

1 Influence of urban spatial configuration and sea breeze on land surface temperature on
2 summer clear-sky days

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14

15 **Abstract**

16 Understanding the spatiotemporal characteristics of land surface temperature (LST), a physical quantity that
17 greatly contributes to the formation of urban thermal environment, is essential for constructing a sustainable urban
18 environment. This study explores the relationship between urban space, atmospheric conditions, and diurnal change
19 characteristics of LST on summer clear-sky days in an urban area of Osaka Plain using a 10 min LST dataset retrieved
20 from the Japanese geostationary satellite (Himawari-8) data. The principal temporal change patterns forming the
21 diurnal cycle of LST were extracted through a principal component analysis, and their spatial variabilities were
22 investigated through the exploitation of ground-based meteorological observation and land-use data. The result
23 showed two temporal change patterns, identified as the amplitude and phase of the diurnal change, being extracted
24 in different clear-sky days. A hierarchical cluster analysis confirmed that the spatial variations of such patterns
25 corresponded to the sea breeze and urban land-use. Moreover, LST change at daytime was found to be larger in the
26 high-density areas rather than in the low-density areas of low buildings, thereby suggesting the influence of building
27 density on LST at daytime.

28

29 Keywords: Himawari-8; urban heat island; land surface temperature; urban land use; sea breeze

30

31 **1. Introduction**

32 Urban heat island which is characterized by deterioration of a hot environment on an urban scale, has recently
33 become an increasingly serious phenomenon, in addition to the rise in temperature due to global warming (e.g.,
34 Kusaka et al., 2012). The foreseeable impact of urban heat islands is exemplified by the case of the Japanese
35 metropolitan areas, where approximately 50,000 heat-stroke patients are being sent to hospitals via ambulance yearly,
36 with a particularly high number of 90,000 during the summer of 2018 (Japan Fire and Disaster Management Agency,
37 2018). In addition to this direct effect, the drop in pressure accompanying the temperature rise may cause enhanced
38 precipitation (Seino et al., 2018) and retention of air pollutants, such as NO_x and SO_x, in the environment (Saitoh et
39 al., 1996).

40 Thermal indices have been proposed for evaluating the thermal environment in an urban area, with such examples
41 as wet-bulb globe temperature and universal thermal climate index, that consider various meteorological elements
42 i.e., surface air temperature (SAT), radiant heat, humidity, and wind speed (Budd, 2008; Jendritzky et al., 2012). SAT
43 and land surface temperature (LST) are likewise effective in evaluating the intensity of urban heat island and in
44 obtaining a detailed spatiotemporal characteristic of the thermal environment (Memon et al., 2009). SAT and LST
45 exhibit different spatiotemporal dynamics because the involved processes are dissimilar. SAT is generally observed
46 in ground weather stations. In urban areas with available dense observation networks, the spatial variability
47 characteristics of diurnal and seasonal changes of SAT have been investigated through a multivariate analysis (Kim
48 and Beik, 2005; Suzuki and Iizuka, 2009; Oku and Masumoto, 2014).

49 On the other hand, LST is retrieved by polar-orbiting satellite/aircraft sensors owing to their high spatial resolution
50 (Memon et al., 2009; Weng, 2009). The obtained LST data have been used to investigate their relationships with
51 urban vegetation and land use. Vegetation-wise, the focus is on changes in the LST from the increase/decrease of
52 green space (Li et al., 2009; Tsunematsu et al., 2016; Honjo et al., 2017), along with the vegetation phenology (Li et
53 al., 2011; Xu et al., 2013). Vegetation status is represented by vegetation coverage based on land use and land cover
54 (LULC) or remotely sensed vegetation indices, such as the normalized difference vegetation index (NDVI) and
55 fractional vegetation cover. Numerous studies have confirmed the negative relationship between LST and the
56 vegetation indices (Weng, 2009).

57 Land-use-wise, the focus is on correlations between LST and landscape metrics quantifying the horizontal urban
58 structures (Li et al., 2011; Connors et al., 2013). The metrics measure density, shape complexity, diversity, and
59 adjacency, among others, in LULC types. Additionally, relationships between LST and vertical urban structures were
60 also assessed using land-use data which have building height information (Li et al., 2011; Tsunematsu et al., 2016).
61 These studies revealed that both the cover composition and spatial configuration of the urban land surface affect the
62 LST. Yamamoto and Ishikawa (2018a) reported that the LST in coastal cities may be affected by the sea/land breeze
63 (i.e., a local meteorological field with diurnal cycle); thus, they stressed the inclusion of local meteorological effects
64 when studying the influence of urban spatial configuration on LST. In the evaluation of thermal environment using
65 instantaneous LST values obtained by polar-orbiting satellite/aircraft observations, the influences of the local
66 meteorological field on LST may pose a challenge.

67 This study presents LST observations in an urban area, through Himawari-8, a Japanese third-generation
68 geostationary satellite that can retrieve LST with high spatial (approximately 2 km) and temporal (10 min) resolutions.
69 Yamamoto and Ishikawa (2018a) attempted to apply the Himawari-8 LST product to an urban area and confirmed
70 that its spatial and temporal resolutions are effective in obtaining the thermal characteristics of land use and local
71 meteorological field in that area. Such advantage, especially temporal resolution, was utilized by this research to
72 understand the influence of urban spatial configuration and sea breeze on the temporal changes of LST during the
73 summer clear-sky days in an urban area within Osaka Plain, Japan.

74

75 **2. Data and method**

76 **2.1 Analyzed area and data**

77 Fig. 1 shows the urban study area of Osaka Prefecture, which is known to have the respective lowest and highest
78 ratios of forest and habitable areas (the latter is calculated by subtracting the area of forests, grasslands, and bodies
79 of water from the total area of the prefecture) in Japan prefectures (Japan Statistics Bureau, 2019). Additionally, the
80 prefecture has the largest number of heat-stroke patients sent to hospitals via ambulance in recent years (Japan Fire
81 and Disaster Management Agency, 2018). From the land use map in Fig. 1, there is a bay (Osaka Bay) on the west
82 side of the analyzed area. The central urban area is occupied by high buildings, and the surrounding areas are occupied
83 by low buildings (high or low density).

84 LST was retrieved by a nonlinear three-band algorithm (Yamamoto et al., 2018). The algorithm requires brightness

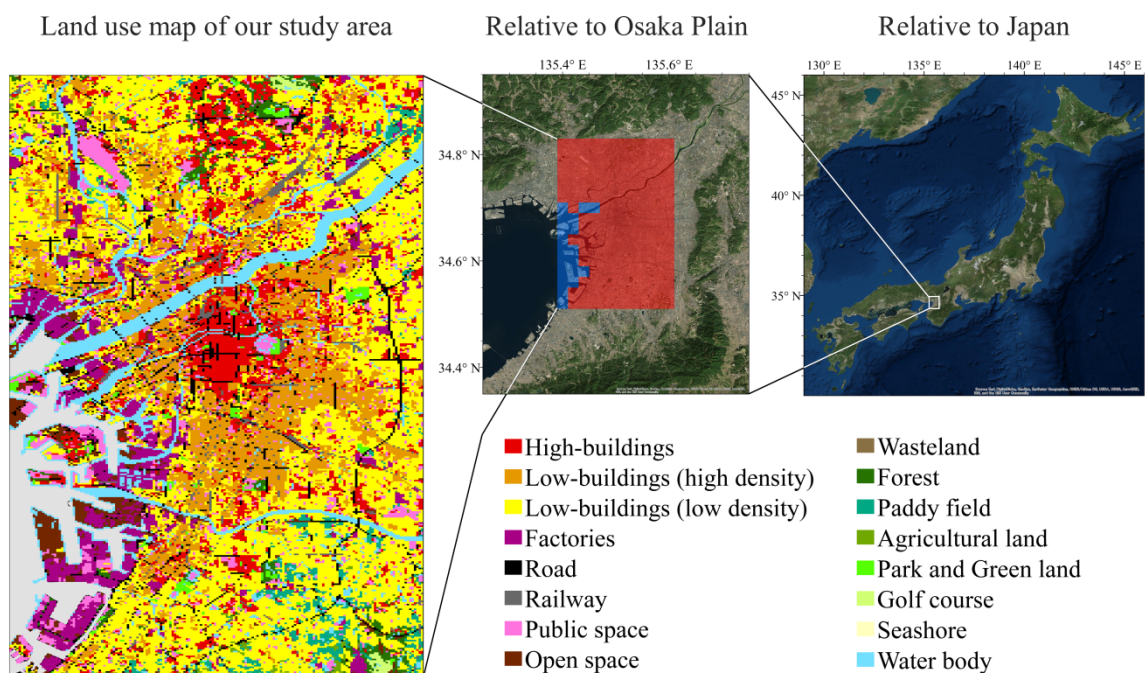


Fig. 1. Location map and land use of the study area (34.51°N–34.83°N, 135.39°E–135.61°E). In the middle figure, LST pixels used in the analysis are shown in red, whereas pixels with a percentage of water surface above 50% are shown in blue.

85 temperatures and land surface emissivities of bands 13, 14, and 15 (centered at 10.4, 11.2 and 12.4 μm , respectively)
 86 of Himawari-8 in the thermal infrared region. The brightness temperature data of Himawari-8 were provided by the
 87 Center for Environmental Remote Sensing, Chiba University, whereas the land surface emissivities were estimated
 88 by a semi-empirical method based on the land cover information and NDVI (Yamamoto and Ishikawa, 2018b). LST
 89 retrieval was limited to a clear-sky condition where the upward thermal infrared radiance from the land surface can
 90 be observed by the Himawari-8 sensor. The spatial and temporal resolutions of the retrieved LST data were 0.02°
 91 and 10 min, respectively. A total of 157 terrestrial pixels which have water surface percentage of less than 50% were
 92 used in all pixels of the study area (Fig. 1). The satellite zenith angle in the study area ranges from 40.47°– 40.86°.
 93 Therefore, all pixels used in the analysis were observed at approximately the same angle. The days when the diurnal
 94 change of LST of the 157 pixels was retrieved (i.e., days with zero cloud cover) were chosen for focal analysis on
 95 the temporal change characteristics of LST. The target season was summer. The presence of clouds was judged by a
 96 cloud detection method employing Himawari-8's visible and infrared data (Yamamoto et al., 2018). The selection
 97 period was the summer season, from July 2015 of the onset of the Himawari-8 operation to August 2018. The final
 98 selected dates were August 12, 2016, July 19, 2018, and August 4, 2018.

99 The urban area land-use segmentation mesh data of national land numerical information provided by the Ministry
 100 of Land, Infrastructure and Transport of Japan (<http://nlftp.mlit.go.jp/ksj/gml/datalist/KsjTmplt-L03-b-u.html>; see
 101 Fig. 1) were used for investigating the relationship between LST and urban land use. The spatial resolution was 100
 102 m; i.e., it was resolved with approximately 400 grid of land-use information per pixel of the LST data. Furthermore,
 103 the SAT, wind direction, and wind speed data from the Automated Meteorological Data Acquisition System
 104 (AMeDAS) operated by the Japan Meteorological Agency, were exploited for investigating the relationship between
 105 LST and meteorological field. The observation points included Osaka, Toyonaka, Sakai, and Yao (see Figs. S1–S3)
 106 at 10 min temporal resolution.

107 2.2 Method of analysis

108 The influence of urban spatial configuration and sea breeze on the temporal changes of LST, not the instantaneous
 109 LST, was the highlight of this study. Such temporal changes were defined through a summary of the original diurnal
 110 changes of LST using the principal component analysis (PCA). In other words, they were the principal temporal
 111 change characteristics forming the diurnal changes of LST.

112 The temporal resolution of the LST data was 10 min, leading to a total of 144 LSTs retrieved for each pixel
 113 throughout the day. As the PCA was applied, the diurnal changes of LST were converted into the diurnal changes of
 114 the deviation by subtracting the daily mean LST of each pixel. The diurnal changes of the deviation of all pixels were
 115 expressed as the following 157×144 matrix \mathbf{dT} :

$$\mathbf{dT} = \begin{pmatrix} dT_{1,1} & \cdots & dT_{1,144} \\ \vdots & & \vdots \\ dT_{157,1} & \cdots & dT_{157,144} \end{pmatrix}, \quad (1)$$

116 and the 144×144 variance-covariance matrix \mathbf{V} of \mathbf{dT} was calculated as follows:

$$\mathbf{V} = \begin{pmatrix} s_{1,1} & \cdots & s_{1,144} \\ \vdots & & \vdots \\ s_{144,1} & \cdots & s_{144,144} \end{pmatrix}, \quad (2)$$

117 where the variances and covariances contained in \mathbf{V} were given by:

$$s_{ij} = s_{ji} = \frac{1}{157} \sum_{k=1}^{157} (dT_{ik} - \overline{dT}_i)(dT_{jk} - \overline{dT}_j), \quad (3)$$

118 where \overline{dT}_i (\overline{dT}_j) is a column average of elements at time i (j) in \mathbf{dT} . By solving the eigenvalue problem for \mathbf{V} ,
 119 the eigenvalues ($\lambda_1, \dots, \lambda_k, \dots, \lambda_{144}$) and the corresponding eigenvectors $\mathbf{a}_k = (a_{k1}, \dots, a_{k144})$ were obtained.

Table 1. Eigenvalues and contribution ratios from PC1 to PC5 derived by the PCA on the three analysis days.

Principal Component	August 12, 2016		July 19, 2018		August 4, 2018	
	Eigenvalue	Contribution Ratio (%)	Eigenvalue	Contribution Ratio (%)	Eigenvalue	Contribution Ratio (%)
PC1	108.15	85.26	142.08	78.08	129.85	84.95
PC2	8.76	6.91	21.84	12.00	10.60	6.93
PC3	4.97	3.92	6.17	3.39	4.95	3.24
PC4	1.77	1.39	4.73	2.60	2.40	1.57
PC5	0.84	0.66	2.34	1.29	1.30	0.85

120 Therefore, the number of principal components (PCs) is 144. The 1×157 PC score of each PC was calculated as the
 121 scalar product of $d\mathbf{T}$ and \mathbf{a}_k . 157 PC scores were normalized for each PC. In this analysis, the PC score represents
 122 the reflection degree of the temporal change characteristic represented by each PC in each pixel. The eigenvector
 123 represents the weight corresponding to the LST deviation at each time in the calculation of the PC score.
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 125

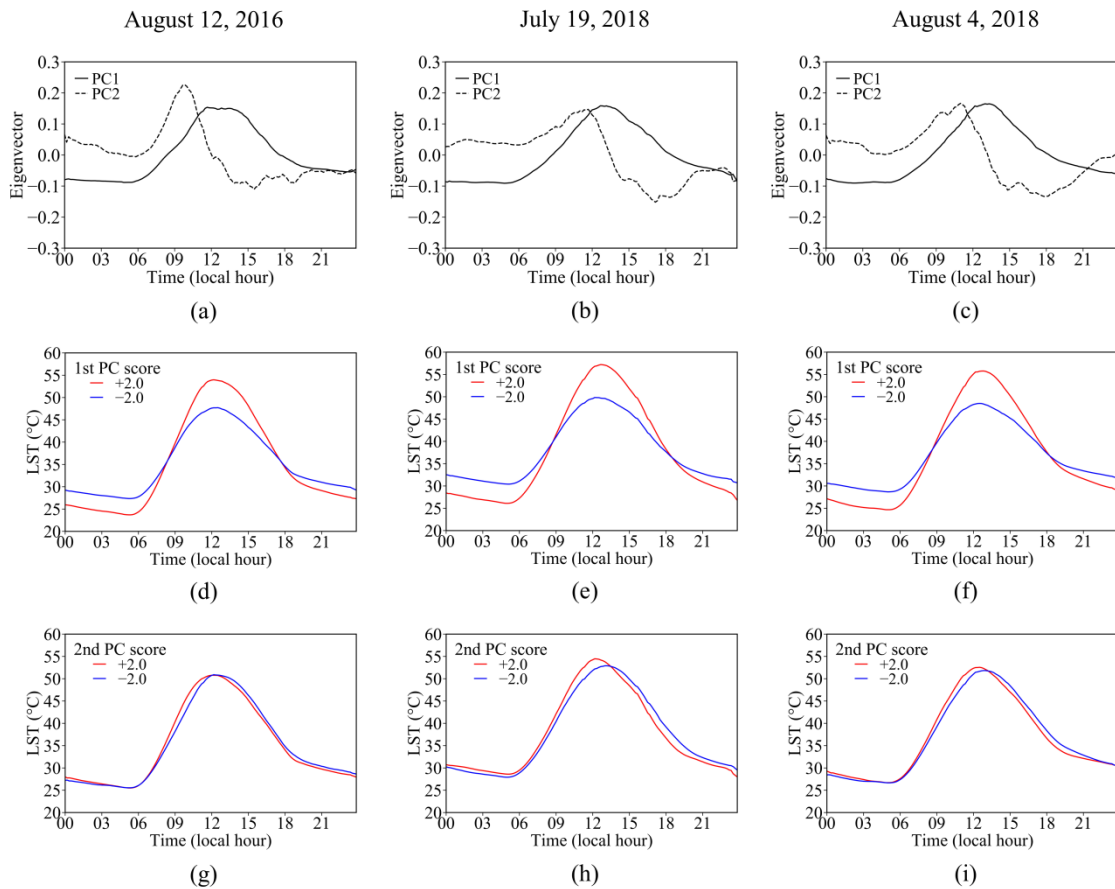


Fig. 2. (a)–(c) PC1 and PC2 eigenvectors for the three analysis days. (d)–(f) Effect of PC1 on diurnal change of the average LST of the study area. (g)–(i) Effect of PC2 on diurnal change of the average LST of the study area.

126 **3. PCA results**

127 Table 1 shows the eigenvalues and contribution ratios from the first PC (PC1) to the fifth PC (PC5) derived by the
128 PCA. The contribution ratios of PC1 were dominant, and the cumulative contribution ratios of PC1 and PC2 were
129 more than 90% on all 3 days. In other words, the diurnal cycle of LST in a summer clear-sky day could be adequately
130 described by the temporal change characteristics represented by PC1 and PC2. This led to the investigation of the
131 influence of the urban spatial configuration and meteorological field on the temporal change characteristics of LST
132 represented by PC1 and PC2.

133 Figs. 2(a) to 2(c) show the eigenvectors of PC1 and PC2. The former had negative and positive components at
134 nighttime and daytime, respectively, on all 3 days, with the positive maximum value reached around noon. Note that
135 the temporal change characteristic represented by PC1 was emphasized around noon. Figs. 2(d) to 2(f) show the
136 responses of diurnal change of the average LST of 157 pixels in the study area when only the normalized PC score
137 of PC1 was changed by ± 2.0 . For the average LST deviation of the study area, the PC scores from PC1 to PC144
138 were 0. Thus, by changing only the PC1 score, the temporal change characteristics represented by PC1 can be
139 emphasized. From Figs. 2(d) to 2(f), the amplitudes of the diurnal cycle were observed to vary with the variation of
140 the PC score on all 3 days, implying that PC1 represents the amplitude of the diurnal cycle on the clear-sky days.

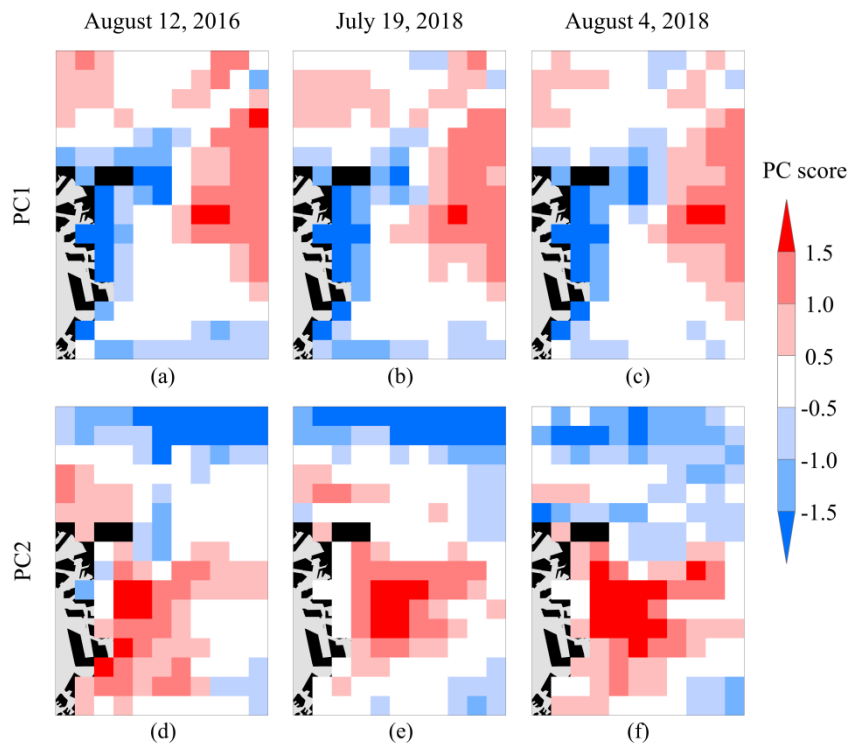


Fig. 3. Spatial distributions of the normalized PC scores of PC1 (a–c) and PC2 (d–f) derived by the PCA on the three analysis days.

141 As shown in Figs. 2(a) to 2(c), the components of PC2 eigenvectors tended to be positive in the forenoon (i.e., the
142 time between midnight and around noon) and negative in the afternoon (i.e., the time between around noon and
143 midnight). The vector components took a positive maximum value from 9:00 to 12:00 when the temperature rise
144 became large, whereas they took a negative maximum value from 15:00 to 18:00 when the temperature drop was
145 large. Moreover, the components were close to 0 around noon when the vector components of PC1 took an extreme
146 value. Figs. 2(g) to 2(i) display a visualization of the temporal change characteristic represented by PC2, similar to
147 Figs. 2(d) to 2(f). These figures imply that PC2 represents the phase shifting by the waveform leaning
148 forward/backward and independent of the amplitude.

149 As described above, a large (small) PC1 score indicates that the amplitude of the diurnal cycle is relatively large
150 (small). A large (small) PC2 score indicates that the phase of the diurnal cycle is relatively advanced (delayed). Fig.
151 3 shows the spatial distributions of the scores of PC1 and PC2, with the former at a positive tendency in the inland
152 area and at a negative tendency in the coastal area. On the other hand, the spatial variability of the PC2 score tended
153 to be geographically opposite to that of PC1. Nevertheless, both score distributions were not diametrically opposite
154 and seemingly corresponding to land-use distribution. For example, the scores of both PCs were low in a pixel
155 occupied by high buildings (see Fig. 1), regardless of the area being coastal or inland.

156

157 **4. Discussion**

158 Fig. 4 shows the scatter diagrams of PC1 and PC2 scores for the three analysis days, with the addition of dominant
159 land-use type and coastal distance information. Moreover, in order to objectively classify the characteristics of the
160 distributions in Fig. 4, we performed cluster analysis (Ward's method) using the 1st and 2nd PC scores, coastal
161 distance (pixels) and proportions of high-buildings, low-buildings (high density), low-buildings (low density),
162 factories, vegetation, and water body of each pixel as input variables. These variables were standardized as pre-
163 processing. The clusters were classified with a threshold Euclidean distance of 10. Figs. S1 – S3 are the dendrograms
164 obtained by the cluster analysis, and Table 2 shows the characteristics of the obtained clusters for each analysis day.
165 The clusters consistently obtained on all analysis days were characterized as follows: pixels with a high proportion
166 of high-buildings (see (a) cluster1, (b) cluster4, and (c) cluster5 in Table 2); pixels with a high proportion of low-
167 buildings (high density) (see (a) cluster7, (b) cluster7, and (c) cluster2 in Table 2); coastal pixels with a high
168 proportion of low-buildings (low density) (see (a) cluster8, (b) cluster6, and (c) cluster4 in Table 2); inland pixels

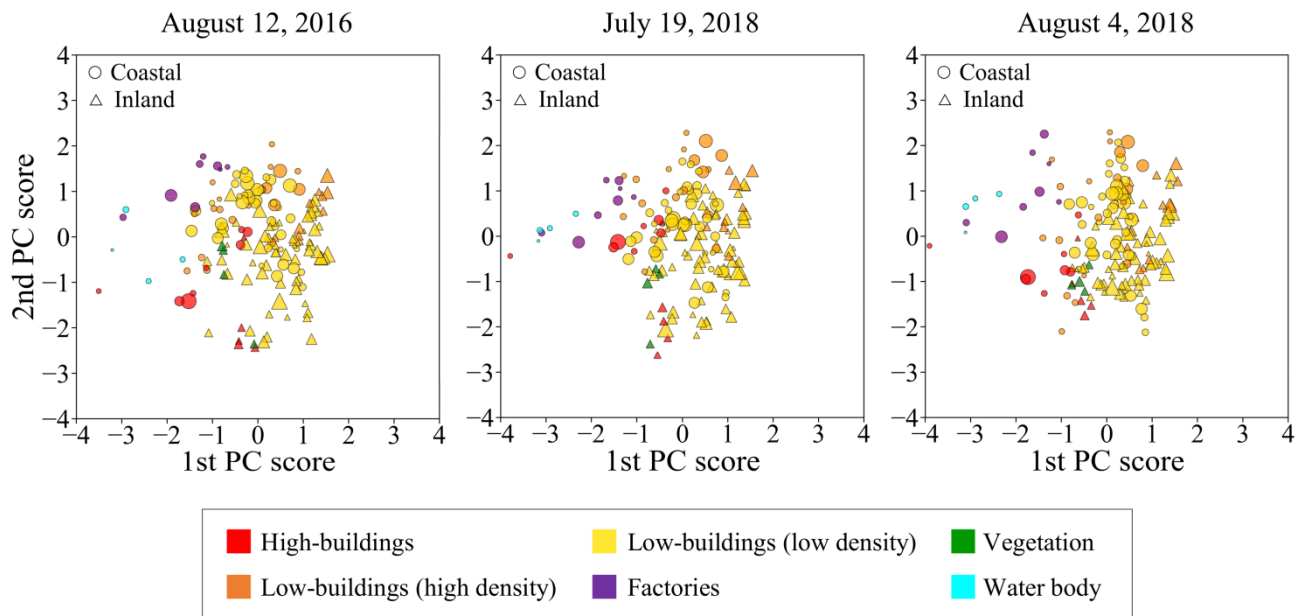


Fig. 4. Scatter diagrams between the PC1 and PC2 scores of the pixels for the three dates in the study area. Of all pixels for each date, half of the pixels close to the coast are indicated by circles as the coastal pixels, and half of the pixels far from the coast are indicated by triangles as the inland pixels. Each pixel is color coded according to land-use type with the highest occupancy, and the marker size corresponds to the occupancy rate.

169 with a high proportion of low-buildings (low density) (see (a) cluster4, (b) cluster2, and (c) cluster7 in Table 2);
 170 coastal pixels with a high proportion of factories (see (a) cluster5, (b) cluster5, and (c) cluster1 in Table 2); inland
 171 pixels with a high proportion of vegetation (see (a) cluster3, (b) cluster3, and (c) cluster6 in Table 2); pixels with a
 172 high proportion of water body (see (a) cluster6, (b) cluster1, and (c) cluster3 in Table 2). Only on August 12, 2016,
 173 pixels with a relatively close Euclidean distance from cluster1 but are located inland were classified as cluster2. This
 174 section discusses the consistent or inconsistent characteristics of PC1 and PC2 in three analysis days in terms of the
 175 land-sea breeze circulation and urban land use.

176 4.1 Relationship between LST and sea breeze

177 As shown in Fig. 4, coastal pixels (indicated by circles) tended to have a low PC1 score and a high PC2 score.
 178 Conversely, inland pixels (indicated by triangles) tended to have a high PC1 score and a low PC2 score. Such
 179 tendencies were aligned with the observation of Yamamoto and Ishikawa (2018a) who considered these to be the
 180 results of land and sea breeze circulation, although concrete evidence of the relationship is yet to be found. The
 181 variability of the land and sea breeze circulation was accounted herein through a more robust analysis of three
 182 different meteorological cases.

Table 2. Clusters resulting from cluster analysis and their mean characteristics of the input variables on (a) August 12, 2016, (b) July 19, 2018, and (c) August 4, 2018.

(a) August 12, 2016

Cluster	1st PC score	2nd PC score	Coastal distance [pixels]	High-buildings [%]	Low-buildings (high density) [%]	Low-buildings (low density) [%]	Factories [%]	Vegetation [%]	Water body [%]
1	-1.25	-0.67	4.67	43.71	13.95	7.55	0.64	2.48	5.17
2	0.31	-1.98	8.74	27.53	2.42	42.80	1.44	10.14	2.03
3	-0.72	-0.56	8.58	3.09	0.24	46.07	3.09	33.68	4.78
4	0.71	-0.23	8.62	8.61	4.80	63.33	2.75	4.80	4.34
5	-1.17	1.36	2.29	2.25	2.87	10.18	40.19	1.47	16.31
6	-0.98	-0.24	5.26	11.28	10.32	20.93	9.41	1.32	27.05
7	0.78	0.64	6.63	11.74	47.51	22.68	0.66	1.97	1.73
8	-0.11	0.52	4.09	9.52	15.10	42.33	6.92	3.99	6.69

(b) July 19, 2018

Cluster	1st PC score	2nd PC score	Coastal distance [pixels]	High-buildings [%]	Low-buildings (high density) [%]	Low-buildings (low density) [%]	Factories [%]	Vegetation [%]	Water body [%]
1	0.11	-0.36	7.72	15.93	10.40	30.72	3.01	0.93	26.18
2	0.44	-0.40	7.50	11.49	5.36	57.66	2.46	5.89	4.38
3	-0.60	-0.94	8.74	3.01	0.26	43.82	3.34	35.31	4.88
4	-0.71	-1.03	6.65	50.54	3.46	13.53	0.02	10.11	3.33
5	-2.27	0.45	2.08	5.75	2.52	11.25	29.50	1.96	19.42
6	-0.10	0.73	3.34	6.84	23.63	34.22	9.78	1.92	7.64
7	0.65	0.87	6.65	14.48	46.78	19.46	0.63	1.86	1.92

(c) August 4, 2018

Cluster	1st PC score	2nd PC score	Coastal distance [pixels]	High-buildings [%]	Low-buildings (high density) [%]	Low-buildings (low density) [%]	Factories [%]	Vegetation [%]	Water body [%]
1	-2.27	0.82	2.08	5.75	2.52	11.25	29.50	1.96	19.42
2	0.65	0.60	6.58	13.63	48.23	19.51	0.58	1.90	1.77
3	-0.20	-0.93	5.49	11.17	18.56	27.99	7.17	1.10	19.42
4	0.05	0.85	4.28	8.89	13.47	47.55	5.68	4.61	5.65
5	-0.93	-1.14	6.65	50.54	3.46	13.53	0.02	10.11	3.33
6	-0.54	-0.73	8.91	2.96	0.05	37.10	4.68	40.43	6.00
7	0.56	-0.49	8.33	11.47	5.14	57.79	2.60	6.45	3.52

183 Fig. 5 shows the penetration time and duration of the sea breeze at each ground observation point for the 3 days.
 184 Penetration time and end time of the sea breeze were judged on the basis of changes in wind direction/speed and
 185 SAT. Thus, the westerly wind ($270^{\circ}\pm 22.5^{\circ}$) at Sakai and Yao, and southwesterly wind ($225^{\circ}\pm 22.5^{\circ}$) at Osaka and
 186 Toyonaka were regarded as the sea breeze, respectively. The penetration time was defined as the time when the wind
 187 speed increased, its direction continued for more than 30 min, and the SAT decreased within 30 min. The end time
 188 was judged as the time when the wind speed decreased and the wind direction began to change in a direction other
 189 than that of the sea breeze. More detailed wind direction/speed and SAT time-series changes are shown in Figs.
 190 S4–S6.

191 Sea breeze causes reduction in both the SAT and the LST. Thus, the trend of the low PC1 score of coastal pixels
 192 was thought to be due to the suppression of LST rise by a cool sea breeze, which reduced PC1's amplitude. Moreover,
 193 the differences between the coastal and inland PC2 scores were thought as a result of the penetration time and duration
 194 of the sea breeze. The sea breeze characteristics for the August 12, 2016, case showed the least amount of regional

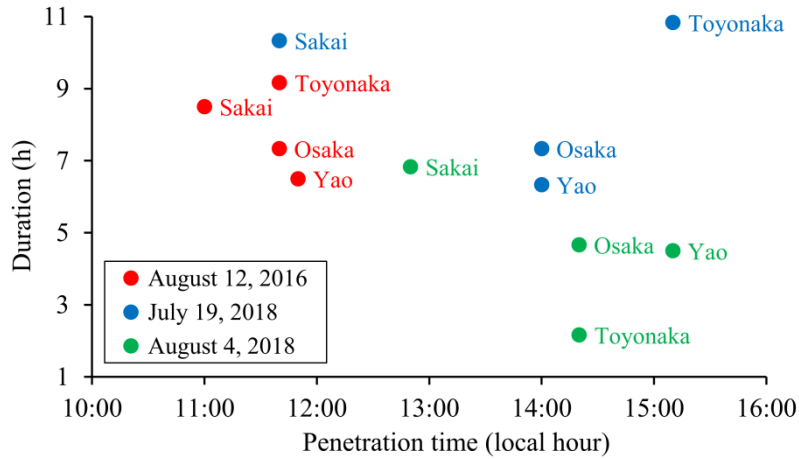


Fig. 5. Penetration time and duration of the sea breeze observed at the stations of AMeDAS in the study area on the three analysis dates.

195 variability compared with the other two cases (see Fig. 5). As shown in Fig. 2a, the components of PC2 eigenvector
 196 on August 12, 2016, did not have a clearly defined minimum value in the afternoon. In other words, the leaning of
 197 the wave on this day was affected by the heating rate in the morning rather than the cooling rate in the afternoon.
 198 This result can be explained by the fact that not much regional difference occurred in the penetration time and duration
 199 of the sea breeze. Thus, the cooling effect of the sea breeze without regional difference may have reduced the regional
 200 difference of LST cooling rate during the afternoon.

201 For the other 2 days, regional differences in the penetration time were large (Fig. 5), especially on July 19, 2018,
 202 whose duration was longer than on August 4, 2018. The components of PC2 eigenvectors in the afternoon on these
 203 2 days took smaller minimum values than those on August 12, 2016, which can be attributed from the regional
 204 difference in penetration time and duration of the sea breeze contrary to the case on August 12, 2016. In addition, the
 205 duration on July 19, 2018, was long, potentially causing large regional difference in the degree of the LST drop. As
 206 a result, PC2 contribution ratio may have become large (Table 1), and the PC2 score of the coastal pixels likewise
 207 became large, as compared with those of the other 2 days (Fig. 4).

208 4.2 Relationship between LST and urban land use

209 As shown in Fig. 4 and Table 2, the pixels occupied by high-buildings, factories, vegetation, or bodies of water
 210 tended to exhibit a low PC1 score, whereas that of pixels occupied by low-buildings (high/low density) tended to be
 211 high. Such tendency could be attributed to the thermal inertia of the surface material. Thermal inertia is a thermal
 212 and physical property indicating the degree of resistance to temperature change of a surface material (Sugawara and

213 Kondo, 1995; Haffner and Kidder, 1999; Zhan et al., 2012). As water surface has large thermal inertia, the amplitude
214 of the diurnal cycle at pixels occupied by the water bodies is small. In urban land-use areas, the amplitudes of the
215 diurnal cycle may reflect the thermal inertia of building roof and wall materials rather than the road materials since
216 the satellite zenith angle of Himawari-8 in the analysis area is about 40°. High-buildings and factories mainly consist
217 of concrete, glass and metal, which similarly show large thermal inertia. By contrast, low-buildings mainly consist
218 of wooden materials and slate roof that have small thermal inertia. Thus, differences in the building materials were
219 thought to have caused variation in the PC1 amplitude. A similar interpretation was made by previous studies using
220 polar-orbiting satellite/aircraft sensors, from the viewpoint of the thermal inertia or heat capacity of building materials
221 (e.g., Li et al., 2011; Tsunematsu et al., 2016). However, depending on the width among building clusters and street
222 orientation, more road surfaces are observed, and it is necessary to consider surface materials such as asphalt with
223 high thermal inertia. As shown in Table 2 (a) cluster5, (b) cluster5, and (c) cluster1, the pixels occupied by the
224 factories and water bodies may have been affected by the cooling effect of the sea breeze, being proximal to the coast.
225 Regarding the vegetated areas, transpiration was believed to have suppressed LST rise at daytime.

226 With regard to PC2, pixels occupied by the low-buildings (high density) tended to have high scores regardless of
227 the area as coastal and inland. The values of thermal inertia were thought to have varied from large to small, as the
228 land use changed from high-buildings, low-buildings (high density), and to low-buildings (low density) with the
229 decrease in concrete buildings and the increase in wooden buildings (Haffner and Kidder, 1999; Tsunematsu et al.,
230 2016). The larger the thermal inertia, the smaller was the amplitude (i.e., lower PC1 score) and the later the phase
231 (i.e., higher PC2 score). Nonetheless, only those pixels occupied by the low-buildings (high density) were biased to
232 the first quadrant in Fig. 4 (i.e., high PC1 and PC2 scores, shown in Table 2 (a) cluster7, (b) cluster7, and (c) cluster2).
233 This possibly be because more road surfaces with high thermal inertia have been observed in low-buildings (low
234 density). In addition, differences in surface roughness between high/low building densities may have different heat
235 balance processes in the diurnal cycle on the clear-sky days.

236 The shadows of buildings during the daytime suppress the temperature rise of the land surface and cause a decrease
237 in the amplitude of the diurnal cycle. Depending on the street orientations, the timing affected by the shadows may
238 vary, which may change the PC2 score. The shadow effects are considered to have been most significant in the high-
239 buildings, which have high building height and large variations of the building height. The shadow effects also vary
240 with the number of buildings and the width between buildings. It is challenging to discuss which is more affected by
241 shadows, the low-buildings (high density) with many buildings or the low-buildings (low density) with wide spaces

242 between buildings. In order to clarify the relationship between building arrangement and the amplitude/phase in the
243 diurnal cycle, the LST and land-use data with higher spatial resolutions (less than 100 m) are needed.

244 Fig. 6 shows the diurnal LST changes calculated using average PC1 and PC2 scores and daily mean LST of the
245 coastal pixels occupied by high-buildings, low-buildings (high density), and low-buildings (low density). LST at
246 nighttime was lower in the order of high-buildings, low-buildings (high density), and low-buildings (low density),
247 which could be due to the thermal inertia of the building materials as described above. However, the order during
248 daytime was not reversed of that during nighttime, and the low-buildings (high density) had the highest LST, implying
249 an increase in internal reflections of solar and infrared radiations, along with wind speed reduction between the
250 buildings as their density increased. More complex internal reflection of the radiations implies higher absorptivity
251 and thereby causes higher heating rate at daytime. Moreover, at a location where variation in building height was
252 small and that for building density was high, the wind above the group of buildings hardly blew into spaces between
253 the buildings (Hagishima et al., 2004). The convective heat transfer between the surface within street canyons (mainly
254 building walls and roads) and the atmosphere near it decreases under weak wind conditions. The LST retrieved from
255 Himawari-8 in the analysis area reflects the surface temperatures of more building roofs and walls in the street
256 canyons as the building density increases. Thus, these effects in radiation and wind fields may have resulted in a
257 larger rate of LST increase in the morning, and the waveform of the diurnal change leaned forward (i.e., the PC2
258 score increased) in the pixels occupied by low-buildings (high density).

259 LST differences between low-buildings (high density) and low-buildings (low density) shown in Fig. 6 may also
260 be due to differences in vegetation cover since vegetation reduces daytime and nighttime LST by transpiration.

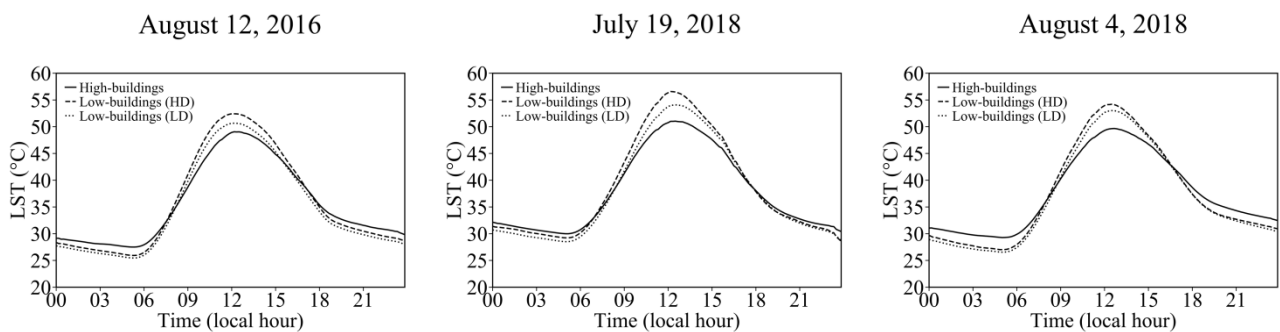


Fig. 6. Diurnal LST changes calculated with average PC1 and PC2 scores and daily mean LST of the coastal pixels where occupancy rate of high-buildings, low-buildings (high density), and low-buildings (low density) are higher than 70%. The solid, dashed, and dotted lines show the diurnal LST changes of high-buildings, low-buildings (high density), and low-buildings (low density), respectively.

261 However, the mean proportions of vegetation class of pixels used in Fig. 6 were 1.07% for low-buildings (high
262 density) and 3.28% for low-buildings (low density). The mean normalized difference vegetation index (NDVI) values
263 on three analysis days were 0.30, 0.29, and 0.35 for low-buildings (high density), 0.32, 0.29, and 0.34 for low-
264 buildings (low density). These small differences in vegetation cover are unlikely to have a significant effect on the
265 diurnal LST changes shown in Fig. 6.

266 The difference in the surface material may also be a cause of LST differences between low-buildings (high density)
267 and low-buildings (low density) shown in Fig. 6. **In Japan, as distance to the city center decreases the proportion of**
268 **concrete buildings increases whilst that of wood buildings decreases (Tsunematsu et al., 2016).** The trend suggests
269 that the effective thermal inertia is greater in the low-buildings (high density), but this is inconsistent with the larger
270 amplitude of the low-buildings (high density) shown in Fig. 6. Besides, if the building materials in low-buildings
271 (high density) were the same as those in low-buildings (low density), surface in low-buildings (high density) **is**
272 **covered by more roofs** and walls (e.g., slate and wooden materials) and fewer roads (e.g., asphalt and concrete)
273 compared with that in low-buildings (low density). In this case, the effective thermal inertia of the low-buildings
274 (high density) is smaller than that of the low-buildings (low density), so the temperature should drop significantly
275 during nighttime. This is also inconsistent with the nighttime LST shown in Fig. 6. Since the spatial resolution of
276 Himawari-8 is low, it is challenging to discuss the relationship between a temporal LST change and surface materials
277 in detail. However, it is difficult to interpret the results showing that the high-density areas of low-buildings were the
278 hottest during the daytime, based solely on the thermal inertia of the surface materials.

279

280 **5. Conclusions**

281 This study exploited the high spatiotemporal resolution of Himawari-8 sensor for investigation of the influence of
282 sea breeze and urban spatial configuration on the temporal changes of LST. The study area was an urban area of
283 Osaka Plain, Japan, **for target clear-sky days in midsummer.** Temporal changes of the LST were defined as the
284 principal temporal change characteristics forming the diurnal changes of LST, extracted via the PCA. The extracted
285 characteristics included the amplitude and phase (waveform leaning) of the diurnal cycle, whose spatial variabilities
286 were investigated using ground-based meteorological observation and land-use data of the study area.

287 As for the relationship with meteorological data, a reduction of the diurnal range and a forward-leaning of the
288 waveform corresponding to the penetration time and duration of the sea breeze were observed, which suggests the

289 importance of considering the effect of land and sea breeze circulation when observing LST in coastal cities. The
290 eigenvalues (contribution ratio) and scores of PC2 have a potential application in assessing the effects of sea breeze
291 on LST and urban heat island patterns. Note that on cloudy days or in the winter season, the diurnal changes of LST
292 are small, and the characteristics of the land-sea breeze circulation are different from those on summer clear-sky
293 days. On such days, different temporal change characteristics **from those in this analysis are expected.**

294 As for the relationship with land-use data, the variations of PC1 and PC2 scores were confirmed to correspond to
295 the thermal inertia of the building materials, which is consistent with previous studies. Furthermore, pixels occupied
296 by the low-buildings (high density) yielded high scores for both PC1 and PC2, regardless of the area as coastal and
297 inland. At daytime, low-buildings (high density) tended to exhibit higher LST than high-buildings and low-buildings
298 (low density), which suggests the impact of building density on diurnal LST change. The findings of this study
299 provide new facts on the relationship between LST and urban spatial configuration, as has been suggested in previous
300 studies (Li et al., 2011; Connors et al., 2013). Furthermore, such findings contribute to a deeper understanding of the
301 hot environment. **The combined knowledge of a spatial LST pattern derived from an aircraft or polar-orbiting satellite
302 and the diurnal LST patterns, including the effects of building density and sea breeze, derived from a new-generation
303 geostationary satellite, would help urban planners mitigate its deterioration due to urbanization.** For example,
304 applying the presented analysis method to the rapidly urbanizing Asia-Megacities would contribute to improve the
305 thermal environment measures of each city.

306

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310

311 **Declarations of interest:** none.

312

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Supplementary materials

Influence of urban spatial configuration and sea breeze on land surface temperature on summer clear-sky days

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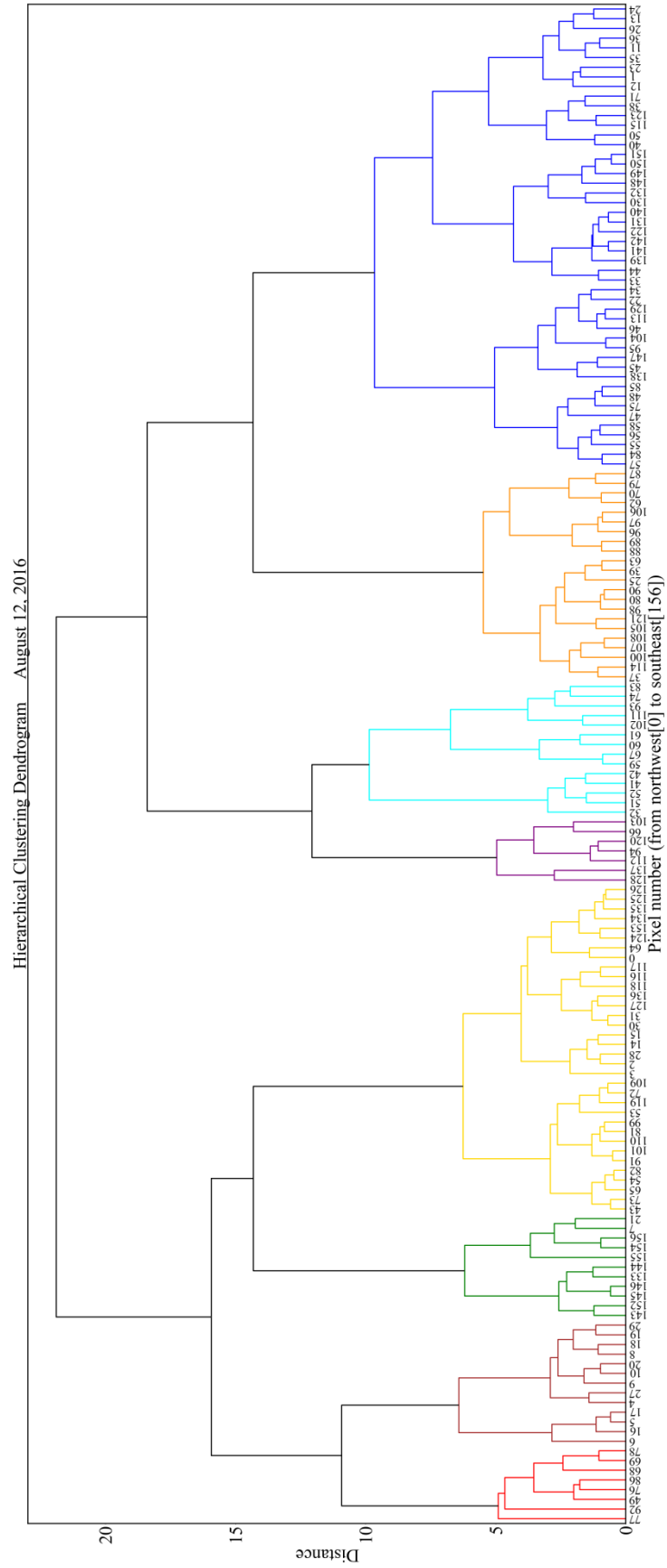


Fig. S1. Dendrograms obtained by the cluster analysis (Ward's method) on August 12, 2016. The horizontal axis represents the pixel number, and the vertical axis represents the Euclidean distance.

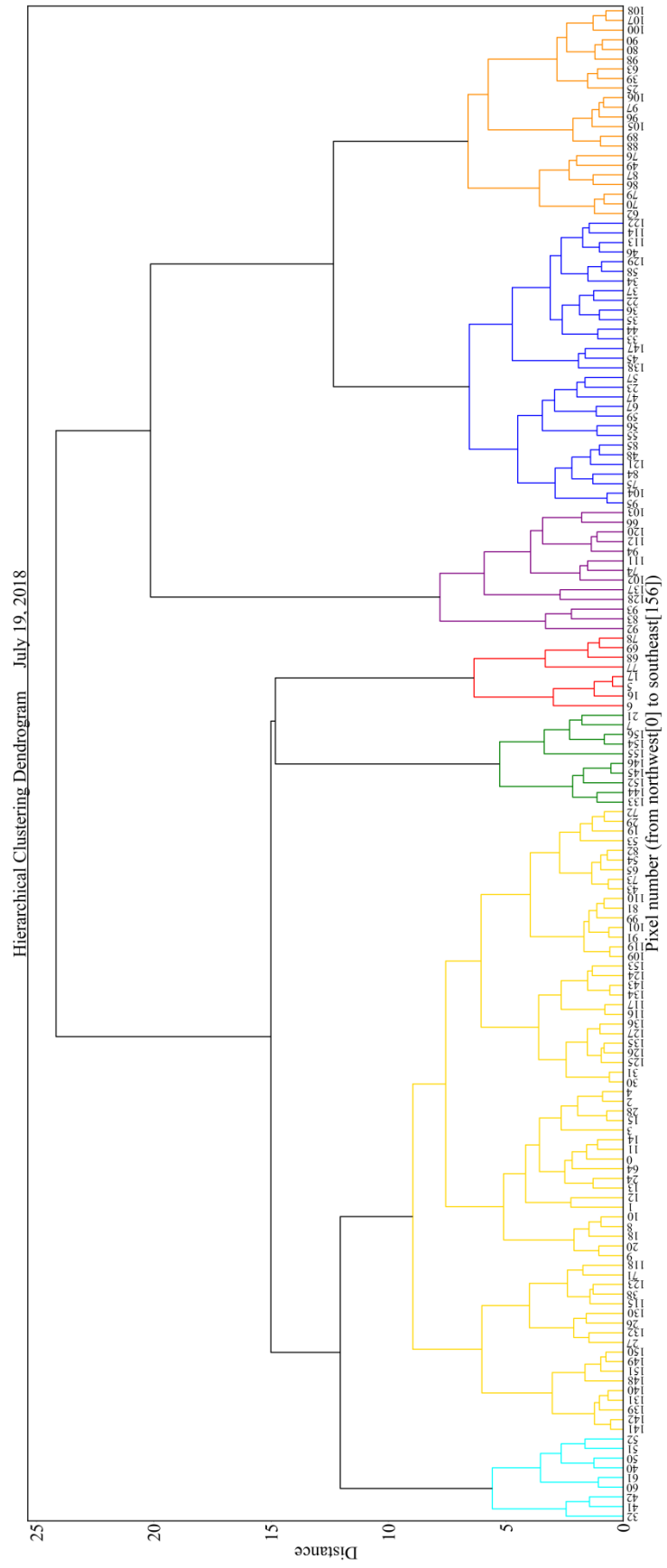


Fig. S2. Similar to Fig. S4, but for July 19, 2018.

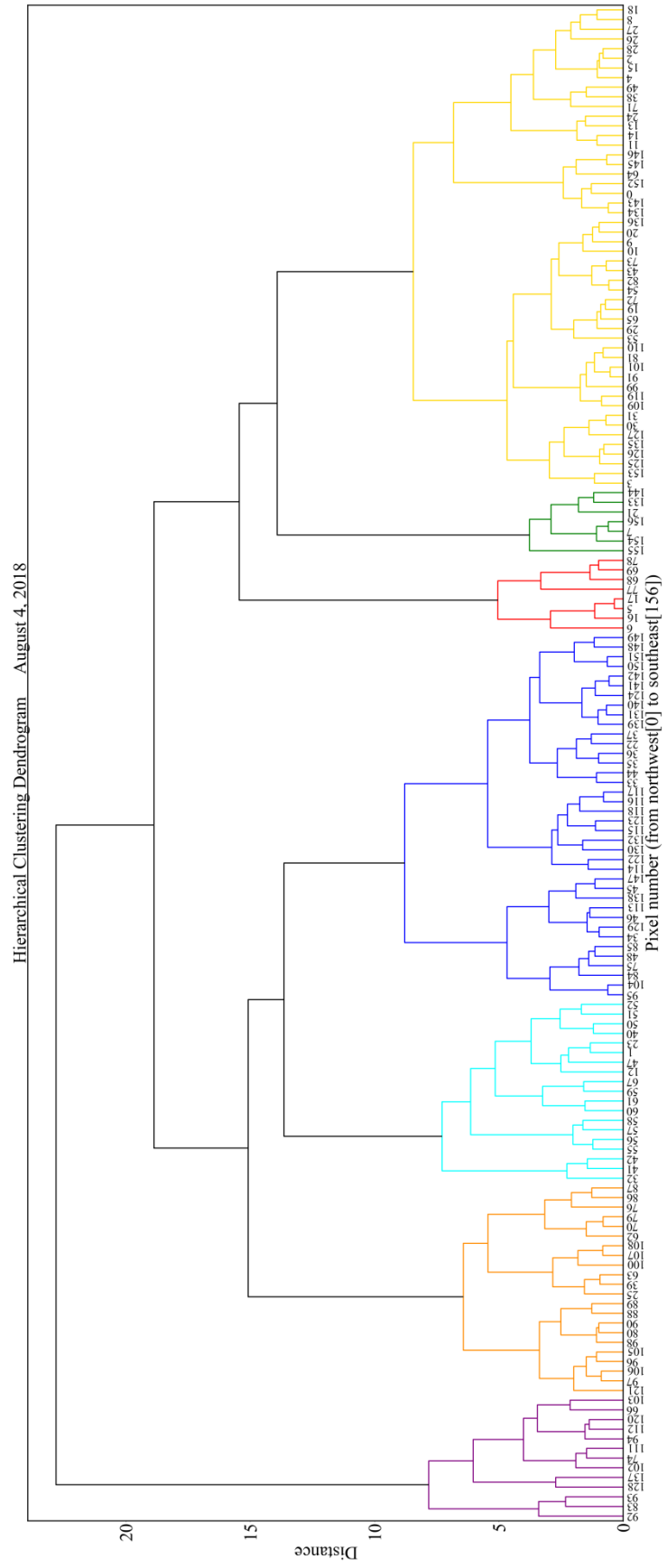


Fig. S3. Similar to Fig. S4, but for August 4, 2018.

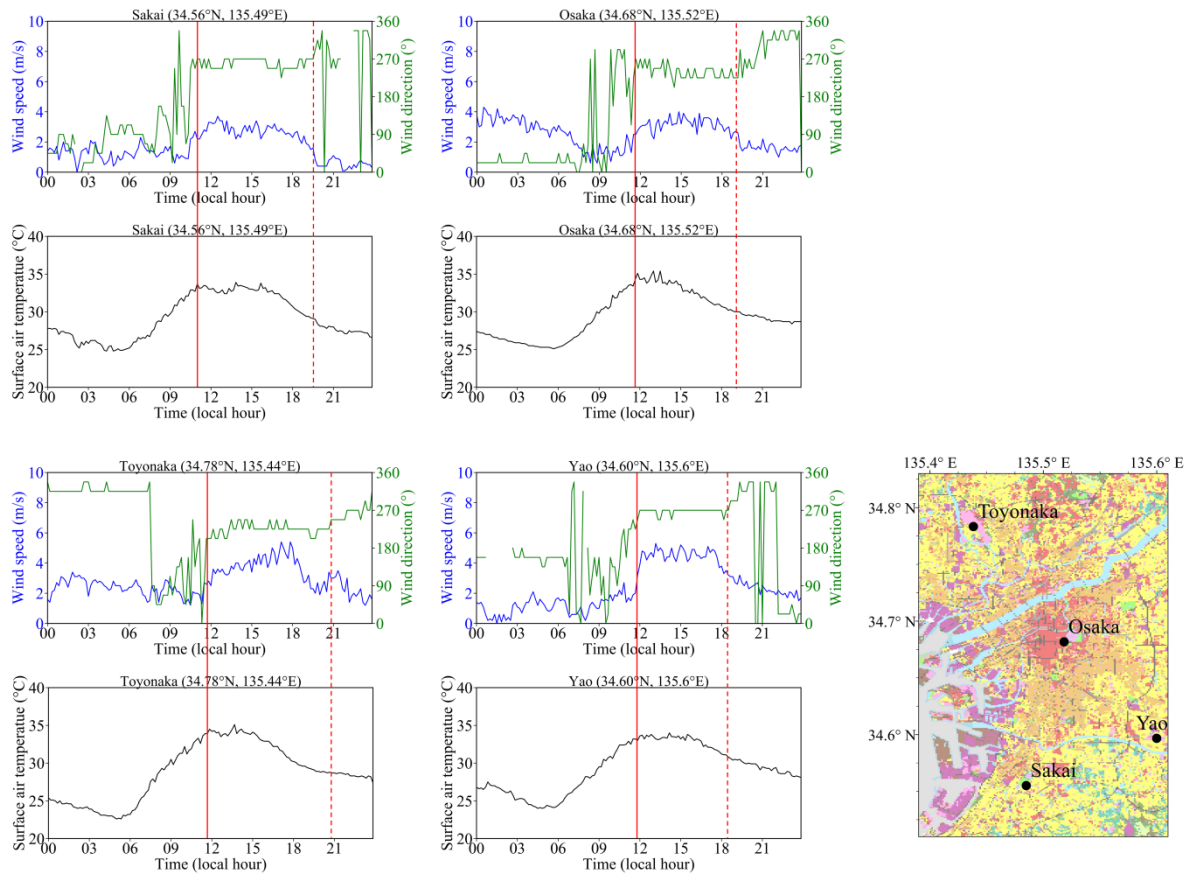


Fig. S4. Time series data of wind speed/direction and SAT observed at the stations of AMeDAS in the study area on August 12, 2016. The lower right figure shows the position of each observation point in the study area. The red solid lines indicate penetration times of the sea breeze roughly judged by changes in wind direction, increase in wind speed, and decrease in SAT. The end times of the sea breeze judged by the change in wind direction and wind speed are indicated by a red dashed line.

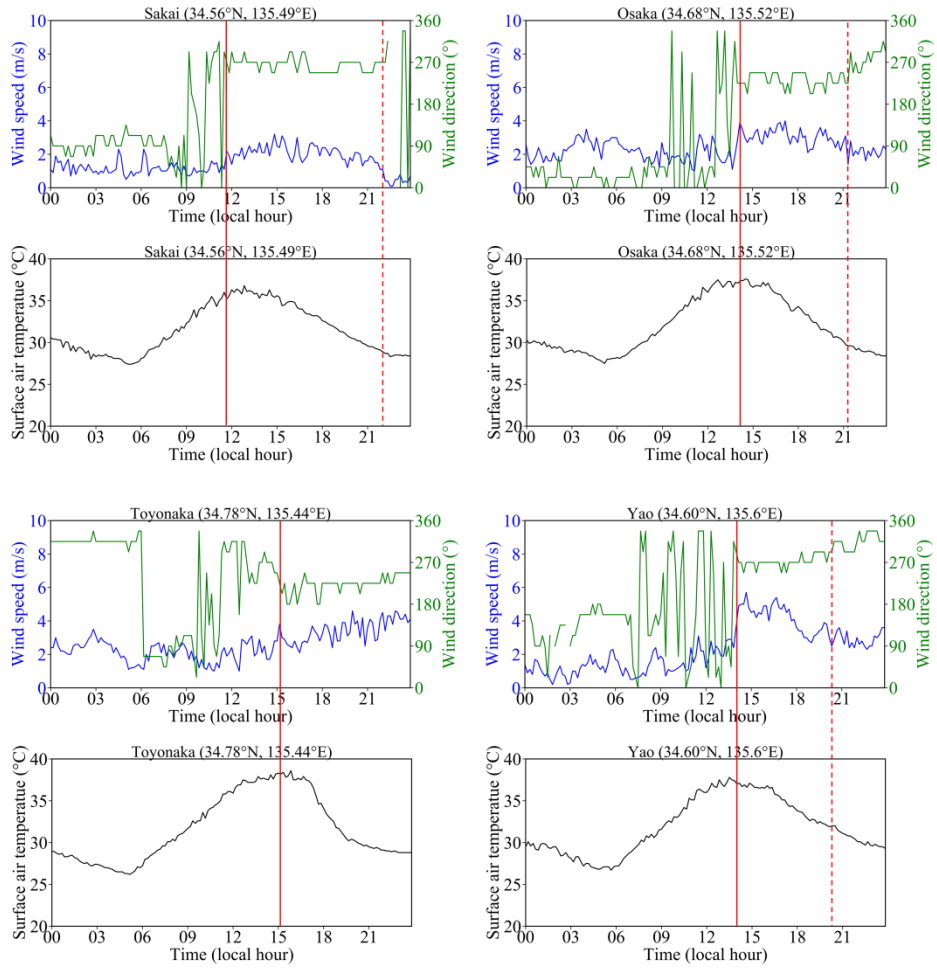


Fig. S5. Similar to Fig. S1, but for July 19, 2018.

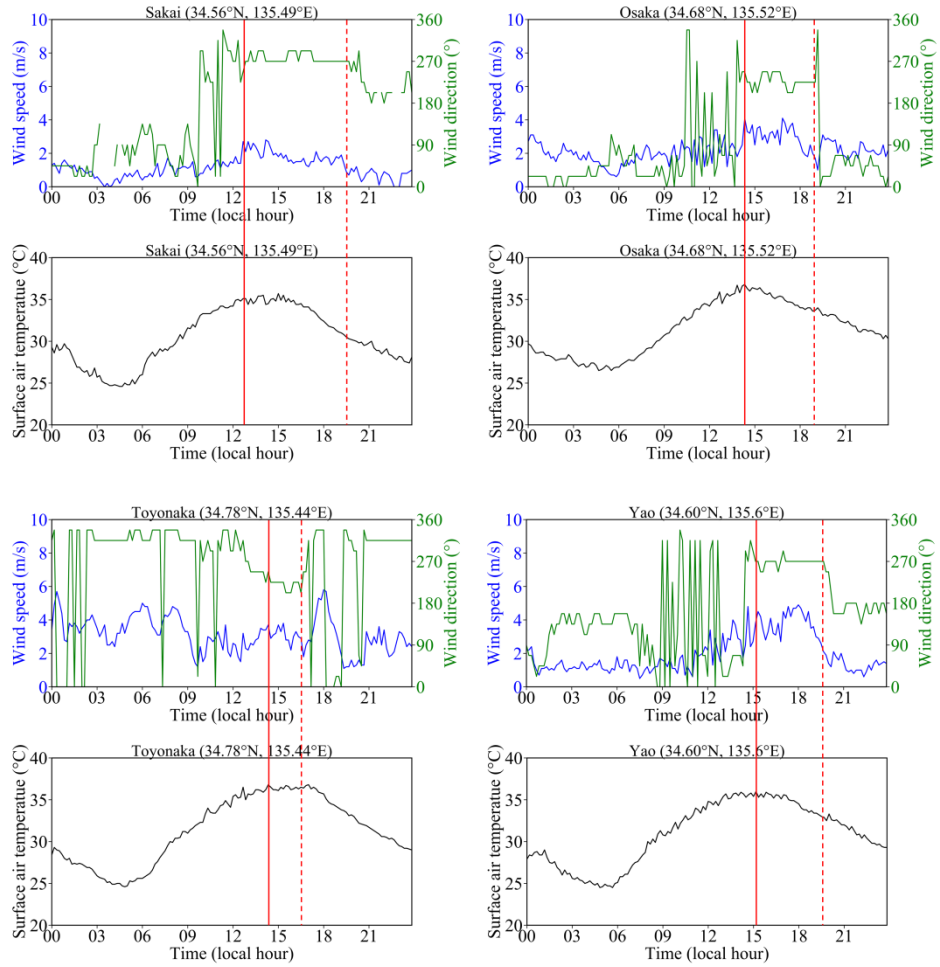


Fig. S6. Similar to Fig. S1, but for August 4, 2018.

Declaration of interests

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

Declarations of interest: none

CRedit author statement

Yuhei Yamamoto: Conceptualization, Data curation, Formal analysis, Investigation, Methodology, Project administration, Software, Validation, Visualization, Writing - original draft, Writing - review & editing.

Hirohiko Ishikawa: Funding acquisition, Resources, Supervision

Urban Climate

Conflict of interest declaration and author agreement form

Title of Paper: Influence of urban spatial configuration and local meteorological field on land surface temperature on summer clear-sky days

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