to the ratio on the state boundary surface at the same mean stress, started from an initial value and gradually converged to zero.

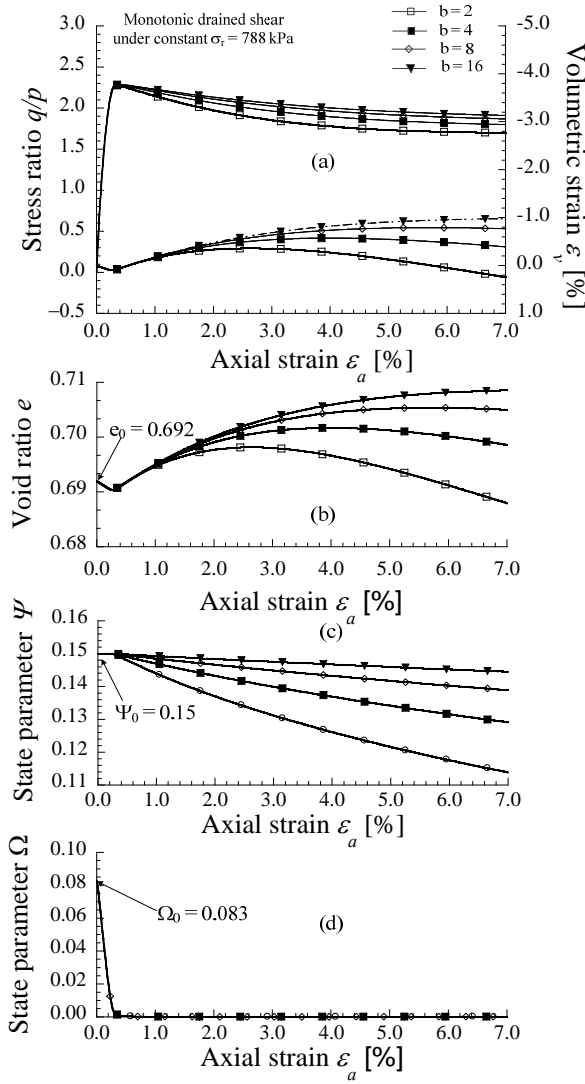


Figure 1: Simulation results of the effect of parameter b on the decay rate in monotonic triaxial shear test under the consolidated drained condition.

5. Conclusion

In summary, this study has highlighted the effects of structural healing and decay on the stress–strain characteristics of soft rock and presented a potential approach for describing the healing and decay of a rock structure in the formulation of an elastoplastic constitutive model. The model performance was validated by comparing simulations with the experimental results of drained, monotonic triaxial shearing tests and drained SHS triaxial shearing tests on soft sedimentary rock. Unlike the existing models for structured soil and

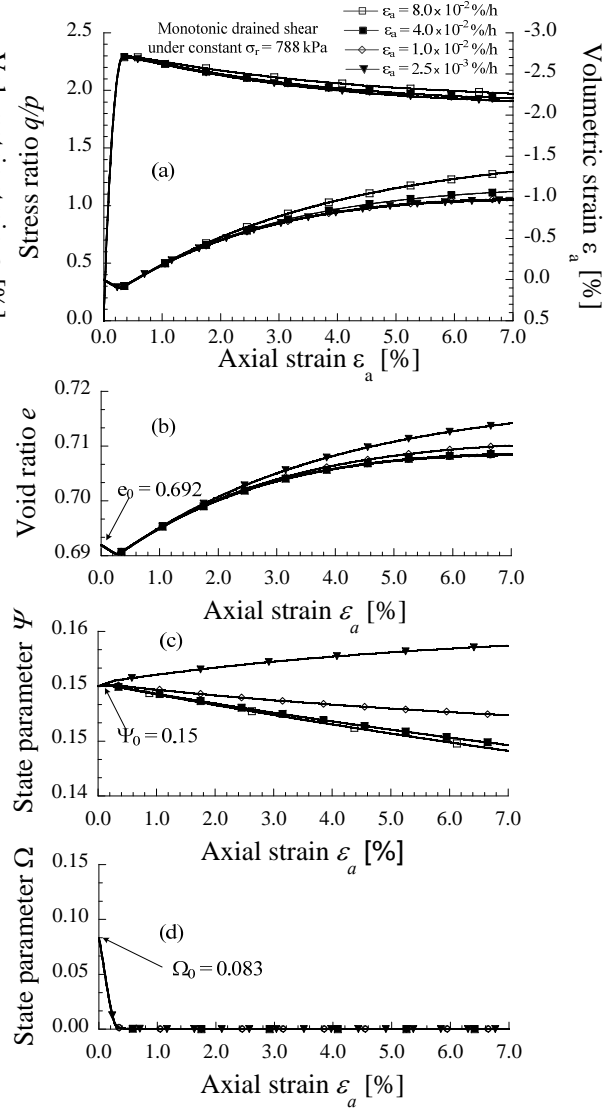


Figure 2: Simulation results of the effect of the strain rate on the decay rate in monotonic triaxial shear test under the consolidated drained condition.

soft rock, our model is capable of describing the time-healing effect of the structure as well as the decay of the structure due to deformation. Our constitutive model is formulated based on a general stress–strain tensor. Thus, it is easy to implement the model in a finite element method to analyse any geotechnical problems considering the long-term behavior of soft sedimentary rocks.

Acknowledgments

Funding: This work was supported by JSPS KAKENHI [grant numbers 26249139, 15K14030]. The second author was financially supported by the Monbukagakusho (MEXT) scholarship. We thank Mr. Takenori Araki and Mr. Masato Ohno, formerly graduate students at Kyoto University, for their support in conducting the experiments.

Appendix

The following notations and symbols are used in the present paper: bold letters denote vectors and matrices; the symbol “ \cdot ” denotes an inner product of two vectors (e.g., $\mathbf{a} \cdot \mathbf{b} = a_i b_i$) or a single contraction of the adjacent indices of two tensors (e.g., $(\mathbf{c} \cdot \mathbf{d})_{ij} = c_{ik} d_{kj}$); the symbol “ $:$ ” denotes an inner product of two second-order tensors (e.g., $\mathbf{c} : \mathbf{d} = c_{ij} d_{ij}$) or a double contraction of the adjacent indices of tensors of rank two and higher (e.g., $(\mathbf{e} : \mathbf{c})_{ij} = e_{ijkl} c_{kl}$); “ \otimes ” denotes a tensor product of two vectors (e.g., $(\mathbf{a} \otimes \mathbf{b})_{ij} = a_i b_j$) or a tensor product of two second-order tensors (e.g., $(\mathbf{c} \otimes \mathbf{d})_{ijkl} = c_{ij} d_{kl}$); “ $\| \quad \|$ ” denotes the norm of a vector (e.g., $\|\mathbf{a}\| = \sqrt{\mathbf{a} \cdot \mathbf{a}} = \sqrt{a_i a_i}$) or a second-order tensor (e.g., $\|\mathbf{c}\| = \sqrt{\mathbf{c} : \mathbf{c}} = \sqrt{c_{ij} c_{ij}}$); $\mathbf{1}$ is the second-order identity tensor; \mathbf{I} is the fourth-order identity tensor $\left(I_{ijkl} = \frac{1}{2}(\delta_{ik} \delta_{jl} + \delta_{il} \delta_{jk}) \right)$; “ $\langle \quad \rangle$ ” is that Macaulay bracket that denotes the ramp function as $\langle x \rangle = \begin{cases} x & \text{if } x > 0 \\ 0 & \text{if } x \leq 0 \end{cases}$; over-dot “ $\dot{\quad}$ ” denotes the time derivative; and a zero subscript denotes an initial state.

467 **Reference**

- 468 [1] Shah DL, Shroff, AV. Rock structure and physical properties. In *Soil mechanics and*
 469 *geotechnical engineering*, CRC Press; 2003; Chapter 2: 6-21.
- 470 [2] Banks DC, Strohm WE, DeAugulo M, Lutton RJ. Study of clay-shale slopes along the
 471 Panama Canal. Report No. 3, Engineering analyses of slides and strength properties of
 472 clay shales along the Gaillard Cut. Technical report S-70-9. US Army Engineers
 473 Waterways Experiment station, Vicksburg, Miss; 1975.
- 474 [3] Addis MA, Jones ME. Mechanical behavior and strain-rate dependence of high
 475 porosity chalk. In *Proc Int Chalk Sympo*, Brighton, London: Thomas Telford; 1990;
 476 111-116.
- 477 [4] Aversa S, Evangelista A. The mechanical behaviour of a pyroclastic rock: yield strength
 478 and 'destruction' effects. *Rock Mech Rock Eng* 1998; 31(1): 25-42.
- 479 [5] Aversa S, Evangelista A, Leroueil S, Picarelli L. Some aspects of the mechanical
 480 behaviour of 'structured' soils and soft rocks. In *Proc Int Sym Geotechnics of Hard Soils*
 481 *- Soft Rocks*, Athens: Balkema; 1993; 359-366.
- 482 [6] Pellegrino A. Mechanical behavior of soft rocks under high stresses. In *Proc 2nd Int*
 483 *Congr Rock Mech*, Belgrade; 1970; 2: 173-180.
- 484 [7] Leroueil S, Vaughan PR. The general and congruent effects of structure in natural soils
 485 and weak rocks. *Géotechnique* 1990; 40(3): 467-88.
- 486 [8] Kavvadas M. General report: Modelling the soil behaviour – Selection of soil
 487 parameters. In *Proc Int Symp Geotechnics of Hard Soils-Soft Rocks*, eds. Evangelista
 488 & Picarelli, Naples: Balkema; 1998; 3: 1441-1481.
- 489 [9] Shao JF, Henry JP. Development of an elastoplastic model for porous rocks. *Int J Plast*
 490 1991; 7(1-2): 1-13.
- 491 [10] Lade V. Elasto-plastic stress-strain theory for cohesionless soil with curved yield
 492 surfaces. *Int J Solids Struct* 1977; 13: 1019-35.

- [11] Gens A, Nova R. Conceptual bases for a constitutive model for bonded soils and weak rocks. In *Proc Int Sym Geotechnics of Hard Soils - Soft Rocks*, Athens: Balkema; 1993; 485–494.
- [12] Kavvas M, Anagnostopoulos A, Kaltezionis N. A framework for the mechanical behavior of the cementer Crinith marl. In *Proc Int Sym Geotechnics of Hard Soils - Soft Rocks*, Athens: Balkema; 1993; 577–583.
- [13] Adachi T, Oka F. An elasto-plastic constitutive model for soft rock with strain softening. *Int J Numer Anal Meth Geomech* 1995; 19:223–7.
- [14] Lagioia R, Nova R. An experimental and theoretical study on the behavior of a calcarenite in triaxial compression. *Géotechnique* 1995; 45(4): 613–648.
- [15] Dieterich JH, Kilgore BD. Direct observation of frictional contacts: new insights for state-dependent properties. *PAGEOPH* 1994; 143: No. 1/2/3: 283–302.
- [16] Hashiguchi K, Ueno M. Elastoplastic constitutive laws of granular materials, Constitutive Equations of Soils. In *Proc Int Soc Soil Mech Geotech Eng Spec Ses 9*. eds Murayama S, Schofield A N. JSSMFE. Tokyo; 1977; 73–82.
- [17] Adachi T, Ogawa T. Mechanical properties and failure criterion of soft sedimentary rock. *J. JSCE* 1980; 295: 51–63 (in Japanese).
- [18] Novello E A, Johnston L W. Geotechnique materials and the critical state. *Géotechnique* 1995; 45(2): 223–35.
- [19] Hicher P Y and Shao J F. Constitutive modeling of soils and rocks. John Wiley & Sons, 2013
- [20] Roscoe K H, Schofield A N, and Thurairajah A. Yielding of clays in state wetter than critical. *Géotechnique* 1963; 13(3): 211–40.