

Empirical Equations for Estimation of Transmission Losses Based on Field Measurements in Ephemeral Streams

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Transmission losses is caused by infiltration into the streambed in ephemeral streams. The conventional methods for flood routing in wadis is impossible to achieve due to transmission losses. The Muskingum routing procedure in its basic form has two parameters, the channel time lag, K_m , and the weighting parameter, x . However, both parameters do not consider transmission losses of floods in channels. O'Donnell 1985 introduced a third parameter, α , in the continuity equation to allow for the lateral movement of floodwater. Elfeki et. Al (2014) carry some modifications of the assumptions such that the negative sign of the parameter α represents transmission losses. In this research, the third-parameter α has been investigated as a tool for estimation of transmission losses using data from Yiba catchment in the Kingdom of Saudi Arabia. A spreadsheet model will be developed to deduce the equations.

EMPIRICAL EQUATIONS FOR ESTIMATION OF TRANSMISSION LOSSES BASED ON FIELD MEASUREMENTS IN EPHEMERAL STREAMS

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Outline

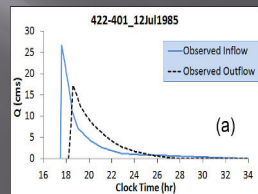
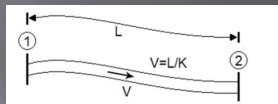
1. Problem Statement.
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3. Study Area.
4. Field Observations and Measurements.
5. Muskingum Method incorporating Transmission losses
6. Model results and Parameter Estimation
7. Results
8. Conclusions

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Problem Statement

- o In the literature, researchers usually use three parameter Muskingum method for river flood routing with lateral inflow (O'Donnell, T. (1985).
- o Recently the method has been applied on ephemeral channels to incorporate transmission losses (α) with success by Elfeki, et al. (2014).



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Research Objective

1. To Investigate The parameter (α) as a tool for estimation of transmission losses using data from Yiba catchment in the Kingdom of Saudi Arabia.
2. To develop A spreadsheet model to deduce the empirical equations for estimation (α) in ephemeral channels

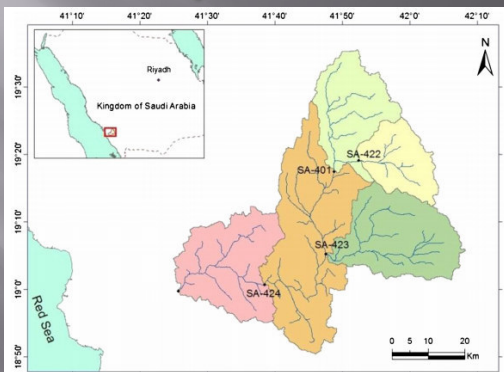
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Study Area

Yiba basin is located in Makkah Al-Mukaramah region, in the southwestern part of Saudi Arabia.

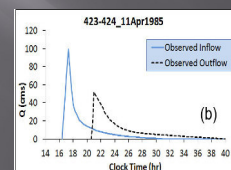
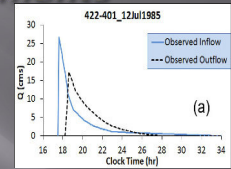
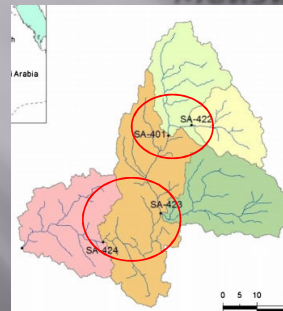
It has an area of 2346 km² and drains its water towards the Red Sea.



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Field Observations and Measurements

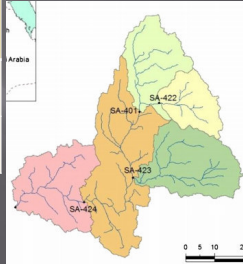


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Catchments and Channel Characteristics

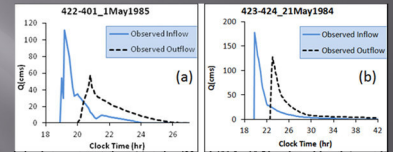
Station	Area (km ²)	Length of the main stream (km)	Basin Slope (m/m)	Channel reach length (km)
422	306.01	23.3	0.3322	10.5
401	764.96	35.07	0.2833	
423	612.83	36.2	0.2657	
424	2346.27	51.8	0.1969	



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Muskingum Method incorporating Transmission losses



Observed hydrograph transmission between two stations: (a) station 422 and 401 for 10.5 km channel length (event 1st of May 1985), (b) station 423 and 424 for 16.1 km channel length (event 21st of May 1984).

O'Donnell, (1985) expressed the three parameter model as follows:

$$I(1 + \alpha) = O - \frac{dS}{dt} \dots\dots\dots(1)$$

where:

S is the channel storage,

I is the inflow hydrograph into the channel segment,

O is the outflow hydrograph from the channel segment, and

α is the coefficient of lateral flow.

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The empirical formula for the channel storage is given by

$$S = K_m [(1 + \alpha)xI + (1 - x)O] \dots\dots\dots(2)$$

where,

K_m is the channel time lag (the hydrograph movement time), and

x is a weighting parameter

$$O_{t+\Delta t} = d_1 I_t + d_2 I_{t+\Delta t} + d_3 O_t \dots\dots\dots(3)$$

where, d₁, d₂, and d₃ are the coefficients that can be obtained by the matrix inversion least square solution

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The three parameters, K_m, x and α, are related to d₁, d₂, and d₃ by the following formulas (O'Donnell, 1985),

$$K_m = \Delta t \frac{d_1 + d_2 d_3}{(1 - d_3)(d_1 + d_2)} \dots\dots\dots(4)$$

$$x = 0.5 \left(1 - \frac{d_2 + d_2 d_3}{d_1 + d_2 d_3} \right) \dots\dots\dots(5)$$

$$\alpha = \frac{d_1 + d_2 + d_3 - 1}{1 - d_3} \dots\dots\dots(6)$$

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Parameter Estimation

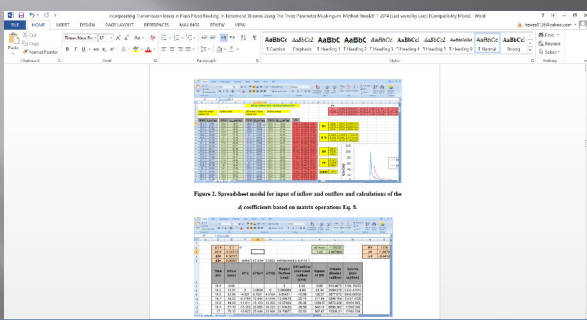
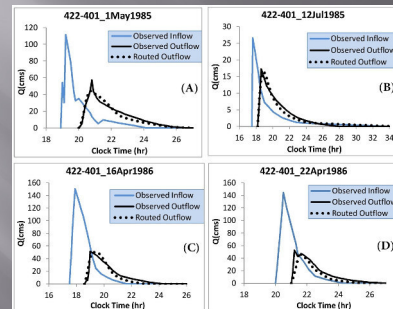


Figure 1. Spreadsheet model for input of inflow and outflow and calculation of the α coefficient based on matrix operation Eq. 5.

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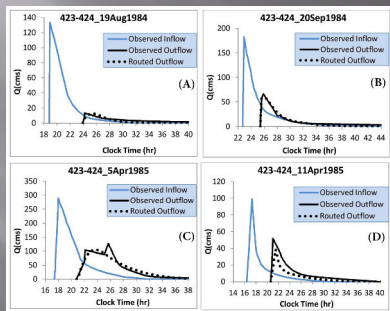
Model Results (Station 422-401)



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Model Results (Station 423-424)



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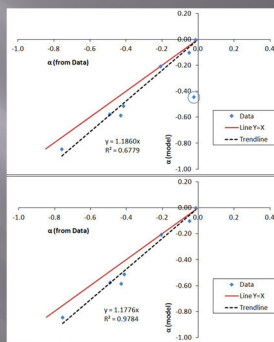
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Summary of Parameter Estimation

Reach	Date	d ₁	d ₂	d ₃	K _r (hr)	x	α (model)	α ₀ (from Data)	T _{lag} (hr)	Q _p (Obs.)	Q _p (Routed)	RMSE	K _t (hr)	K _t +K _r (hr)
Inflow 422	19May1985	0.05	0.15	0.80	0.42	-0.06	0.00	-0.02	1	17.35	16.38	0.56	0.7	1.12
Outflow 401	16Apr1986	0.22	-0.11	0.75	0.50	0.31	-0.58	-0.49	1.3	51.75	48.91	2.10	1.1	1.60
	22Apr1986	-0.23	0.30	0.82	0.15	-2.33	-0.58	-0.43	0.7	52.76	48.30	2.95	1	1.15
Average		0.02	0.11	0.81	0.43	-0.53	-0.31	-0.25	1.15	44.80	39.50	1.76	0.98	1.41
SD		0.18	0.17	0.05										
Upper Limit		0.21	0.28	0.85										
Lower Limit		-0.16	-0.07	0.76										
Inflow 423	19Aug1984	-0.01	0.02	0.90	0.88	-0.12	-0.84	-0.76	5.3	13.35	12.63	1.60	5.1	5.98
Outflow 424	20Sep1984	-0.25	0.30	0.89	0.33	-1.50	-0.51	-0.41	3.0	66.29	59.19	2.21	2.7	3.03
	5Apr1985	-0.01	0.04	0.96	2.49	-0.04	-0.21	-0.21	7.7	127.20	104.86	8.49	3.4	5.89
	11Apr1985	-0.35	0.37	0.97	1.04	-1.85	-0.44	-0.03	4.2	51.94	38.70	5.02	4.2	5.24
Average		-0.09	0.12	0.92	1.23	-0.55	-0.52	-0.46	5.33	68.94	58.89	4.10	3.73	4.97
SD		0.14	0.16	0.04										
Upper Limit		0.05	0.28	0.95										
Lower Limit		-0.23	-0.03	0.88										

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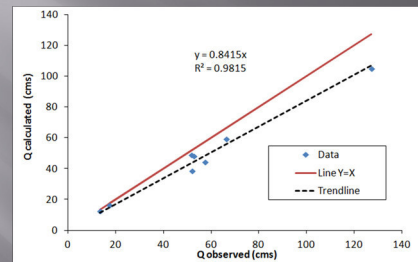
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Comparison between observed and calculated α -parameter (top) with all data points, (bottom) after omitting the extreme value as shown in Table 2 (bold face number in α -parameter columns).

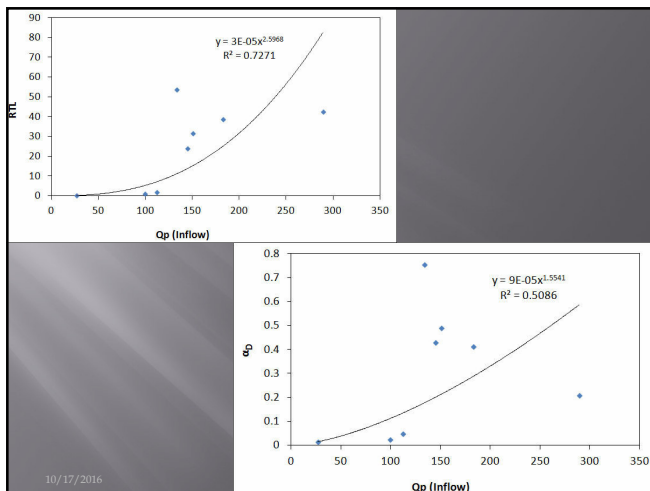
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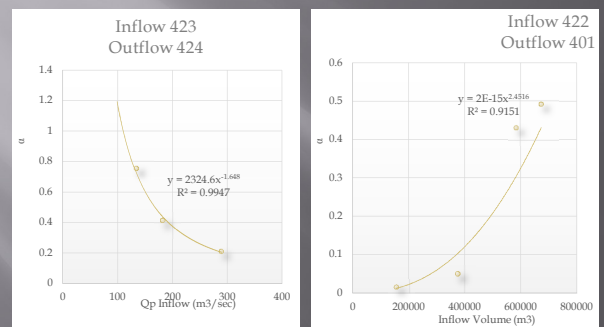
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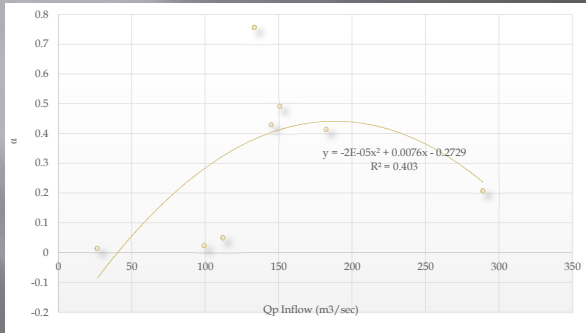
Transmission losses versus Qp



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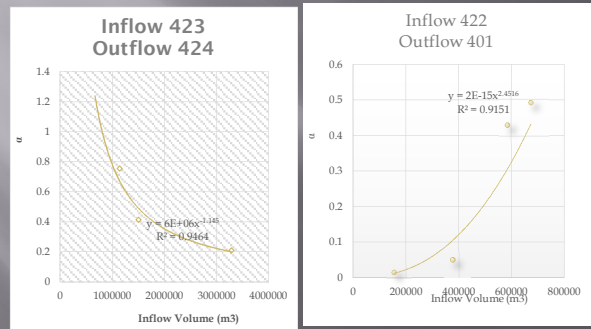
Transmission losses versus Qp



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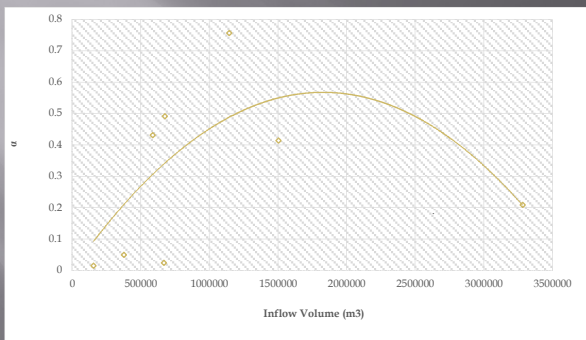
Transmission losses versus volume



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Transmission losses versus volume

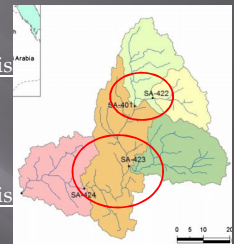


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Conclusions

- Based on the proposed model, The average percentage of transmiss
Reach 422-401 is 0.31
Reach 423-424 is 0.52
- Based on the filed data, The average percentage of transmiss
Reach 422-401 is 0.25
Reach 423-424 is 0.46
- Both reaches show different agreement in terms of flood volume and peaks, that need further investigation.



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