

CONTINENTAL SLOPE STRUCTURAL HIGH IGNEOUS ACTIVITY AND ITS BEARING ON THE GEOTECTONIC HISTORY OF THE PHILIPPINE SEA

Tsunemasa SHIKI* and Yasuyuki MIYAKE**

* *Department of Geology and Mineralogy, Faculty of Science,
Kyoto University, Kyoto, Japan*

** *Department of Geology, Faculty of Science, Matsue, Japan*

Some characteristic magmatism on the structural highs of continental slopes indicates initiation of subduction of a new, relatively warm slab. An instance is the plagiogranitic and tholeiitic magmatism that occurred on the Middle Miocene topographic and structural highs of the continental slope of the outer zone of Southwest Japan. It was simultaneous with the initiation of subduction of the new Shikoku Basin slab, just following the off-ridge volcanism in the basin. The magmas intruded upward through the remnant fissure of the transform fault which had separated the Southwest Japan arc from the Philippine Sea plate until the cessation of the Shikoku Basin opening.

Application of the idea of "continental slope structural high magmatism" to the igneous rocks of the Daito Ridge Group provides a new hypothetical scenario for the tectonic history of the Northern Philippine Sea. Late Cretaceous collision of the "Kuroshio Paleo-island" with the Japan arc resulted in suspension of northward subduction of the Paleo-Philippine Sea plate, temporal southward subduction of a part of the plate under the Paleo-Daito arc, and the structural high magmatism on the Amami Plateau, the present northernmost ridge in the Daito Ridge Group.

1. Introduction

Occurrence of "continental slope igneous complex" (SHIKI and MISAWA, 1979: or "forearc igneous complex," SHIKI and MISAWA, 1982) is one of the most significant findings in the recent studies on the structure and history of the continental slope-trench region.

Much recent information supports the essential part of the generalized structural model of the Japanese continental slope regions proposed by SHIKI and MISAWA (1979, 1980, 1982) except for a few minor points. Especially, the important role played by thrust movement, which is directly manifested by the tectonic activity of the Wadati-Benioff zone, is gradually being elucidated, though the idea of generation and intrusion of magma along the earthquake zone and the thrust fault is still hypothetical.

One the other hand, however, the idea of the "Nankai Structural High Magmatic Zone" and the "structural high magmatism" (MIYAKE and HISATOMI, 1985) or "forearc magmatism" (MIYAKE, 1985 a) was derived from the study of the Shiono-Misaki (see Fig. 3 for location) igneous complex which was formed in a structural high belt of the Miocene age (MIYAKE and HISATOMI, 1985). The most interesting result of the study of the Shiono-Misaki complex was the marked resemblance in the petrochemical feature of its basaltic rocks with that of the Shikoku Basin off-ridge tholeiite.

The purposes of the present paper are as follows:

- 1) To stress the importance of "(continental slope) structural high igneous complex" and its bearing on subduction-collision tectonic history in the Western Pacific.
- 2) To discuss the geotectonic history of the Northern Philippine Sea, and hypothetical scenario for the change of tectonic setting due to a Late Cretaceous collision at the northern margin of the sea.
- 3) To provide some suggestion on the genetic history of the marginal seas and basins in the transitional zone.

2. Continental Slope Structural High Igneous Complex

Figure 1 shows the topographic and structural section and Fig. 2 shows the geological and geophysical section respectively, of the continental slope of Southwest Japan and Nankai Trough (SHIKI and MISAWA, 1980; SHIKI, 1982). Figure 2 was drawn based on many seismological and geological data including seismic profiling data and dredge samples (HADA *et al.*, 1982; HADA and SUZUKI, 1983 a, b; HILDE *et al.*, 1969; KAGAMI, 1984, 1985; KAGAMI *et al.*, 1983; KIMURA,

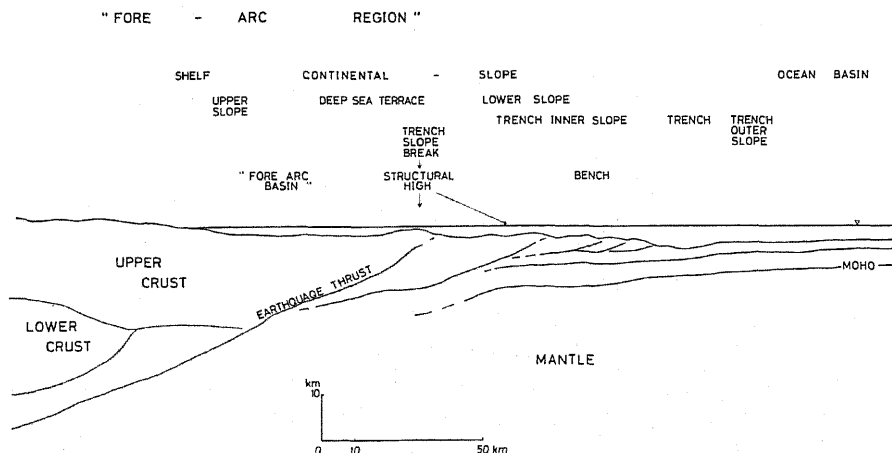


Fig. 1. Simplified topographical and structural section on the Outer Southwest Japan continental slope and Nankai Trough.

1979; KIMURA and OKANO, 1980; NASU *et al.*, 1982; OKANO and KIMURA, 1981; OKUDA, 1977; OKUDA *et al.*, 1979; SHIKI *et al.*, 1971; TAMANO *et al.*, 1983; YOSHII *et al.*, 1973).

Several structural highs, forearc basins, and deep-sea terraces develop between the upper and lower continental slopes. The "(continental slope) structural high igneous complex" occurs exactly on these structural highs (MIYAKE and HISATOMI, 1985). A basaltic rock specimen was recently dredged from one of the structural highs off Shikoku (HONZA *et al.*, 1984). Many concealed igneous masses were suggested in the acoustic data and positive magnetic anomalies were found along modern structural highs (OKUDA *et al.*, 1978; OSHIMA *et al.*, 1980; WATANABE and HATTORI, 1980). The Shiono-Misaki igneous complex, cropping out onshore now, is an excellent example of the continental slope structural high igneous complex of past geological age.

Based on stratigraphic and sedimentological studies, HISATOMI and MIYAKE (1981) made it clear that a topographic barrier developed autochthonously with basin genesis, to dam up the sediments of the Miocene Kumano Group. The Shiono-Misaki igneous complex occurs on this topographic barrier of the Miocene age. This setting of the basin, the topographic high, and the igneous complex in Miocene time resembles well that of the modern age.

The Miocene upheaval zone through the Shiono-Misaki area is very close to one of the modern structural highs (Fig. 3). It suggests that the modern structural high has developed from one of the Miocene magmatic upheaval zones.

MIYAKE (1981, 1985a) examined the sequence of igneous activity and petrochemistry of the Shiono-Misaki igneous complex, and found that all of these igneous rocks are classified as non-alkaline rock. They are markedly low in potassium. All the petrochemical characters of the basaltic rocks of the Shiono-Misaki complex, that is, TiO_2 versus FeO/MgO ratio, titanium versus chromium, chondrite normalized rare earth element values, and K_2O versus SiO_2 , have similarities to the off-ridge basalt (tholeiite) of the Shikoku Basin. They possibly were formed from the source material common with that of the Shikoku Basin off-ridge basalt, but by a rather smaller degree of melting than the latter. This must be noted together with the fact that the magmatism occurred just after the cessation of opening of the Shikoku Basin, that is 15–16 Ma (MIYAKE, 1981).

In the section of Fig. 2, the structural high rock might seem as if it had intruded along an inclined thrust from the deep place beneath the crust. However, it must be considered that no accretionary wedge existed in the Miocene time. The pass of magma intrusion in the Miocene time could be much closer to the Shikoku Basin and steeper. The figure shows present geological section after the accretion of sediments and deformation.

It can be pointed out that a transform fault separated the Southwest Japan arc from the Philippine Sea plate during the time of opening of the Shikoku Basin (Fig. 4). The off-ridge volcanism occurred at the end of the opening in the Shikoku Basin. It seems natural that the fissure, originated from the transform fault, provided the

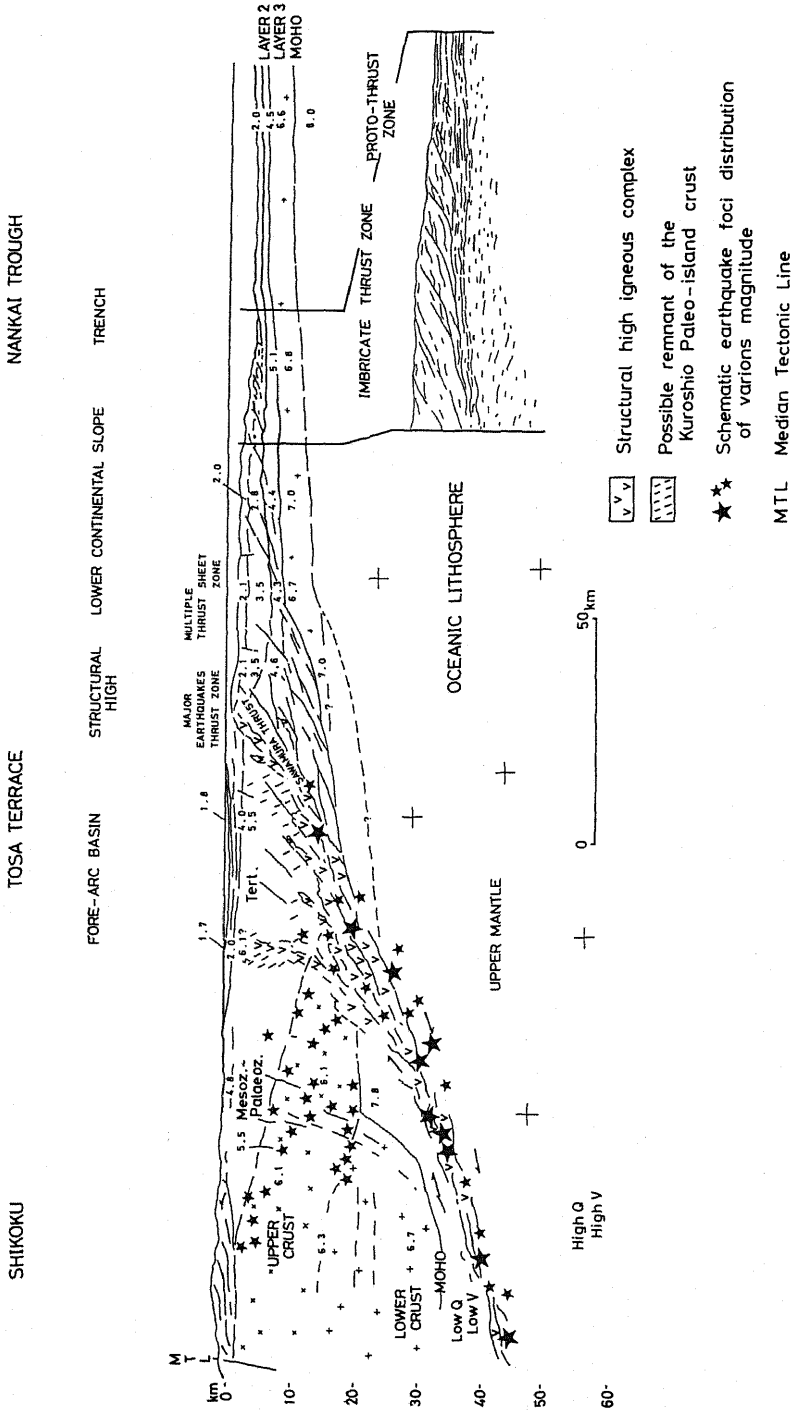


Fig. 2

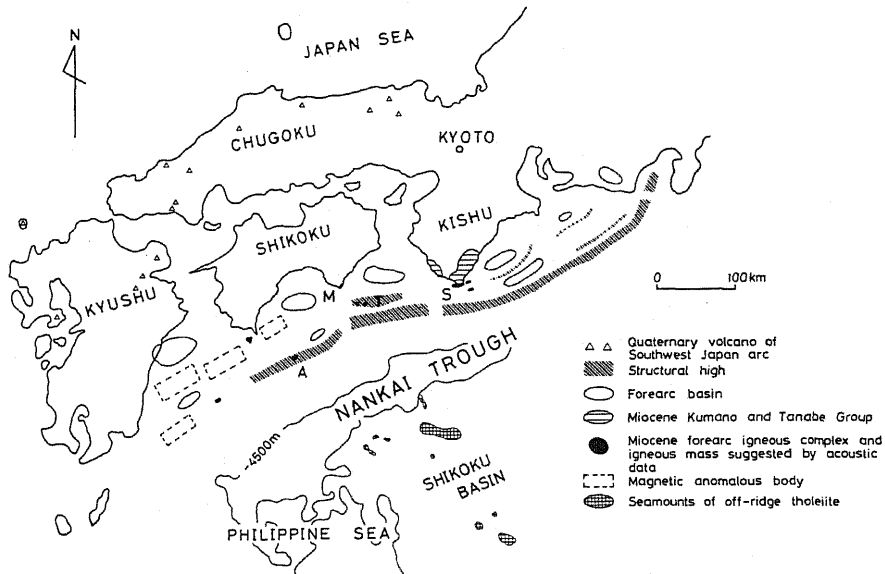


Fig. 3. Map showing the geologic setting of "continental slope structural high igneous complex." Modified from MIYAKE and HISATOMI (1985). S, Shionomisaki; M, Muroto-Misaki; A, Dredge station of basaltic rock.

pass of upward intrusion of magmas just before the initiation of subduction of the new slab of the Shikoku Basin (SHIKI, 1985). MIYAKE (1985b) discussed possible lateral intrusion of materials of the Shikoku Basin asthenosphere under the Southwest Japan arc lithosphere assuming relatively warm and light physical properties of the Shikoku Basin asthenosphere (Fig. 4).

In short, the Miocene continental slope structural high magmatism is a characteristic incident which records the initiation of subduction in the past. There are also reported igneous masses in some other modern continental slope structural highs, though their petrological characters are different from each other. Some of them consist of abyssal tholeiite gabbro and accompanied by low potassic acidic rocks (JAKES and MIYAKE, 1984; MIYAKE, 1985a).

Ophiolitic rocks such as abyssal tholeiite, gabbro, and amphibolite associated

Fig. 2. Geological and geophysical section across the Outer Southwest Japan continental slope, and Nankai Trough showing structure of crust and accretionary prism, earthquake foci distribution, and continental slope structural high igneous complex. Modified from SHIKI (1982). Seismic velocities are based mainly on seismic refraction information by HILDE *et al.* (1969); KIMURA (1979), KIMURA and OKANO (1980), and multi-channel seismic reflection data by NASU *et al.* (1982). Stars: Earthquake foci by OKANO and KIMURA (1981).

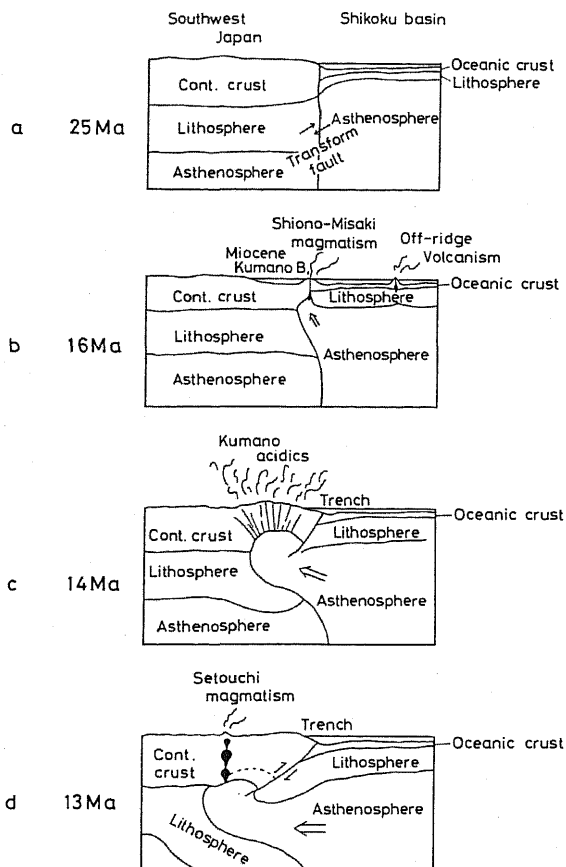


Fig. 4. Sketches of lateral intrusion process of the Shikoku Basin asthenosphere under the crust of Southwest Japan arc, and magmatism in the Miocene.

with granitic rocks occur in much older tectonic belts in the world. Permo-Triassic rocks in the Maizuru Belt and the Kurosegawa Belt in Southwest Japan (Fig. 5) are good examples. It should be pointed out that a large amount of granitic rocks of noticeable petrochemical character, that is, poor in potassium, low in initial $^{87}\text{Sr}/^{86}\text{Sr}$ ratio, were formed there. Some of the rocks have protoclastic texture (SHIKI and MISAWA, 1982).

Tectonic settings of the granitic and other rocks of these tectonic belts were discussed by several authors. MARUYAMA (1978), and HADA and SUZUKI (1983 a, b) proposed the idea of occurrence of ophiolitic rocks along the suture zone between the island arc crust and the oceanic plate. A similar idea was presented by SHIKI (1978, 1982), and SHIKI and MISAWA (1979, 1980) in connection with subduction and accretion tectonics and activity of the Wadati-Benioff zone.

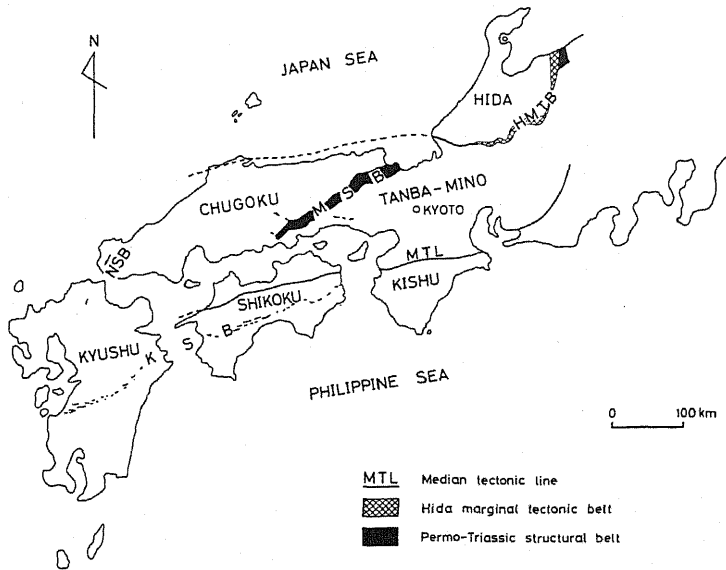


Fig. 5. Permian—Triassic structural belt in Southwest Japan. MSB, Maizuru Structural Belt; KSB, Kurosegawa Structural Belt; NSB, Nagato Structural Belt.

3. Geotectonic History of the Northern Philippine Sea

On the basis of a comprehensive study of numerous data of the Philippine Sea and the margin of Southwest Japan arc including the petrochemical character of rocks stated above, SHIKI (1985) proposed a hypothetical scenario for the geotectonic history of the Northern Philippine Sea. This scenario implies subduction of marginal sea floors which existed between the Southwest Japan arc and the Kuroshio Paleo-islands, and the Daito Paleo-islands.

As was reported by AOKI *et al.* (1975, 1976), SHIKI *et al.* (1977, 1979), potassium-poor granite and gabbro occur on the Amami Plateau, the northernmost member of the Daito Ridge Group. This granite has a notable resemblance with that of the Shiono-Misaki, that is, it is poor in potassium, low in initial $^{87}\text{Sr}/^{86}\text{Sr}$ ratio. This fact, together with its association with gabbro and marginal situation in the Daito Ridge Group Region, suggests the possibility that the rocks belong to the structural high igneous complex of the "Daito Ridge Group arc" (SHIKI, 1985).

High P and low T metamorphic rocks develop in the Daito Ridge, which forms the central part of the ridge group (MIZUNO *et al.*, 1975, 1976; SHIKI *et al.*, 1985). In the case of the Plagiogranite-bearing "ophiolitic rocks" of the Maizuru Belt and Sangun metamorphics, the former is situated on the outer side (ocean side). As pointed out by SHIKI (1985), if there is a similarity in tectonic setting between the Daito Ridge Group and the Maizuru Belt, the ocean basin (marginal sea basin) must have existed in the northern side of the ridge group, and the Cretaceous

subduction must have occurred southward from the north. This polarity of subduction is opposite to the general direction of the movement of the Philippine Sea plate during the time of spreading of the West Philippine Basin stated by many authors (e.g., KARIG, 1975; SENO and MARUYAMA, 1984; UYEDA and BEN-AVRAHAM, 1972). The southward subduction was an exceptional incident which happened following a collision at the southern margin of the Southwest Japan arc.

Modern geological structure and tectonic history of the continental slope of Southwest Japan and Nankai Trough have long been investigated. Here we introduce a pre-Miocene geotectonic model of 1) Continental slope structural high igneous activity at the beginning of subduction of the Shikoku Basin, and 2) collision of remnant of the "Kuroshio Paleo-islands" with the Southwest Japan arc.

Members of the Shimanto Research Group proposed that a piece of land had extended from the south of the Southwest Japan arc to the east of the Ryukyu arc from Late Mesozoic to Paleogene, judging from the direction of paleocurrent and the distribution of orthoquartzite gravels observed in the Shimanto Major Belt in Southwest Japan (HARATA and TOKUOKA, 1974; SUZUKI and TATEISHI, 1975). They named the land "Kuroshio Paleoland." However, no evident remnant crust of the Kuroshio Paleoland has been found in the Philippine Sea floor. Most probably, the remnant had collided and joined with the crust of the Southwest Japan arc and the Ryukyu arc as suggested by NAKAZAWA *et al.* (1983). Facies change of sediments of the Shimanto Major Belt through the Late Mesozoic to the Paleogene indicates such a tectonic event. Probable trench sediments appear in formations older than Santonian, whereas much coarser and younger sediments including endolistostrome derived from the south reflect the upheaval of the southern provenance.

It must be noted that the inferred time of collision of the Kuroshio Paleo-island is very close to or correlative with that of the magma intrusion at the northern margin of the Daito Paleo-island arc. In this connection, the structural high magmatism is a special event which occurred only immediately prior to or at the initial stage of subduction of relatively warm oceanic plate.

Putting these pieces of information together, a hypothetical scenario, as shown in Fig. 6, can be proposed for the changes of the tectonic setting in the Northern Philippine Sea from the Late Cretaceous to the Miocene. Subduction of the unnamed part of the Paleo-Philippine Sea plate of the Mesozoic time (Kula plate?) was suspended when the Kuroshio Paleo-islands collided with the Japan arc. Subduction of a part of the plate to the south under the Paleo-Daito arc occurred in the Latest Cretaceous time.

The volcanic activities of the Paleocene and Earliest Eocene time on and near the Daito Ridge (MCKEE and KLOCK, 1980; OZIMA *et al.*, 1980; TOKUYAMA *et al.*, 1980) was related also with the subduction tectonics from the north. This subduction, together with older northward subduction, was related with the opening of the West Philippine Basin.

The southward subduction probably ceased soon. New subduction of the

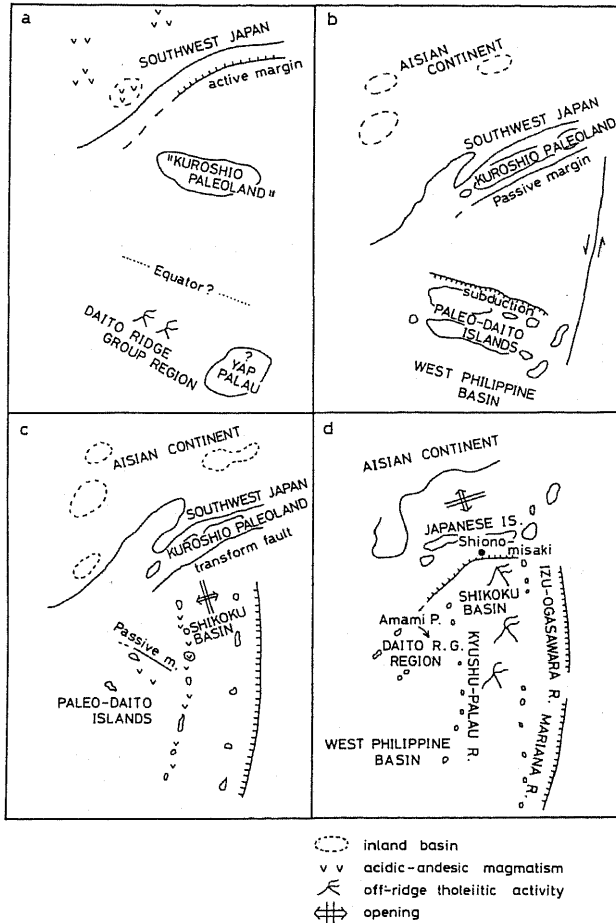


Fig. 6. Schematic scenario for change of the geotectonic setting of subduction in the Northern Philippine Sea. a, Late Cretaceous; b, Paleocene-Eocene; c, Oligocene; d, Middle Miocene. Modified from SHIKI (1985).

Philippine Sea plate began from the south to the north under the Kuroshio Paleoland mass of the Southwest Japan arc.

Opening of the Shikoku Basin in the Oligocene and Early Miocene time has been well clarified (KOBAYASHI and NAKADA, 1978; KOBAYASHI, 1983). It was followed by off-ridge volcanism in the basin, subduction of the new slab of the Shikoku Basin, and upward intrusion of magma through the remnant fissure of transform fault of the Shikoku Basin opening.

It is noteworthy that the time of clockwise rotation of the Southwest Japan (almost 15 Ma; OTOFUI and MATSUDA, 1983; OTOFUI *et al.*, 1985, 1986) was very

close to that of the Shiono-Misaki igneous activity. Discussion of this problem, however, is beyond the purposes of the present paper.

After the consolidation of the Shiono-Misaki igneous complex, subduction zone shifted southward slightly, to the outer side of the complex. Thus, younger structural highs and structural high magmatic zone are situated slightly south of the older structural high.

In this paper, we stressed the importance of continental slope structural high complex ("forearc igneous complex") and provided a hypothetical scenario of the geohistory of the Northwestern Philippine Sea from the Late Cretaceous to the Present. It is clear that extensive further investigations and studies are required to check the hypothetical history stated above.

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