TITLE:

EVALUATION OF ICHNODIVERSITY BY IMAGE-RESAMPLING METHOD TO CORRECT OUTCROP EXPOSURE BIAS

AUTHOR(S):

KIKUCHI, KAZUKI; NARUSE, HAJIME; KOTAKE,
NOBUHIRO

CITATION:

KIKUCHI, KAZUKI ...[et al]. EVALUATION OF ICHNODIVERSITY BY IMAGE-
RESAMPLING METHOD TO CORRECT OUTCROP EXPOSURE BIAS.
PALAIOS 2018, 33(5): 204-217

ISSUE DATE:

2018-05-01

URL:

<http://hdl.handle.net/2433/231110>

RIGHT:

This is the accepted manuscript of the article, which has been published in final form at <https://doi.org/10.2110/palo.2017.090>; The full-text file will be made open to the public on 01 May 2019 in accordance with publisher's 'Terms and Conditions for Self-Archiving'; This is not the published version. Please cite only the published version.; この論文は出版社版ではありません。引用の際には出版社版をご確認ご利用ください。

**EVALUATION OF ICHNODIVERSITY BY IMAGE-RESAMPLING METHOD TO
CORRECT OUTCROP EXPOSURE BIAS**

KAZUKI KIKUCHI¹, HAJIME NARUSE¹, and NOBUHIRO KOTAKE²

¹*Department of Geology and Mineralogy, Graduate School of Science, Kyoto University,
Kitashirakawaoiwake-cho, Sakyo-ku, Kyoto, 606-8502, Japan*

²*Department of Earth Sciences, Graduate School of Science, Chiba University, 1-33,
Yayoi-cho, Inage-ku, Chiba, 263-8522, Japan*

email: kikuchi@kueps.kyoto-u.ac.jp

RRH: NEW EVALUATION METHOD OF ICHNODIVERSITY

LRH: K. KIKUCHI ET AL.

ABSTRACT

We propose a new method to evaluate the diversity of ichnofossils from outcrop. Ichnodiversity (defined here as the number of ichnotaxa) characterizes paleoenvironmental conditions. However, the apparent numbers of ichnotaxa observed in outcrops are significantly affected by differences in areas of exposed outcrops. This study proposes a new method to evaluate ichnodiversity, independent of outcrop exposure bias, by using an image-resampling technique combined with the shareholder quorum subsampling method. In this method, the relationship between observed and detected numbers of ichnotaxa is estimated by subsampling from existing outcrop images. The relative diversity of ichnotaxa is obtained at a given value of the estimated coverage parameter, representing the ratio of the observed number of ichnotaxa to the actual diversity. The method was verified by analyzing artificial images of ichnoassemblages, and the method successfully estimated reasonable values of relative diversity of ichnotaxa. It was also suggested that the spatial distribution patterns of ichnofossils on the bedding planes does not affect the estimated intensity of ichnodiversity when using this method. This method was also applied to field data pertaining to deposits of the submarine channel-levee complex in the Oligocene Izaki Olistolith of the Nichinan Group, southwest Japan. As a result, the ichnodiversity of the successions in the Izaki Olistolith was reconstructed to be relatively high in channel deposits and low in levee deposits.

INTRODUCTION

33   Ichnofossils represent a record of the response of ancient benthic animals to
34 changes in environmental conditions. Therefore, ichnological data are important for
35 assessing sedimentary environments of the sea floor and also to understand the ethology of
36 ancient benthic animals.

37   Ichnodiversity, which is regarded as the number of ichnotaxa at the ichnogenus
38 level, is an important and useful parameter for evaluating sedimentary environments of the
39 seafloor. Many studies have assessed not only the physicochemical disturbances including
40 sedimentation rate, energy level, sediment properties, salinity, and pore water oxygenation
41 (Bromley and Ekdale 1984; Buatois et al. 1997; Knaust 2007; Heard and Pickering 2008;
42 Hauck et al. 2009; Cummings and Hodgson 2011; Gingras et al. 2011; Phillips et al. 2011;
43 Callow et al. 2013; Heard et al. 2014; Bayet-Goll et al. 2015; Timmer et al. 2016), but also
44 biological factors such as organic matter input (Pearson and Rosenberg 1978; Wetzel and
45 Uchman 1998; Hyland et al. 2005; Callow et al. 2014). In addition, ichnodiversity is
46 strongly controlled by differences in sedimentary facies: the diversity of ichnotaxa found in
47 the submarine channel facies seems to be lower than that in the levee facies (Heard and
48 Pickering 2008; Cummings and Hodgson 2011; Phillips et al. 2011; Callow et al. 2013;
49 Heard et al. 2014). Furthermore, ichnodiversity can also provide informative data for the
50 evolutionary history of ethological strategies of benthic communities (Uchman 2003;
51 Buatois et al. 2016).

52   Previous studies have discussed ichnodiversity based on qualitative data using the
53 apparent number of ichnotaxa at the ichnogenus level obtained from outcrop observations
54 (Buatois and Mángano 2013). The apparent numbers of ichnotaxa strongly reflect not only

the actual diversity in activities of benthic animals, but also exposed areas of observed outcrops (Orr 2001). In the case of outcrop observation, the numbers of ichnotaxa must be simply increased by increasing the area of the observed bedding plane. The observational biases mentioned above can be treated as types of sample size effects. Thus, quantitative approaches are required for evaluation of ichnodiversity (Buatois et al. 2016).

Several established methods are available for correcting sample size effects in studies of ecology and paleobiology. For instance, the rarefaction method (Sanders 1968) is a traditional method to compare biodiversity among assemblages with respective sample sizes. In this method, rarefaction curves for each assemblage, which show relationships between sample size and expected species richness, are described to standardize diversity at a given sample size (e.g., Hurlbert 1971). This method is suitable when size of collected samples is significantly large and the diversity of communities is relatively low (Alroy 2010c; Chao and Jost 2012). However, this method tends to underestimate biodiversity, especially in highly diverse communities because estimated biodiversity standardized by sample size strongly depends on species-abundance distributions of the real communities (Alroy 2010c; Chao and Jost 2012). The estimated biodiversity of a community with a small number of taxa can be saturated at relatively small sample sizes, but this is not adequate for a highly diverse community. Therefore, comparison of biodiversity based on the sample size is not usually “fair” because the sample diversity of one community might cover nearly all of the real diversity, whereas the sample diversity of the other community might represent only a part of the real diversity (Chao and Jost 2012). Recently, Alroy (2010b, 2010c) proposed the shareholder quorum subsampling (SQS) method to solve this

problem. In this method, the biodiversity of communities are compared with each other at the same value of the sample coverage parameter, which represents the sum of frequencies of each taxon included in hypothetical populations (Alroy 2010a, 2010b, 2010c; Chao and Jost 2012). The sample coverage is estimated by Good-Turing frequency estimation (Good 1953) or the slope of the tangential lines of the rarefaction curves (Chao and Jost 2012). Even though the species richness value estimated by the SQS method is always lower than the real diversity, the ratio of species richness values between any two samples is expected to match the ratio of the real diversity of those communities. Thus, the SQS method can be used for estimating relative variation of biodiversity.

Ichnodiversity, however, cannot be analyzed by established methods for biodiversity described above. Although these methods require the number of individuals of each taxon for obtaining rarefaction curves or values of sample coverage, it is difficult to count the number of individuals of ichnofossils because of their morphological characteristics. The apparent number of ichnofossil specimens strongly depends on their structural organization. Ichnofossils that appear as discrete scattered spots are counted as large numbers, whereas those which make connected networks appear as a single specimen, even though they are extremely large in size. For example, planar-formed, regular network graphoglyptids such as *Megagraption* or *Paleodictyon* may be preserved as multiple spots on the bottom of a single sandstone bed. However, it is almost impossible to judge if multiple spots on a bedding plane actually represent multiple individuals or fragments of a single specimen. Even if there is intact preservation, it is difficult to recognize a single individual from planar-formed ichnofossils that occupy very large areas of the bedding

plane because the exposed region of the ichnofossil may represent only a part of the entire morphology of an individual specimen.

In previous studies dealing with ichnodiversity, most authors attempted to attenuate the effect of the outcrop exposure and observe the bedding plane as large as possible (e.g., Heard and Pickering 2008; Cummings and Hodgson 2011). In this approach, there is no criterion to judge if the area of observed outcrops is adequate for fair evaluation of ichnodiversity. Recently, three different methods were applied for evaluating ichnodiversity: the rarefaction method (Buatois et al. 2016), the ichnoabundance method (Knaust et al. 2014), and the Gini Index method (Gianetti and McCann 2010). These methods require the relative abundance of each ichnotaxon, which was estimated from the numbers reported previously (Buatois et al. 2016) or apparent numbers of sections of tubular structures on the bedding planes. However, as mentioned above, these methodologies are not suitable for estimating the abundance of ichnofossils at outcrop scale. In addition, the two methods described above are known to be problematic for estimation of biodiversity (Alroy 2010c; Chao and Jost 2012). Thus, methods for assessing ichnofossil diversity are not yet well established.

We propose a new method to evaluate ichnodiversity, independent of exposed area of outcrops, by using an image-resampling technique with the application of the SQS method. This method is applicable to images such as vertical successions, top or bottom surfaces of beds, or polished sections of sedimentary rocks. In this method, (1) the relationship between observed and detected numbers of ichnotaxa is obtained by subsampling from existing outcrop or polished section images; (2) the coverage parameter

of the sampled number of ichnotaxa and real ichnodiversity is then estimated from the fit curve of subsampled data; and (3) the sample diversity of ichnotaxa is calculated from a given value of the coverage parameter. We verified the method by applying it to artificial data of ichnoassemblages. In this test, effects of variation on distribution patterns of ichnofossils on bedding planes and shortage of outcrop exposure areas were examined. Finally, we applied the new method to field data obtained from the Oligocene submarine channel-levee complex to provide the first example of ichnological data independent of outcrop exposure bias.

METHODOLOGY OF IMAGE-RESAMPLING

We named this method “Measurement of Ichnofossil Diversity by Image-Resampling Technique” (MIDIRT) (Fig. 1). The step-wise procedure of this method is listed below.

1. Image Acquisition

First, outcrop photographs are taken in the field or laboratory. The photographs should be taken at random and perpendicular to the outcrop surfaces containing ichnofossils. Ichnofossils may be observable on various types of surfaces, such as bottom surfaces of sandstone beds or polished sections. This methodology is applicable for any type of surface, but only photographs of the same type of surface can be compared with each other. Additionally, photographs should not be taken to preferentially show specific ichnotaxa.

143

144

2. Identification of Ichnotaxa

145

146

147

148

149

150

3. Image Resampling

151

152

153

154

155

156

157

158

159

160

161

4. Curve Fitting to Resampled Data

162

163

164

Each ichnotaxon is identified from the acquired outcrop images. The regions of ichnofossils in images are colored with a specific grayscale or RGB color value for each ichnotaxon, and regions where ichnofossils are absent are colored white. Regions outside the outcrop surfaces are colored black.

The line-of-interest in data resampling is randomly set in the acquired outcrop image. If the resampling line protrudes from the outcrop, the line turns up to the next row of pixels. The number of ichnotaxa on the line is then counted. As the length of the resampling line (L) increases, the number of counted ichnotaxa (N_{is}) is expected to increase. The length of the resampling line (L) ranges from one to the maximum value of outcrop area. For each length, the resampling process is repeated 100 times, and then the mean value of N_{is} ($\overline{N_{is}}$), is regarded as the representative value of the resampled data at the given resampling length. This value shows the expected number of ichnotaxa when the outcrop is explored by the resampling line that is L in length.

Repetition of the resampling process provides the relationship between the observed length and number of ichnotaxa. Alroy (2010b, 2010c) attempted to estimate the sample coverage, which was expressed by the sum of the frequencies of taxa that

hypothetical populations included, from chances of occurrence of undiscovered taxa using Good-Turing estimation (Good 1953). This implies that the completeness of the sample (sample coverage) can be evaluated by differences between unity and chance of occurrence of undiscovered taxa. Although the chance of occurrence of undiscovered taxa cannot be directly measured, the ratio of the number of singletons in the sample when the relative frequencies of each taxon are based on a binomial distribution can be approximated. In addition, Chao and Jost (2012) indicated that sample coverage for biodiversity is equal to the slope of the tangential line of rarefaction curves in that condition. For the analytical calculation of the slope of tangential lines as an estimator of the chance of occurrence of undiscovered taxa, we employed a fitting function to the resampled data. The chance of occurrence of undiscovered taxa can be estimated from using differential calculus of the fitting function. Although the observed number of ichnotaxa ($\overline{N_{is}}$) increases as the resampling length (L) increases, the increasing rate of $\overline{N_{is}}$ is expected to decline gradually. Therefore, we employed the following as the fitting function (Mauffrey et al. 2007):

$$E_s = a_1 \ln(1 + a_2 L) \quad (1)$$

where E_s denotes the expected number of ichnotaxa and a_1 and a_2 are fitting coefficients. The parameter a_1 is a dimensionless variable, and a_2 denotes a coefficient for describing an increase rate of observed numbers of ichnotaxa, which has a dimension of the inverse number of the length scale. The obtained curve of Equation 1 can be regarded as the equivalent of rarefaction curves of biodiversity. The residual standard error of the curve

186 fitting is defined as follows:

187
$$R_{se} = \sqrt{\frac{1}{n-p} \sum_{i=1}^n R_i^2} \quad (2)$$

188 where n and p denote the number of data points and fitting coefficients ($p = 2$ in this study),
189 respectively; $n - p$ represents degrees of freedom; and R_i is the residual of the i th data
190 point. Thus, R_{se} indicates the average distance of the data points from the fitted curve,
191 which can be interpreted as the goodness of fit of the curve. Although there is no certain
192 criterion for R_{se} , an excessively large R_{se} (i.e. > 1) implies that the assumption of the
193 methodology may be violated. Therefore, application of the method should be reconsidered
194 in such cases.

195

196 5. Estimation of Ichnodiversity

197 Next, the ichnotaxa coverage parameter is defined. This parameter describes
198 completeness of sampled diversity. Here, the slope of the tangential lines of Equation 1 at
199 any given L (S_L) is calculated as follows:

200
$$S_L = \frac{a_1 a_2}{1 + a_2 L} \quad (3)$$

201 Chao and Jost (2012) argued that the sample coverage (C) can be estimated as follows:

202
$$C = 1 - S_L \quad (4)$$

203 The sample coverage (C) becomes 1 when sampled diversity is equivalent to the real
204 diversity, and thus the slope of the rarefaction curve is zero, and is expected to be positive
205 because the parameter C is the proportion of the total number of individuals in the

206 hypothetical population. However, the slope of the curve of Equation 1 (S_L) is not always
207 less than 1, so that C may be negative when Equation 4 is used. Therefore, the slope S_L is
208 normalized by the slope of Equation 1 at $L = 0$ (S_0):

$$209 \quad \overline{S}_L = \frac{S_L}{a_1 a_2} \quad (5)$$

210 For the biodiversity rarefaction curve, the number of taxa always becomes one when the
211 first specimen is sampled, and thus the slope of the discretized biodiversity rarefaction
212 curve is 1 at the interval where the number of samples is from 0 to 1. Because of the
213 normalization represented by Equation 5, the normalized slope of Equation 1 becomes 1 at
214 $L = 0$. Here the ichnotaxa coverage parameter $\overline{C}$ is defined as follows:

$$215 \quad \overline{C} = 1 - \overline{S}_L \quad (6)$$

216 The parameter $\overline{C}$ becomes the minimum value 0 when $L = 0$, and converges to 1 as L
217 increases. Equation 6 can be recalculated as:

$$218 \quad \overline{C} = \frac{a_2 L}{1 + a_2 L} \quad (7)$$

219 Equation 1 is recalculated as follows, with the aid of Equation 7:

$$220 \quad E_s = a_1 \ln \left(1 + \frac{\overline{C}}{1 - \overline{C}} \right) \quad (8)$$

221 Equation 8 indicates that the ratio of E_s between two samples at any given $\overline{C}$ is always
222 constant. Consequently, fluctuation of the ichnodiversity in the outcrop image data at any
223 given ichnotaxa coverage parameter is obtained independent of differences in the exposed

224 area of outcrops.

225

226 VERIFICATION OF METHODOLOGY

227 The method proposed in this study was applied to artificial data of
228 ichnoassemblages to verify the methodology, specifically the effects of distribution patterns
229 of ichnofossils on bedding planes and shortage of outcrop exposure area. Four artificial
230 outcrop images were produced with different distribution patterns and various types of
231 ichnofossils. In addition, incomplete outcrop images were also produced by gradually
232 decreasing the area of these four outcrop images. Then, the ichnodiversity estimated from
233 the artificial images was compared with the true values for methodology verification.

234

235 Artificial Data

236 Artificial ichnofossil images were allocated on virtual bedding planes to generate
237 artificial outcrop images showing ichnoassemblages. Chance of occurrence of each
238 ichnotaxon was set to the prescribed value (Fig. 2). In this series of experiments, two types
239 of spatial distribution patterns of ichnofossils were examined: uniform and patchy
240 distributions. In addition, two types of maximum number of ichnotaxa were also examined
241 (10 and 5 ichnotaxa) (Fig. 3). The procedures of allocating the ichnofossil images were as
242 follows: (1) 150 ichnofossil images were chosen based on their chance of occurrence (Fig.
243 2). (2) Ichnofossil images were allocated onto a white colored image (6000 × 4500 pixels).
244 The coordinates of each ichnofossil image were determined by the following equation:

$$p_p = \left(\frac{d_{\min}}{D} \right)^{2k} \quad (9)$$

where p_p denotes the probability of whether an ichnofossil image is allocated at a point (p) that was chosen by a uniform random number; D is a specific distance from the point p ($D = 200$ pixels in this study); $d_{\min}$ is the minimum value of the distances between p and other points the ichnofossil images were already allocated; and k is a coefficient that determines the distribution pattern of the ichnofossil images ($k = -1, 1$). When $k = 1$, the point p that is far from other points in which the ichnofossil images were already allocated tends to be adopted, and thus the distribution pattern becomes uniform. In contrast, when $k = -1$, the point p that is close to other points tends to be adopted so that the ichnofossil images are allocated in proximity with each other and the patchy distribution pattern is established. Resolution of each image was set at 50 pixels/cm, thus, the maximum area of the outcrop images was $10,800 \text{ cm}^2$. All ichnofossil images were 100×100 pixels ($2 \times 2 \text{ cm}$).

These outcrop images were then partially and progressively covered by black coloration to produce images of the reduced areas, which were analyzed in order to verify the effects on areas with outcrop exposure.

261

262 Results

Results of the image-resampling method are summarized in Fig. 4. After 100 repetitions of the resampling process, the mean number of observed ichnotaxa ($\overline{N_{is}}$)

against each length of the resampling line (L) was estimated. The 95% confidence intervals for each $\overline{N}_{is}$ were calculated by bootstrapping, replicating 10,000 times, with normal approximation. Confidence intervals of the estimated numbers of ichnotaxa were larger in patchy distribution patterns than in uniform distribution patterns. For all distribution patterns, the estimated numbers of ichnotaxa increased as the lengths of resampling lines increased, and they approached their maximum numbers of ichnotaxa when the resampling lines filled the whole outcrop images. As a whole, the relationship between the lengths of resampling lines and the estimated number of ichnotaxa were well-fitted to the function expressed in Equation 1. Parameters of curve fitting are summarized in Table 1.

As a result of analyses, the fitting coefficients a_1 and a_2 were estimated as follows, respectively: 1.726 and 0.036 in the uniform distribution with 10 ichnotaxa; 0.835 and 0.047 in the uniform distribution with 5 ichnotaxa; 1.797 and 0.029 in the patchy distribution with 10 ichnotaxa; and 0.900 and 0.031 in the patchy distribution with 5 ichnotaxa (Table 1). Using these values, the coverage parameters ($\overline{C}$) were calculated. The relationships between the coverage parameters ($\overline{C}$) and the expected numbers of ichnotaxa (E_s) are shown in Fig. 5. The shapes of curves, which were based on data of equivalent maximum number of ichnotaxa, were similar to each other. Ratios of E_s among artificial data are shown in Table 2.

The results of the numerical experiments for reduced outcrop exposure are shown in Figure 6 and Table 1. The shapes of the $\overline{C}$ -based ichnofossil rarefaction curves were well maintained if the exposure area decreased, especially in diverse artificial data (uniform

and patchy distribution patterns with 10 ichnotaxa) (Fig. 6A, C). In contrast, the E_s in the case with 5 ichnotaxa tended to be under- or overestimated when the total exposure area was smaller than 50% of the original image (Fig. 6B, D; Table 1).

APPLICATION TO FIELD DATA

The MIDIRT method was applied to the field data measured in deposits of the submarine channel-levee complex in the Oligocene Izaki Olistolith of the Nichinan Group (Sakai et al. 1987). The variation of ichnodiversity in the channel-fill deposit and the levee deposit was evaluated with MIDIRT.

Geological Setting

The Oligocene to lower Miocene Nichinan Group is distributed on the southeastern part of Kyushu, southwestern Japan (Fig. 7). The Nichinan Group is composed of various sized coherent blocks and intensely deformed beds. They are interpreted as the deposits of the olistostrome which was caused by gravitational instability from the restart of subduction of the Philippine Sea plate in 21–17 Ma (Sakai 1988a, 1988b, 1988c). The Izaki Olistolith distributed in Izaki-bana is considered as one of the coherent blocks that were originally deposited in the deep-sea setting (Sakai et al. 1987) (Fig. 7). It is mainly composed of alternating beds of turbidite sandstone and mudstone and is interpreted to be a deposit of a submarine channel-levee complex (Yumi and Ishihara 2012). The Izaki Olistolith can be divided into three stratigraphic units based on lithology. The lower and upper parts of the Izaki Olistolith (units A and C; Fig. 7D) are comprised of thin-bedded

turbidite sandstone and mudstone beds. The sandstone beds in the units A and C are mainly 1–20 cm thick, and climbing ripple and convolute lamination are observable. The mudstone beds are 10–20 cm thick. Taking this into consideration, the deposits of the units A and C are interpreted as submarine levee deposits (Arnott 2010). In contrast, the middle part of the Izaki Olistolith (unit B; Fig. 7D) consists of thick-bedded turbidite sandstones and thin-bedded mudstones. The thickness of the sandstone beds in the unit B ranges from 5 to 200 cm. The current ripple, climbing ripple, convolute lamination, and parallel lamination are observable on the top of sandstone beds. The sole marks, such as flute cast or groove cast, are commonly found on the bottom surface of thick-bedded (more than 100 cm thick) sandstone beds. The mudstone beds in the middle part are less than 10 cm thick. The alternating beds in the unit B show an upward-thinning succession. These characteristics indicate that the unit B of the Izaki Olistolith is the submarine channel-fill deposit (Arnott 2010).

Ichnoassemblage of the Izaki Olistolith

The ichnoassemblage in the Izaki Olistolith is mainly composed of graphoglyptids (Fig. 8). This study investigated the number of ichnogenera and measured the exposed area of the bottom surfaces of each turbidite sandstone bed. For the image-resampling method, photographs of the bottom surfaces of sandstone beds were obtained through field work, and then were colored appropriately for each ichnogenus. We observed 5,960 cm² and 33,520 cm² of bottom surfaces of the turbidite sandstone beds in the levee deposit and the channel-fill deposit, respectively.

A total of 11 ichnogenera were recognized on the sole surface of sandstone beds in the levee deposit (Table 3). Abundant *Phycosiphon incertum* and *Gordia marina* are characteristics of this deposit. Graphoglyptids, such as *Megagraption irregulare* or *Paleodictyon strozzii*, were rarely observed in relatively thin-bedded sandstones (5–15 cm thick). In contrast, the thick-bedded sandstones (50–200 cm thick) contain fewer ichnogenera even though the bottom surfaces of these beds are widely exposed.

A total of 22 ichnogenera were found on the sole surface of sandstone beds in the channel-fill deposit (Table 3). Various types of graphoglyptids were observed in these deposits. *Helminthorhapse japonica* and *Paleodictyon strozzii* were common. *Desmograption inversum*, *Punctorhapse parallela*, and *Spirorhapse involuta* are rarely observed. Thick-bedded sandstones in the lower part of the channel-fill deposits (unit B1; Fig. 7D) yielded fewer ichnogenera, whereas various ichnogenera occurred in the upper part (unit B2; Fig. 7D).

Results

Results of our image-resampling method are summarized in Fig. 9. As with the artificial data, the numbers of ichnogenera in each sedimentary environment increased as observed area increased. The plots were well-fitted to Equation 1. The residual standard error (R_{se}) was 0.180 in the channel-fill deposit and 0.300 in the levee deposit (Table 4).

The fitting coefficients a_1 and a_2 in the channel-fill deposit were 4.916 and 0.003, respectively. In contrast, those in the levee deposit were 2.825 and 0.007, respectively (Table 4). Using these values, the coverage parameters ($\bar{C}$) were calculated.

The relationship between the coverage parameters ($\bar{C}$) and the expected numbers of ichnotaxa (E_s) were then estimated (Fig. 10). E_s , when $\bar{C} = 0.8$ in the channel-fill and levee deposits was 7.913 and 4.547, respectively. Therefore, the ichnodiversity in the channel-fill deposit was 1.740 times higher than in the levee deposit.

DISCUSSION

Selection of the Fitting Function

This study employed Equation 1 as the fitting function of the rarefaction curves, according to Mauffrey et al. (2007). Mauffrey et al. (2007) used and evaluated three models, which included the Exponential Dependence model (recast to Equation 1), Clench model, and Linear Dependence model, to fit the individual-based rarefaction curves for species richness extrapolation of the small-mammal communities in a French Guianan rainforest. They concluded that Equation 1 was the most suitable for extrapolation of the rarefaction curves because the estimated diversity based on Equation 1 showed the most similar value to the known local biodiversity in their study area based on previous trapping missions, although goodness of fit for these three models was not significantly different. In contrast, van Rooijen (2009) estimated snake species richness of the Santubong Peninsula in Borneo by extrapolating the individual-based rarefaction curve. He applied two exponential models to fit the rarefaction curve, the negative exponential and Weibull functions. The Weibull function was recast as the following equation:

$$Y = A(1 - \exp(b - ct)) \quad (10)$$

where Y denotes the expected number of species; A is total number of species; t is sample size; and b and c are constants that denote the ease which species are found. van Rooijen (2009) argued that the Weibull function exhibited higher goodness of fit than the negative exponential function, and expected species richness based on the Weibull function corresponded with the value estimated by the Chao I estimator (Chao 1984). Therefore, he concluded that the Weibull function is suitable for extrapolation of the rarefaction curve.

For ichnodiversity, there are some problems in fitting the Weibull function to the relationship between L and E_s . First, goodness of fit of the Weibull function is lower than that of Equation 1. The Akaike's Information Criterion (AIC) for the uniform distribution with 10 ichnotaxa was 24.44 when fitted to the Weibull function, whereas 1.38 when fitted to Equation 1. Figure 11 shows the result of fitting the Weibull function for uniform distribution with 10 ichnotaxa. The fitted Weibull function tended to overestimate when L was an intermediate value (approximately 4,000 to 5,000 cm) and underestimate when L was small or large (Fig. 11). Second, the asymptote of the Weibull function ($Y = A$) strongly depended on the observed maximum number of ichnotaxa. The maximum number of ichnotaxa estimated by the fitted Weibull function was 9.701 for 100% exposed artificial data of the uniform distribution with 10 ichnotaxa. In contrast, it was 8.224 in for 50% exposed data in which exactly 8 ichnotaxa were observed. Therefore, estimation using the Weibull function is unsuitable to correct outcrop exposure bias.

Application of Equation 1 leads an infinite value of ichnodiversity when $L \rightarrow \infty$ or $\bar{C} = 1$ because Equation 1 does not have an asymptote (eq. 1 and 8). This implies that ichnodiversity in the hypothetical population cannot be estimated by extrapolation of

Equation 1. However, E_s standardized by $\bar{C}$ in the interpolation interval can be compared fairly, regardless of the difference in distribution pattern and outcrop exposure bias (see below for discussion). Furthermore, the ratio of E_s among any two datasets was always constant at any given $\bar{C}$ (eq. 8). Taking this into consideration, Equation 1 is valid as the fitting function for the relationship between L and E_s , if the analysis aim is comparison of E_s in the interpolation interval and not extrapolation.

Validity of the Method

Verification of the method by applying artificial outcrop image data indicated that the MIDIRT is a valuable method to evaluate ichnodiversity regardless of outcrop exposure area, the total number of ichnotaxa, or distribution patterns of ichnofossils on the bedding planes. Each $\bar{C}$ -based fitting curve of artificial data corresponded to another curve which had an equal maximum number of ichnotaxa (Fig. 5). As the $\bar{C}$ -based fitting curves are the logarithmic functions passing an origin (eq. 8), ratios of E_s in each artificial image are constant even if E_s is normalized at any ichnotaxa coverage parameter. For example, the ratio of E_s in the uniform distribution pattern with 10 ichnotaxa and uniform distribution pattern with 5 ichnotaxa was constantly 2.067 (Table 2). In contrast, the ratio of the maximum numbers of ichnotaxa in these conditions was 2.000. Therefore, it was considered that the ratio of E_s was reflected in the maximum numbers of ichnotaxa in each condition. In other combinations of conditions, the ratio of E_s also showed the same

tendency (Table 2). The $\bar{C}$ -based fitting curves showed that E_s obtained from the images showing patchy distribution patterns with 10 ichnotaxa were slightly overestimated when compared with E_s in conditions with equivalent maximum number of ichnotaxa (Fig. 5; Table 2). The increasing rate of observed number of ichnotaxa (a_2) in the patchy distribution with 10 ichnotaxa was smaller than that of other conditions (Table 1). This was likely due to a larger area of bedding plane in which no ichnotaxa were observed, and thus, $\bar{C}$ at a given L was also smaller. This may have led to the overestimation of E_s in the patchy distribution with 10 ichnotaxa. However, differences between the ratios of E_s and that of the maximum numbers of ichnotaxa were small enough to disregard. These results indicate that ichnodiversity can be compared fairly, independent of outcrop exposure bias using MIDIRT.

Application to outcrops

We tested our methodology to the outcrops of the channel-levee system in the Izaki Olistolith, and suggested that the effect of outcrop exposure bias cannot be ignored. The result of our method shows that the ratio of E_s in the channel-fill deposit and levee deposit is 1.740 although that of the raw numbers of ichnotaxa is higher value (2.000). In the case of the Izaki Olistolith, the channel-fill deposits expose better than the levee deposits, which leads to the apparently larger ichnodiversity. The decrease in the ratio of the ichnodiversity indicates that, therefore, our method corrected this bias.

In general, however, it is estimated that the ichnodiversity in channel-fill deposits

tends to be underestimated than in levee deposits. The number of observable beds in a certain interval is generally larger in levee deposits than in channel-fill deposits because levee deposits are composed of thin-bedded turbidite sandstone beds in which erosive bottom surfaces are rare. On the other hand, channel-fill deposits are dominated by thick-bedded turbidite sandstone beds which often show amalgamated bottom surfaces. Therefore, the observed number of ichnotaxa in channel-fill deposits is expected to be smaller than in levee deposits. This kind of outcrop exposure bias in each sedimentary environment may have not been adequately considered in previous studies, and therefore our methodology is significant for evaluating not only the temporal variation of the paleoenvironments but also the comparison of ichnodiversities over various lithofacies.

In contrast, it is generally considered that the repeated scouring on channel floors may negatively affect the establishment of K-selected populations such as those responsible of producing graphoglyptids (e.g. Heard and Pickering 2008; Cummings and Hodgson 2011; Heard et al. 2014). It implies a possibility that the higher ichnodiversity in channel-fill deposits than that of levee deposits may be a peculiar phenomenon in the Izaki Olistolith. Further investigations of ichnodiversity in channel-levee complexes independent of outcrop exposure bias will clarify the inconsistency between the trend in previous works and that in the Izaki Olistolith.

Future application of the method

Further application of our method to ichnological data around the world will reveal the paleoenvironmental variation through the evaluation of the ichnodiversity eliminating

the effect of observation bias. Wetzel (1991) suggested that pore water oxygenation level and benthic food content strongly affect ichnodiversity. Cummings and Hodgson (2011) argued that fluctuation in number of ichnotaxa in submarine fan deposits of the Basque Basin, northern Spain was induced by dysoxic/anoxic conditions. In addition, biodiversity of benthic communities can be affected by benthic food content on the seafloor. For instance, it is well known that biodiversity increases are associated with decreasing organic matter input, whereas biodiversity decreases in oversupplied organic matter conditions (P-R model; Pearson and Rosenberg 1978). However, ichnodiversity does not directly correspond with benthic biodiversity (Buatois and Mángano 2013), so that further examination is needed for understanding relationship between benthic food content and the resultant variation in ichnodiversity.

Although there are some issues and room for development, the MIDIRT method is applicable for both characterizing depositional facies and temporal variations of ichnodiversity. There have been attempts to construct the model for estimation of paleoenvironmental conditions based on fluctuation of ichnodiversity (e.g. Heard and Pickering 2008; Cummings and Hodgson 2011; Phillips et al. 2011; Callow et al. 2013). These models will be more reliable with fair comparison of ichnodiversity by the MIDIRT method.

CONCLUSIONS

We proposed a new method, MIDIRT, to evaluate ichnodiversity from outcrop records. Although the number of ichnotaxa is generally affected by outcrop exposure bias,

existing methods to correct sample size bias cannot be applied to ichnofossil analyses as the relative abundance of ichnofossils is difficult to measure because of the variations in morphological characteristics. The method we proposed corrects this bias by using an image-resampling technique combined with the SQS method (Alroy 2010b, 2010c). The method was verified by applying it to four types of artificial data. As a result, the ratio of the estimated sample ichnodiversity approximated the ratio of the real ichnodiversity with each dataset. Results also suggested that ichnodiversity can be compared fairly, regardless of area of outcrop exposure and distribution patterns of ichnofossils on the bedding planes. The method was also applied to field data of the ichnoassemblage in the channel-levee complex of the Oligocene Izaki Olistolith of the Nichinan Group. The result of the analysis indicated that ichnodiversity, independent of outcrop exposure bias, was higher in the channel-fill deposit than the levee deposit. In contrast, previous studies showed an inverse trend of increasing ichnodiversity from channel-axial to marginal environments. Evaluation of ichnodiversity by the MIDIRT method is expected to be useful in reconstructing paleoenvironmental conditions.

ACKNOWLEDGEMENTS

Takao Ubukata (Kyoto University) provided useful comments on statistics and body fossil diversity. The paper benefited from constructive reviews by two anonymous referees, as well as from numerous editorial comments by M. Gabriela Mángano. A part of this study was supported by a grant from the Fukada Geological Institute (2014). All these contributions are gratefully acknowledged.

REFERENCES

- ALROY, J., 2010a, The shifting balance of diversity among major marine animal groups: Science, v. 329, p. 1191–1194, doi: 10.1126/science.1189910.
- ALROY, J., 2010b, Geographical, environmental and intrinsic biotic controls on Phanerozoic marine diversification: Paleontology, v. 53, p. 1211–1235, doi: 10.1111/j.1475-4983.2010.01011.x.
- ALROY, J., 2010c, Fair sampling of taxonomic richness and unbiased estimation of origination and extinction rates, in J. Alroy and G. Hunt (eds.), Quantitative Methods in Paleobiology: The Paleontological Society, Boulder, p. 55–80.
- ARNOTT, R.W.C., 2010, Deep-marine sediments and sedimentary system, in N.P. James and R.W. Dalrymple (eds.), Facies Models 4, Geological Association of Canada, St. John's, NL, Canada, p. 295–322.
- BAYET-GOLL, A., NETO DE CARVALHO, C., MAHMUDY-GHARAEI, M.H., AND NADAF, R., 2015, Ichnology and sedimentology of a shallow marine Upper Cretaceous depositional system (Neyzar Formation, Kopet-Dagh, Iran): palaeoceanographic influence on ichnodiversity: Cretaceous Research, v. 56, p. 628–646, doi: 10.1016/j.cretres.2015.07.008.
- BROMLEY, R.G. AND EKDALE, A. A., 1984, *Chondrites*: a trace fossil indicator of anoxia in sediments: Science, v. 224, p. 872–875, doi: 10.1126/science.224.4651.872.
- BUATOIS, L.A. and MÁNGANO, M.G., 2013, Ichnodiversity and ichnodisparity: significance and caveats: Lethaia, v. 46, p. 281–292, doi: 10.1111/let.12018.

- 524 BUATOIS, L.A., MANGANO, M.G., AND MAPLES, C.G., 1997, The paradox of nonmarine
525 ichnofaunas in tidal rhythmites; integrating sedimentologic and ichnologic data from
526 the Late Cretaceous of eastern Kansas, USA: *PALAIOS*, v. 12, p. 467–481, doi:
527 10.2307/3515384.
- 528 BUATOIS, L.A., MANGANO, M.G., OLEA, R.A., and WILSON, M.A., 2016, Decoupled
529 evolution of soft and hard substrate communities during the Cambrian Explosion and
530 Great Ordovician Biodiversification Event: *Proceedings of the National Academy of*
531 *Sciences*, v. 113, p. 6945–6948, doi: 10.1073/pnas.1523087113.
- 532 CALLOW, H.T.R., KNELLER, B., DYKSTRA, M., and MCILROY, D., 2014, Physical, biological,
533 geochemical and sedimentological controls on the ichnology of submarine canyon and
534 slope channel systems: *Marine and Petroleum Geology*, v. 54, p. 144–166, doi:
535 10.1016/j.marpetgeo.2014.02.016.
- 536 CALLOW, H.T.R., MCILROY, D., KNELLER, B., and DYKSTRA, M., 2013, Integrated
537 ichnological and sedimentological analysis of a Late Cretaceous submarine
538 channel-levee system: the Rosario Formation, Baja California, Mexico: *Marine and*
539 *Petroleum Geology*, v. 41, p. 277–294, doi: 10.1016/j.marpetgeo.2012.02.001.
- 540 CHAO, A. and JOST, L., 2012, Coverage-based rarefaction and extrapolation: standardizing
541 samples by completeness rather than size: *Ecology*, v. 93, p. 2533–2547, doi:
542 10.1890/11-1952.1.
- 543 CHAO, A., 1984, Nonparametric estimation of the number of classes in a population:
544 *Scandinavian Journal of Statistics*, v. 11, p. 265–270, doi: 10.2307/4615964.
- 545 CUMMINGS, J.P. and HODGSON, D.M., 2011, Assessing controls on the distribution of

- 546 ichnotaxa in submarine fan environments, the Basque Basin, Northern Spain:
547 Sedimentary Geology, v. 239, p. 162–187, doi: 10.1016/j.sedgeo.2011.06.009.
- 548 GIANNETTI, A. AND MCCANN, T., 2010, The upper Paleocene of the Zumaya section
549 (northern Spain): review of the ichnological content and preliminary palaeoecological
550 interpretation: Ichnos, v. 17, p. 137–161, doi: 10.1080/10420941003659550.
- 551 GINGRAS, M.K., MACEachern, J.A., AND DASHTGARD, S.E., 2011, Process ichnology and
552 the elucidation of physico-chemical stress: Sedimentary Geology, v. 237, p. 115–134,
553 doi: 10.1016/j.sedgeo.2011.02.006.
- 554 GOOD, I.J., 1953, The population frequencies of species and the estimation of population
555 parameters: Biometrika, v. 40, p. 237–264, doi: 10.1093/biomet/40.3-4.237.
- 556 HAUCK, T., DASHTGARD, S.E., PEMBERTON, S.G., AND GINGRAS, M.K., 2009,
557 Brackish-water ichnological trends in a microtidal barrier island–embayment system,
558 Kouchibouguac National Park, New Brunswick, Canada: PALAIOS, v. 24, p. 478–496,
559 doi: 10.2110/palo.2008.p08-056r.
- 560 HEARD, T.G. and PICKERING, K.T., 2008, Trace fossils as diagnostic indicators of
561 deep-marine environments, middle Eocene Ainsa-Jaca basin, Spanish Pyrenees:
562 Sedimentology, v. 55, p. 809–844, doi: 10.1111/j.1365-3091.2007.00922.x.
- 563 HEARD, T.G., PICKERING, K.T., and CLARK, J.D., 2014, Ichnofabric characterization of a
564 deep-marine clastic system: a subsurface study of the middle Eocene Ainsa System,
565 Spanish, Pyrenees: Sedimentology, v. 61, p. 1298–1331, doi: 10.1111/sed.12101.
- 566 HURLBERT, S.H., 1971, The nonconcept of species diversity; a critique alternative
567 parameters: Ecology, v. 52, p. 577–586, doi: 10.2307/1934145.

- 568 HYLAND, J., BALTHIS, L., KARAKASSIS, I., MAGNI, P., PETROV, A., SHINE, J., VESTERGAARD,
569 O., AND WARWICK, R., 2005, Organic carbon content of sediments as an indicator of
570 stress in the marine benthos: *Marine Ecology Progress Series*, v. 295, p. 91–103, doi:
571 10.3354/meps295091.
- 572 KNAUST, D., 2007, Invertebrate trace fossils and ichnodiversity in shallow-marine
573 carbonates of the German Middle Triassic (Muschelkalk), *in* R.G. Bromley, L.A.
574 Buatois, M.G. Mángano, J.F. Genise, and R.N. Melchor (eds.), *Sediment-Organism*
575 *Interactions: A Multifaceted Ichnology*: SEPM Special Publication, v. 88. p. 221–238.
- 576 KNAUST, D., WARCHOL, M., AND KANE, I.A., 2014, Ichnodiversity and ichnoabundance:
577 revealing depositional trends in a confined turbidite system: *Sedimentology*, v. 61, p.
578 2218–2267, doi: 10.1111/sed.12134.
- 579 MAUFFREY, J. F., STEINER, C. AND CATZEFLIS, F. M., 2007, Small-mammal diversity and
580 abundance in a French Guianan rain forest: test of sampling procedures using species
581 rarefaction curves: *Journal of Tropical Ecology*, v. 23, p. 419–425, doi:
582 10.1017/S0266467407004154.
- 583 ORR, P. J., 2001, Colonization of the deep-marine environment during the early
584 Phanerozoic: the ichnofaunal record: *Geological Journal*, v. 36, p. 265–278, doi:
585 10.1002/gj.891.
- 586 PEARSON, T.H. AND ROSENBERG, R., 1978, Macrobenthic succession in relation to organic
587 enrichment and pollution of the marine environment: *Oceanography and Marine*
588 *Biology: An Annual Review*, v. 16, p. 229–311.
- 589 PHILLIPS, C., MCILROY, D., and ELLIOTT, T., 2011, Ichnological characterization of

- 590 Eocene/Oligocene turbidites from the Grès d'Annot Basin, French Alps, SE France:
591 Palaeogeography, Palaeoclimatology, Palaeoecology, v. 300, p. 67–83, doi:
592 10.1016/j.palaeo.2010.12.011.
- 593 SAKAI, H., 1988a, Toi-misaki olistostrome of the southern belt of the Shimanto Terrane,
594 south Kyushu—I. Reconstruction of depositional environments and stratigraphy
595 before collapse: Journal of Geological Society of Japan, v. 94, p. 733–747, doi:
596 10.5575/geosoc.94.733. (In Japanese).
- 597 SAKAI, H., 1988b, Toi-misaki olistostrome of the southern belt of the Shimanto Terrane,
598 south Kyushu—II. Deformation structure of huge submarine slides and processes of
599 formation: Journal of Geological Society of Japan, v. 94, p. 837–853 doi:
600 10.5575/geosoc.94.837. (In Japanese).
- 601 SAKAI, H., 1988c, Origin of the Misaki Olistostrome Belt and re-examination of the
602 Takachiho Orogeny: Journal of Geological Society of Japan, v. 94, p. 945–961 doi:
603 10.5575/geosoc.94.945. (In Japanese).
- 604 SAKAI, T., KUSABA, T., NISHI, H., KOMORI, M., and WATANABE, M., 1987, Olistostorome of
605 the Shimanto Terrane in the Nichinan area, southern part of the Miyazaki Prefecture,
606 south Kyushu—with reference to deformation and mechanism of emplacement of
607 olistolith: Science Reports, Department of Geology, Kyushu University, v. 15, p. 167–
608 199. (In Japanese).
- 609 SANDERS, H.L., 1968, Marine benthic diversity: a comparative study: The American
610 Naturalist, v. 102, p. 243–282, doi: 10.1086/282541.
- 611 TIMMER, E.R., BOTTERILL, S.E., GINGRAS, M.K., AND ZONNEVELD, J.-P., 2016, Visualizing a

- 612 process ichnology dataset, Lower Cretaceous McMurray Formation, NE Alberta,
613 Canada: Bulletin of Canadian Petroleum Geology, v. 64, p. 251–265, doi:
614 10.2113/gscpgbull.64.2.251.
- 615 UCHMAN, A., 2003, Trends in diversity, frequency and complexity of graphoglyptid trace
616 fossils: evolutionary and palaeoenvironmental aspects: Palaeogeography,
617 Palaeoclimatology, Palaeoecology, v. 192, p. 123–142, doi:
618 10.1016/S0031-0182(02)00682-X.
- 619 VAN ROOIJEN, J., 2009, Estimating the snake species richness of the Santubong Peninsula
620 (Borneo) in two different ways: Contributions to Zoology, v. 78, p. 141–147.
- 621 WETZEL, A., 1991, Ecologic interpretation of deep-sea trace fossil communities:
622 Palaeogeography, Palaeoclimatology, Palaeoecology, v. 85, p. 47–69, doi:
623 10.1016/0031-0182(91)90025-M.
- 624 WETZEL, A., AND UCHMAN, A., 1998, Deep-sea benthic food content recorded by
625 ichnofabrics; a conceptual model based on observations from Paleogene flysch,
626 Carpathians, Poland: Palaios, v. 13, p. 533–546, doi: 10.2307/3515345.
- 627 YUMI, M. AND ISHIHARA, Y., 2012, Characterization of erosional marks in the base of
628 sediment-gravity-flow deposits: special reference to the effect of flow duration for
629 flute mark formation: Journal of Sedimentological Society of Japan, v. 71 p. 173–190,
630 doi: 10.4096/jssj.71.173.
- 631
- 632 Figure and Table Captions
- 633 FIG. 1.—Schematic diagram of Measurement of Ichnofossil Diversity by

Image-Resampling Technique (MIDIRT) procedures. The procedures are as follows:

(1) outcrop images containing trace fossils are acquired; (2) each identified ichnotaxon is illustrated with a particular gray scale value in outcrop images; (3a) one of the outcrop images is selected at random, and (3b) the line of interest that has a given length L is randomly set in the image. Then the number of types of gray scale values are counted along the line of interest; (3c) the processes of resampling (3a, b) are repeated 100 times, and the results are averaged for obtaining an expected number of ichnotaxa corresponding to the length of resampling line L . (4) The relationship between length of the line-of-interest (approximated to observed area) L and expected number of ichnotaxa E_s is estimated by repeating processes (3a–c) with changing sampling length L ; and (5) the coverage parameter, which is an estimate of the ratio of a measured value to actual diversity, is calculated from the slope of the tangential lines of the curve. See details of these procedures in the text.

FIG. 2.—List of illustrated symbols of ichnofossils and their chance of occurrence.

FIG. 3.—Four artificial outcrop images that were generated for verification of the MIDIRT method proposed in this study. **A)** Uniform distribution pattern with 10 ichnotaxa. **B)** Uniform distribution pattern with 5 ichnotaxa. **C)** Patchy distribution pattern with 10 ichnotaxa. **D)** Patchy distribution pattern with 5 ichnotaxa.

FIG. 4.—Relationships between resampling length and number of detected ichnotaxa, based on analysis of artificial outcrop images by MIDIRT. Each plot represents the average number of detected ichnotaxa that was obtained from 100 trials using the given resampling length. The fitted curves and coefficients of determination are also shown.

Error bars indicate the 95% confidence intervals of the average numbers of detected ichnogenera. **A)** Result from the image of uniform distribution pattern with 10 ichnotaxa. **B)** Result from the image of uniform distribution pattern with 5 ichnotaxa. **C)** Result from the image of patchy distribution pattern with 10 ichnotaxa. **D)** Result from the image of patchy distribution pattern with 5 ichnotaxa.

FIG. 5.—Expected number of ichnotaxa E_s against the ichnotaxa coverage parameter used in the MIDIRT method. Shapes of fitted curves for E_s only depend on the actual number of ichnotaxon, and are independent of spatial patterns of ichnofossil distribution.

FIG. 6.— $\bar{C}$ -based fitting curves of the artificial outcrop images showing the effect of outcrop exposure bias based on the result of MIDIRT analysis. **A)** Uniform distribution pattern with 10 ichnotaxa. **B)** Uniform distribution pattern with 5 ichnotaxa. **C)** Patchy distribution pattern with 10 ichnotaxa. **D)** Patchy distribution pattern with 5 ichnotaxa.

FIG. 7.—Maps showing the study area. **A)** Location of study area. **B)** Geological map of the southern part of the Nichinan coastal area, southern part of Kyushu, southwest Japan. Modified after Sakai (1988c). **C)** Lithological map of the Izaki-bana. In addition to coherent alternating beds of turbidite sandstone and mudstone, intensely deformed slumped beds are widely distributed on the Izaki-bana. Various ichnofossils that mainly comprised graphoglyptids are observable at the basal surfaces of turbidite beds. **D)** Schematic columnar section of the Izaki Olistolith. Stratigraphic intervals and interpretation of depositional environment are also shown. st: siltstone, vfs: very fine

678 sandstone, fs: fine sandstone, ms: medium sandstone.

679 FIG. 8.—Ichnofossils occurring on the bottom surface of sandstone beds in the Izaki
680 Olistolith. **A)** *Cosmorhapse parva*. **B)** *Gordia marina*. **C)** *Helminthorhapse japonica*.
681 **D)** *Paleodictyon minimum* (*Pm*) and *P. strozzii* (*Ps*). **E)** *Punctorhapse parallela*. **F)**
682 *Spirophycus bicornis*. Scale bar = 2 cm.

683 FIG. 9.—Relationships between resampling lengths and number of ichnogenera obtained by
684 the MIDIRT method applied to field data from the Izaki Olistolith. Each plot
685 represents the average number of detected ichnogenera of 100 trials for each
686 resampling length. The fitted curves and coefficients of determination are also shown.
687 Error bars indicate the 95% confidence intervals of the average numbers of observed
688 ichnogenera. **A)** Channel-fill deposits. **B)** Levee deposits.

689 FIG. 10.—Variation of expected number of ichnotaxa E_s of the channel-fill deposits and
690 the levee deposits in the Izaki Olistolith against the ichnotaxa coverage parameter.

691 FIG. 11.—Comparison of two fitting functions with the artificial data of the uniform
692 distribution with 10 ichnotaxa.

693 TABLE. 1.—Coefficients of the curves fitted to the resampled data of the artificial outcrop
694 images, parameters showing goodness of fit, and ratio of ichnodiversity. The rightmost
695 column represents ratios between ichnodiversity of each exposure condition and 100%
696 exposed data in the same distribution patterns and maximum numbers of ichnotaxa.

697 TABLE. 2.—Results comprising ratios of estimated numbers of ichnotaxa for artificial
698 outcrop images.

699 TABLE. 3.—List of ichnotaxa occurring on the bottom of sandstone beds in the Izaki

700 Olistolith.

701 TABLE. 4.—Coefficients of the curves fitted to the resampled data of the field data of the

702 Izaki Olistolith and parameters showing goodness of fit.

1. Acquisition of outcrop images

2. Painting with particular gray scale for each ichnotaxon

3a. Selecting a image at random

3c. Repeating 100 times

3b. Counting kinds of gray scale value on the line-of-interest randomly set

4. Curve fitting to resampled data

5. Estimation of ichnodiversity

Fig. 1

Ichnofossil images	Chances of occurrence	
	High diversity	Low diversity
<i>A</i>	0.30	0.64
<i>B</i>	0.20	0.23
<i>C</i>	0.20	0.09
<i>D</i>	0.10	0.03
<i>E</i>	0.06	0.01
<i>F</i>	0.05	-
<i>G</i>	0.03	-
<i>H</i>	0.03	-
<i>I</i>	0.02	-
<i>J</i>	0.01	-

Fig. 2

Fig. 3

Fig. 4

Fig. 5

Fig. 6

Fig. 7

Fig. 8

Fig. 9

Fig. 10

Fig. 11

Distribution patterns	Exposure	α_1			α_2			R_{se}	Ratio
		Estimate	Std. Error	p	Estimate	Std. Error	p		
uniform w/ 10 ichnotaxa	100%	1.726	0.041	1.440E-42	0.036	0.004	4.960E-11	0.236	-
	90%	1.847	0.049	3.700E-22	0.026	0.003	7.010E-08	0.178	1.070
	80%	1.966	0.065	5.890E-20	0.022	0.003	9.820E-07	0.230	1.139
	70%	1.837	0.073	2.730E-18	0.022	0.004	8.690E-06	0.252	1.065
	60%	2.051	0.068	5.490E-20	0.015	0.002	1.010E-07	0.216	1.188
	50%	1.812	0.061	6.910E-20	0.018	0.002	1.330E-07	0.194	1.050
	40%	2.073	0.029	2.060E-28	0.011	0.001	1.480E-16	0.078	1.201
	30%	2.628	0.107	4.990E-18	0.006	0.001	3.820E-09	0.203	1.523
	20%	1.667	0.044	2.700E-22	0.017	0.001	2.340E-11	0.109	0.966
	10%	1.811	0.069	1.190E-18	0.013	0.001	2.280E-10	0.112	1.049
uniform w/ 5 ichnotaxa	100%	0.835	0.035	2.630E-30	0.047	0.011	4.250E-05	0.205	-
	90%	0.836	0.033	1.960E-18	0.052	0.011	9.990E-05	0.124	1.001
	80%	0.853	0.030	2.010E-19	0.051	0.009	2.060E-05	0.114	1.022
	70%	0.857	0.034	2.140E-18	0.058	0.012	7.690E-05	0.127	1.027
	60%	0.878	0.037	9.380E-18	0.059	0.013	1.290E-04	0.138	1.052
	50%	0.907	0.041	4.950E-17	0.060	0.014	2.085E-04	0.151	1.086
	40%	0.954	0.048	6.080E-16	0.058	0.014	4.191E-04	0.174	1.143
	30%	1.144	0.062	3.060E-15	0.031	0.007	1.306E-04	0.200	1.371
	20%	1.374	0.037	3.770E-22	0.017	0.001	2.720E-11	0.090	1.646
	10%	2.064	0.047	1.300E-23	0.010	0.001	1.450E-15	0.066	2.472
patchy w/ 10 ichnotaxa	100%	1.797	0.046	1.380E-40	0.029	0.004	2.030E-10	0.263	-
	90%	1.906	0.065	1.140E-19	0.023	0.004	2.530E-06	0.234	1.061
	80%	1.949	0.072	5.990E-19	0.023	0.004	5.380E-06	0.254	1.085
	70%	2.056	0.085	6.910E-18	0.020	0.004	1.200E-05	0.290	1.144
	60%	2.307	0.072	1.500E-20	0.013	0.002	2.220E-08	0.223	1.284
	50%	2.310	0.076	4.790E-20	0.016	0.002	5.390E-08	0.235	1.286
	40%	2.510	0.069	6.690E-22	0.014	0.001	4.060E-10	0.195	1.397
	30%	2.735	0.050	7.580E-26	0.011	0.001	1.200E-14	0.124	1.523
	20%	2.783	0.083	5.660E-21	0.012	0.001	4.970E-11	0.181	1.549
	10%	6.682	0.255	1.300E-18	0.002	0.000	1.830E-14	0.101	3.719
patchy w/ 5 ichnotaxa	100%	0.900	0.030	2.890E-35	0.031	0.005	1.360E-07	0.169	-
	90%	0.989	0.043	1.900E-17	0.021	0.004	4.340E-05	0.151	1.098
	80%	1.046	0.048	8.100E-17	0.017	0.003	4.730E-05	0.164	1.162
	70%	1.124	0.018	6.260E-27	0.012	0.001	1.020E-13	0.058	1.249
	60%	1.401	0.034	4.710E-23	0.005	0.000	2.580E-12	0.081	1.556
	50%	0.983	0.050	8.270E-16	0.011	0.002	1.570E-05	0.145	1.092
	40%	0.934	0.034	3.650E-19	0.018	0.002	1.690E-07	0.102	1.038
	30%	1.159	0.065	5.900E-15	0.010	0.002	7.610E-06	0.156	1.288
	20%	0.770	0.074	3.490E-10	0.032	0.011	9.245E-03	0.218	0.855
	10%	0.918	0.098	2.450E-09	0.032	0.011	6.280E-03	0.237	1.019

Table 1

	uniform w/ 10 ichnotaxa	uniform w/ 5 ichnotaxa	patchy w/ 10 ichnotaxa	patchy w/ 5 ichnotaxa
uniform w/ 10 ichnotaxa	-	0.484	1.041	0.522
uniform w/ 5 ichnotaxa	2.067	-	2.152	1.078
patchy w/ 10 ichnotaxa	0.961	0.465	-	0.501
patchy w/ 5 ichnotaxa	1.917	0.927	1.996	-

Table 2

Sedimentary environments	Ichnotaxa
Levee (11 ichnogenera 12 ichnospecies)	<i>Bergaueria</i> isp. <i>Gordia marina</i> <i>Gordia</i> isp. <i>Helminthopsis abeli</i> <i>Megagraption irregulare</i> <i>Nereites missouriensis</i> <i>Paleodictyon strozzii</i> <i>Paleophycus</i> isp. <i>Phycosiphon incertum</i> <i>Spirorhaphe involuta</i> Circular trace Radial trace
Channel-fill (22 ichnogenera 24 ichnospecies)	<i>Asteriacites lumbricalis</i> <i>Belorhaphe zickzack</i> <i>Bergaueria</i> isp. <i>Cosmorhaphe parva</i> <i>Desmograption ichtyforme</i> <i>Gordia marina</i> <i>Gordia</i> isp. <i>Halopoa imbriata</i> <i>Helminthopsis abeli</i> <i>Helminthorhaphe japonica</i> <i>Lorenzina</i> isp. <i>Mammilichnis aggeris</i> <i>Megagraption irregulare</i> <i>Nereites missouriensis</i> <i>Paleodictyon minimum</i> <i>Paleodictyon strozzii</i> <i>Paleophycus</i> isp. <i>Phycosiphon incertum</i> <i>Punctorhaphe parallela</i> <i>Spirophycus bicornis</i> <i>Spirorhaphe involuta</i> Bloom like trace Circular trace Radial trace

Table 3

	a_1			a_2			R_{se}
	Estimate	Std. Error	p	Estimate	Std. Error	p	
Channel-fill	4.916	0.035	6.994E-85	2.650.E-03	7.171.E-05	2.733E-46	0.180
Levee	2.825	0.077	2.050E-42	6.722.E-03	5.841.E-04	1.030E-16	0.300

Table 4