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RESEARCH ARTICLE

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Key Points:

- MHD simulation is used to investigate the evolution of substorm under different Earth's dipole magnetic moment M
- For lower M, expansion onset takes place later, auroral electrojet develops slowly, and the maximum auroral electrojet increases
- For lower *M*, auroral oval shifts equatorward during the growth phase and expands more equatorward and poleward during the expansion phase

Supporting Information:

Figure S1

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How Do Auroral Substorms Depend on Earth's Dipole Magnetic Moment?

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Abstract Earth's dipole magnetic moment M is known to decrease by $\sim 9\%$ over the past 150 years. It has been argued that the decrease in M makes the near-Earth space environment different. We investigated how the change in M affects the development of an auroral substorm by increasing and decreasing M by a factor of 1.5 in global magnetohydrodynamics simulation. The ionospheric conductivity decreases with increasing M, in accordance with the aid of empirical relations. When we imposed the southward interplanetary magnetic field, an auroral substorm took place regardless of M, but its development depends largely on M. When M is lower, (1) the expansion onset takes place later, (2) the auroral electrojet develops slowly, and (3) the maximum auroral electrojet increases. The first two consequences are probably associated with the slow magnetospheric convection as manifested by the polar cap potential drop. The third consequence is associated with the nonlinear dependence of substorm-associated field-aligned currents (FACs) on the ionospheric conductivity. The maximum values of the westward auroral electrojet and the net FACs increase with decreasing M, whereas the incident magnetic energy into the magnetosphere decreases with decreasing M. This implies that the efficiency of the generation of the substorm-associated FACs increases with decreasing M. It is also found that, for the lower M-value, the auroral oval shifts equatorward during the growth phase and expands more equatorward and poleward during the expansion phase. Evolution of substorms depends largely on the value of Earth's dipole moment and the ionospheric conductivity.

Plain Language Summary Earth's intrinsic magnetic field inhibits the direct entry of the fast streaming gas coming from the Sun, known as solar wind. Observations have shown that Earth's magnetic field has decreased by ~9% over the past 150 years. The reduction of Earth's magnetic field may make the near-Earth space environment different. Here, we show simulation results of auroral substorm under different strengths of Earth's magnetic field. The auroral substorm is one of the drastic, transient phenomena taking place in near-Earth space environment, which is accompanied with sudden brightening of aurora and sudden intensification of electrojet flowing in the upper atmosphere. If Earth's magnetic field keeps decreasing like this, the auroral electrojet will be larger during the expansion of the auroral substorm, the auroral oval will expand more to north and south, and the auroral electrojet will develop slowly.

1. Introduction

Earth's global magnetic field appeared approximately 3.2 billion years ago (Tarduno et al., 2007). The dipole moment of Earth's magnetic field *M* has decreased by ~9% over the past 150 years (Olson & Amit, 2006). The substantial decrease in the dipole magnetic moment is suggested to have a significant impact on near-Earth space environment. The ionospheric conductance has been thought to increase with decreasing with *M* (Cnossen et al., 2011, 2012; Glassmeier et al., 2004; Takeda, 1996, 2013; Tao et al., 2017). For example, the Pedersen conductivity above 75 km altitude (where electron cyclotron frequency is much larger than collision frequency) is given by (Kelley, 2009, p. 45)

$$\sigma_P = \frac{ne^2}{m_i v_{in} \left(1 + \kappa_i^2\right)},\tag{1}$$

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where *n*, *e*, *m_i*, *v*_{in}, and *κ_i* are the electron density, the elementary charge, the ion mass, the ion-neutral collision frequency, and the ratio of cyclotron frequency to collision frequency, respectively. Assuming that the electron density and the ion composition are independent of *M*, we can find that σ_P increases with decreasing magnetic field strength. Glassmeier et al. (2004) suggested the scaling relations for the height-integrated Pedersen and Hall conductivities to be $\Sigma_P \propto M^{-1}$ and $\Sigma_H \propto M^{-4/3}$. Tao et al. (2017) performed the Ground-to-topside model of Atmosphere and Ionosphere for Aeronomy with various dipole moments and found that the reduced dipole moment results in modifications of the thermospheric dynamo field (αM^1), the ionospheric current (αM^{-1}), and the ionosphere–Thermosphere model shows that the amplitude of the *S*_q current decreases with increasing *M* (Cnossen et al., 2011). Interestingly, the Joule heating power, the global mean thermospheric temperature, and the global mean height of the peak of the ionospheric *F*₂ layer do not show a monotonic increase or decrease with *M*.

The decrease in *M* is also thought to have an impact on the magnetospheric structure. One such impact is the reduction of the stand-off distance of the magnetopause with decreasing *M* (Cnossen et al., 2012; Siscoe & Chen, 1975). The global magnetohydrodynamics (MHD) simulation predicts that the shape of the magnetopause is also changed when *M* is less than 4×10^{22} A m² (Cnossen et al., 2012). Glassmeier et al. (2004) considered the extension of the magnetopause reconnection and suggested that the polar cap potential is proportional to the magnetopause stand-off distance. Assuming that the stand-off distance is proportional to $M^{1/3}$, they proposed that the polar cap potential is proportional to $M^{1/3}$. Cnossen et al. (2012) performed the global MHD simulation and pointed out that the stand-off distance does not simply obey the scaling. Siscoe and Christopher (1975) suggested that the equatorward boundary of the auroral oval moves equatorward with decreasing *M*. They regarded the inner boundary of the convection region as the equatorward boundary of the auroral oval. The net field-aligned currents (FACs) increase with decreasing *M* because of the increase in the ionospheric conductivity (Cnossen et al., 2012; Zieger, Vogt, Ridley, et al., 2006). Glassmeier et al. (2004) considered the scaling relations for the ionospheric conductivity and the electric field and predicted that the intensity of the auroral electrojet decreases with decreasing *M*.

There were many discussions on the influence of *M* on quasi steady state conditions of the near-Earth space environment. As far as we know, little discussion has been made on the influence of *M* on transient phenomena, such as a substorm. The substorm is one of the most drastic, transient disturbances taking place in near-Earth space environment. The intensification of the auroral electrojet is one of the manifestations of it (Akasofu et al., 1965; Kamide & Akasofu, 1975). The Joule dissipation rate in the ionosphere sometimes increases to ~10¹¹ W during the expansion phase of the substorm (Ahn et al., 1983; Kamide et al., 1996; Palmroth et al., 2005; Richmond et al., 1990; Rostoker et al., 1980; Tanskanen et al., 2002; Zhou et al., 2011). The ε parameter (Perreault & Akasofu, 1978) is used to estimate the amount of energy penetrating into the magnetosphere. Observations have shown that when the ε parameter exceeds 10¹¹ W, a substorm occurs (Akasofu, 1980). Originally, the ε parameter was derived on the basis of the Poynting flux in the solar wind interacting with the magnetosphere (Perreault & Akasofu, 1978). Later, the ε parameter is suggested to be identical to the dynamo power from the solar wind to an open magnetosphere (Kan et al., 1980). The amount of intake energy, as well as the energy stored in the magnetosphere, is expected to depend on the scale of the magnetosphere (Gonzalez & Mozer, 1974; Siscoe & Cummings, 1969).

Recent global MHD simulations helps understand the pathway and the conversion of energy from the solar wind to the ionosphere (Ebihara & Tanaka, 2017; Ebihara et al., 2019). According the global MHD simulation, the magnetic energy penetrating into the magnetosphere comes from two sources: magnetic energy of interplanetary magnetic field (IMF) and solar wind kinetic energy. About 33%–88% of the magnetopause passing through the magnetopause is converted from the solar wind kinetic energy near the magnetopause (Ebihara et al., 2019). In the magnetosphere, the large-scale, earthward transfer of the magnetic energy is associated with the FACs. When IMF is southward, dynamo that is responsible for the generation of the large-scale FACs appears in the mantle region (Tanaka et al., 2010). Just before the substorm expansion, additional dynamo appears in the near-Earth region. The near-Earth dynamo is closely associated with the substorm-associated FACs. Snapshots of the FACs obtained by Active Magnetosphere and Planetary Electrodynamics Response Experiment clearly demonstrate the enhancement of the FACs during the substorm expansion (Murphy et al., 2013). The substorm-associated FACs are thought to be responsible for the additional transport of the magnetic energy to the Earth during the expansion phase.



Table 1 Simulation Settings for the Five Runs						
Name of run	Dipole magnetic moment M (10 ²² A m ²)	Ionospheric conductivity				
Run 1	7.76 (standard run)	Standard				
Run 2	11.64	$\Sigma_{\rm P} \propto M^{-1}$ and $\Sigma_{\rm H} \propto M^{-4/3}$				
Run 3	5.17	$\Sigma_{\rm P} \propto M^{-1}$ and $\Sigma_{\rm H} \propto M^{-4/3}$				
Run 4	11.64	Standard				
Run 5	5.17	Standard				

The purpose of this study is to illuminate the possible influence of M on the auroral substorm on the basis of the results obtained by the global MHD simulation. We arbitrarily changed M in the simulation. The quadrupole component is excluded. We admit that the simulation results are somewhat speculative, but we believe that the results are useful to predict possible near-Earth space environment and to understand the solar wind-magnetosphere–ionosphere coupling that happens on Earth today as well as on other magnetized planets.

2. Simulation

We used the global MHD simulation developed by Tanaka (2015). This simulation consists of the solar wind-magnetosphere domain and the ionosphere domain, and they couple with each other. The solar wind-magnetosphere domain ranges from the geocentric distance of 2.6 R_E to 200 R_E at midnight and 600 R_E at noon. The ionosphere domain covers the ionosphere from pole to pole. The ionospheric conductivity is calculated on the basis of the FAC, the plasma pressure, and the temperature at the inner boundary of the solar wind-magnetosphere domain. The ionospheric conductivity increased by the plasma pressure and temperature is regarded as the diffuse aurora, while the conductivity increased by the (upward) FAC is regarded as the discrete aurora. We mapped the FAC from the inner boundary of the magnetospheric domain to the ionospheric domain. To satisfy the requirement of the current continuity, we solved an elliptic partial differential equation with a thin shell approximation and obtained the electric potential. The electric field is mapped from the ionosphere domain to the magnetosphere domain and is used to impose the boundary condition of the magnetosphere domain. The derivation of the ionospheric conductivity is given by Ebihara et al. (2014).

For the purpose of obtaining a quasi-steady condition, we imposed the steady solar wind condition with a speed of 400 km/s, a density of 5 cm⁻³, IMF B_y of 0, and IMF B_z of 3 nT for 2 h. Then, we turned IMF B_z from 3 nT to -5 nT at t = 120 min as a step function in the upstream solar wind. The dipole moment of the Earth M was set to be 7.76×10^{22} A m², which corresponds to the present Earth. This is a standard run, which is hereinafter referred to as Run 1. In Run 2, we increased M by 50%, that is, 11.64×10^{22} A m², which corresponds to the Earth about 2,650 years ago (Knudsen et al., 2008). In Run 3, we decreased M by 50%, that is, 5.17×10^{22} A m². If the rate of change in M remains constant in future, this M-value will correspond to the Earth in about 2850 AD. In Runs 2 and 3, we multiplied the height-integrated Pedersen and Hall conductivities with the aid of the scaling relations, $\Sigma_P \propto M^{-1}$ and $\Sigma_H \propto M^{-4/3}$, suggested by Glassmeier et al. (2004). Runs 4 and 5 are the same as Runs 2 and 3, respectively, except that the ionospheric conductivities in Runs 4 and 5 in the same manner as Run 1. Runs 4 and 5 are probably unrealistic because the ionospheric conductivity is independent of M but are used to isolate the contribution from the ionospheric conductivity. The simulation settings are summarized in Table 1.

3. Results

Figure 1 shows the calculated auroral electrojet indices, SMU and SML. The SMU/SML indices are extension of the AU/AL indices to overcome the ambiguities arising from the limited number of observatories (Newell & Gjerloev, 2011a). We calculated the magnetic disturbances induced by the ionospheric Hall current at magnetic latitudes (MLAT) from 50° to 90° with an interval of 1° in MLAT and 0.5 h in the magnetic local time. The auroral electrojet is known to be closely associated with the auroral activities (Akasofu et al., 1965; Kamide & Akasofu, 1975; Obayashi & Nishida, 1968). Originally, an abrupt intensification of the auroral brightness is used to identify the onset of expansion phase of substorms (Akasofu, 1964). Here, we use the sudden enhancement of the auroral electrojets to identify substorms onset (Akasofu et al., 1965; Newell & Gjerloev, 2011a, 2011b; Rostoker et al., 1980). Following Newell and Gjerloev (2011a), we determined the expansion onset t_0 when all the following criteria are satisfied.

$$SML(t_0 + 1\min) - SML(t_0) < -15\,\mathrm{nT}$$
⁽²⁾

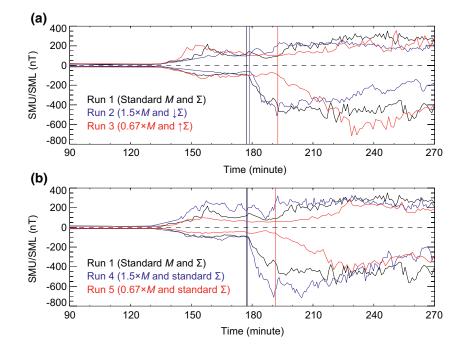


Figure 1. Calculated SMU and SML indices. The vertical lines indicate the time of expansion onset.

$$SML(t_0 + 2\min) - SML(t_0) < -30 \text{ nT}$$
(3)

$$SML(t_0 + 3\min) - SML(t_0) < -45\,\mathrm{nT} \tag{4}$$

$$\frac{1}{26} \sum_{i=4}^{30} SML(t_0 + i \min) - SML(t_0) < -100 \text{ nT}$$
(5)

The expansion onset determined by the above criteria is displayed by the vertical lines in Figure 1. The minimum values of SML within 60 min from the expansion onset are summarized in Table 2.

The black lines in Figure 1 show the SMU and SML indices for Run 1 (which corresponds to the present Earth). The lower line indicates SML, which starts to decrease gradually at $t \sim 136$ min and stays almost constant until $t \sim 177$ min. This period corresponds to the growth phase of the substorm (McPherron, 1970). SML shows a sudden decrease at $t \sim$ 177 min. This moment is identified as the expansion onset, according to the above criteria (Newell & Gjerloev, 2011a). SML reached the minimum value of -579 nT within 60 min from the expansion onset. For Run 2 (in which M is 1.5 times larger than present and the ionospheric conductivity decreases), SML recovers earlier than for Run 1. The minimum value of SML is -513 nT. For Run 3 (in which M is 1.5 times smaller than present and the ionospheric conductivity increases), SML decreases slowly in comparison with those for Runs 1 and 2. The minimum value of SML is -717 nT, which is smaller than that for Runs 1 and 2. The result indicates that the maximum intensity of the westward electrojet increases with decreasing M. When the ionospheric conductivity is independent of M, the situation is different (as shown in the bottom panel of Figure 1). SML reached the minimum value of -714 nT for

Table 2

Minimum SML Within 60 min From the Expansion Onset, Cross-Polar Cap Potential Drop at the Expansion Onset, Maximum Speed of Sunward Flow in Near-Earth Plasma Sheet at Onset, Rate of Magnetic Energy Generated in Near-Earth Region (Near-Earth Dynamo) at Onset, and Incident Magnetic Energy Into the Magnetosphere at Onset

	Maximum speed of				
		Cross-	sunward		Incident
		polar	flow in	Near-	magnetic
		cap	plasma	Earth	energy into
Name		potential	sheet at	dynamo	magnetosphere
of	Minimum	at onset	onset	at onset	at onset
run	SML (nT)	(kV)	(km/s)	$(10^{10} \mathrm{W})$	$(10^{10} \mathrm{W})$
Run 1	-579	75	406	5.7	118
Run 2	-513	86	772	7.1	137
Run 3	-717	57	246	3.9	96
Run 4	-714	83	891	7.4	136
Run 5	-462	72	267	3.5	96



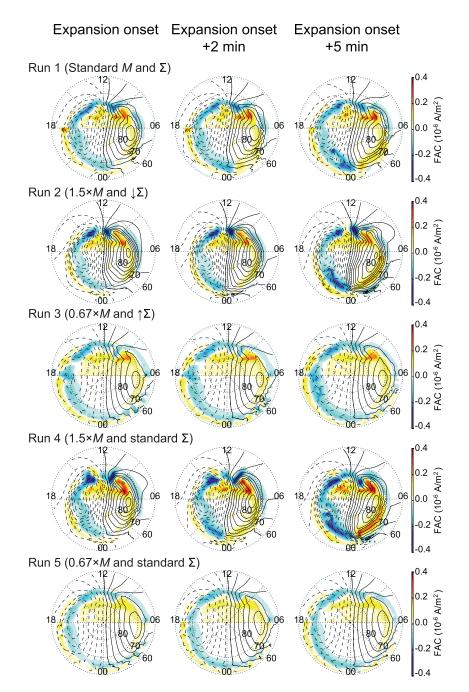


Figure 2. Field-aligned currents at the ionosphere altitude (left) at the expansion onset, (middle) 2 min after the onset, and (right) 5 min after the onset. Positive downward (red) and negative upward (blue). The Sun is to the top. The contour lines indicate the ionospheric electric potential at an interval of 5 kV. The solid lines mean the positive potential, whereas the dashed ones mean the negative potential.

Run 4, whereas it reached the minimum value of -462 nT for Run 5. Namely, the maximum intensity of the westward electrojet increases when we increase the ionospheric conductivity.

Figure 2 summarizes the FACs for the first 5 min from the expansion onset. At a glance, the global distributions of FACs resemble the Region 1 and Region 2 FACs (Iijima & Potemra, 1976) for all the runs. Both the Region 1 and Region 2 FACs are intensified after the expansion onset, in particular, on nightside. This tendency is consistent with observations (Coxon et al., 2014). For Runs 1 and 2, a surge-like structure of the upward FACs appears in the postmidnight sector immediately after the expansion onset. In the simulation,



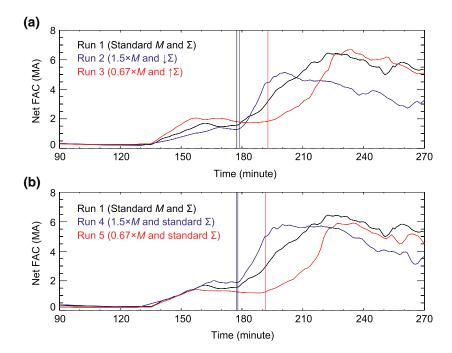


Figure 3. Net field-aligned current flowing into and out of the ionosphere in both the hemispheres. The vertical lines indicate the time of expansion onset.

the surge-like structure traveling westward is caused by the mutual interaction between the magnetosphere and the ionosphere (Ebihara & Tanaka, 2015a, 2018). For Run 3, the surge-like structure appears later (not shown). The FAC distribution for Run 4 is similar to that for Run 2, whereas the intensity of the FACs for Run 4 is somewhat larger than for Run 2. The FAC distribution for Run 5 is also similar to that for Run 3. The intensity of the FACs for Run 5 is smaller than for Run 3. There is a tendency that the intensity of the substorm-associated FACs increases with increasing M.

The contour lines overlaid on the FACs in Figure 2 represent the electric potential. The two-cell convection pattern is confined at high latitudes for the large *M*-value as Siscoe and Christopher (1975) pointed out. The cross-polar cap potential (CPCP) drop in the Northern Hemisphere at the expansion onset (the end of the growth phase) is summarized in Table 2. A clear tendency can be identified that the CPCP increases with increasing *M*. This is consistent with a theoretical study (Siscoe & Chen, 1975) and simulation studies (Cnossen et al., 2012; Zieger, Vogt, & Glassmeier, 2006).

Table 3

Maximum Value of the Net FAC, Joule Heating Rate in the Ionosphere, Rate of Magnetic Energy Generated in Near-Earth Region (Near-Earth Dynamo), and Incident Magnetic Energy Into the Magnetosphere at the Moment When the Net FAC Reaches Maximum

Name of run	Net FAC (MA)	Joule heating rate in the ionosphere (10 ¹⁰ W)	Near-Earth dynamo (10 ¹⁰ W)	Incident magnetic energy into magnetosphere (10 ¹⁰ W)
Run 1	6.5	3.0	48	132
Run 2	5.1	3.4	50	176
Run 3	6.7	3.3	24	109
Run 4	5.8	3.3	48	194
Run 5	5.9	3.2	22	97

Figure 3 shows the net FAC flowing into and out of the ionosphere in both the hemispheres. For Runs 1–3, the net FAC increases with decreasing M during the growth phase. This is consistent with the simulation study (Cnossen et al., 2012). During the expansion phase, the net FAC increases faster and starts to recover earlier for larger M-value. The maximum value of the net FAC decreases with increasing M. When the ionospheric conductivity is independent of M (Runs 1, 4, and 5), the net FAC increases with increasing M during the growth phase. This tendency is different from the cases that the ionospheric conductivity depends on M (Runs 1–3). During the expansion phase, the maximum values of the net FAC are almost independent of M. (They are summarized in Table 3.) These results imply that higher conductivity gives rise to higher net FACs and that the contribution from the ionospheric conductivity significantly regulates the net FACs.

Figure 4 summarizes the FAC, the nondiagonal component of the ionospheric conductivity (Hall conductivity), the intensity of the ionospheric



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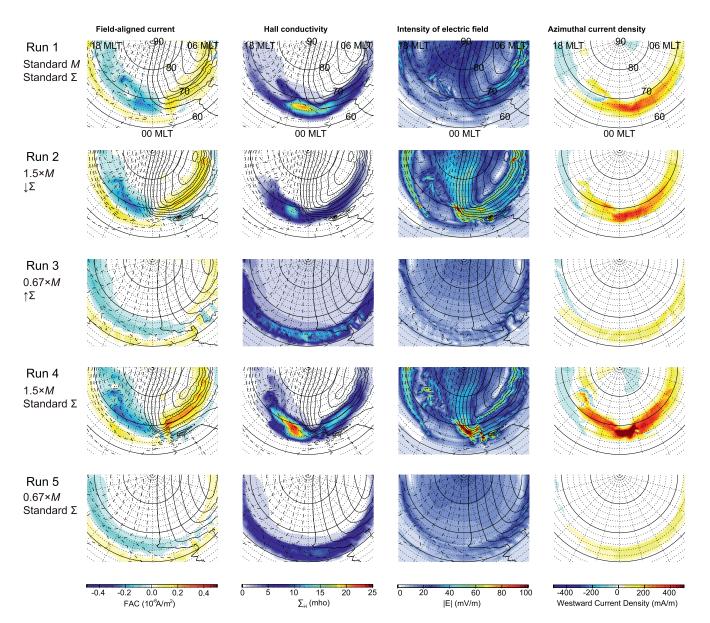


Figure 4. From left to right, field-aligned current (negative upward), nondiagonal component of the ionospheric conductivity (Hall conductivity), intensity of the ionospheric electric field, and azimuthal component of the ionospheric current (positive westward) at 5 min after the expansion onset. The Sun is to the top.

electric field, and the azimuthal component of the ionospheric current at 5 min after the expansion onset. By looking at the results for Runs 1 and 4, one can find that all the quantities (the FAC, the ionospheric conductivity, the ionospheric electric field, and the ionospheric current) increase with increasing M. The large increase in the ionospheric current results in the rapid development of the auroral electrojet as shown in Figure 1. When we introduce the ionospheric conductivity model depending on M ($\Sigma_P \propto M^{-1}$ and $\Sigma_H \propto M^{-4/3}$), the auroral electrojet decreases with increasing M (from Run 4 to Run 2). One simple explanation for the decrease in the auroral electrojet is the decreases in the conductivity and the electric field. (From Run 4 to Run 2, the electric field increases in the polar cap, but it does not directly affect the intensity of the auroral electrojet.) Another explanation, which is probably more essential, is that the auroral electrojet is largely regulated by the FAC. To satisfy the current continuity, when the FAC decreases, the ionospheric current must decrease. However, the situation is not so simple because the FAC, the ionospheric conductivity; the ionospheric electric field, and the ionospheric current are coupled with each other. So far, the reason why the FAC depends on the ionospheric conductivity is still unknown. This is an important subject to be investigated in the future.



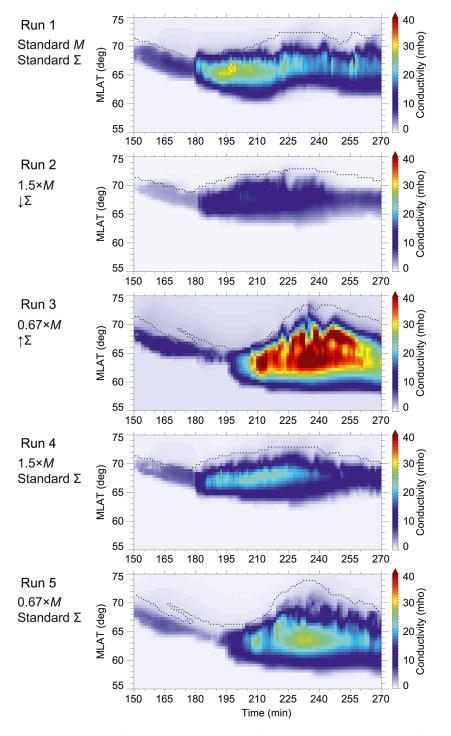


Figure 5. Nondiagonal component of the ionospheric conductivity (which can be regarded as a proxy of auroral brightness) at midnight as a function of magnetic latitude (MLAT) and time. The dashed line indicates the open-closed boundary of the magnetic field line.

Figure 5 shows the nondiagonal component of the ionospheric conductivity (the Hall conductivity) as a function of MLAT and time at magnetic midnight. Hereinafter, we regard the ionospheric conductivity as aurora for the sake of simplicity. The observations show that the auroral brightness is correlated with the ionospheric conductivity (Kamide et al., 1986; Robinson et al., 1989). Noticeable points can be summarized below. During the growth phase, the poleward boundary of the auroral oval and the open-closed boundary

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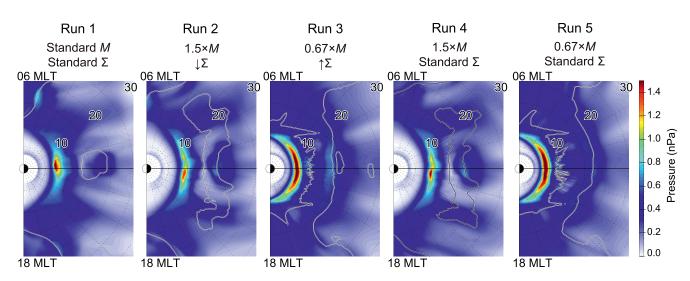


Figure 6. Plasma pressure in the equatorial plane. The black contour indicates that the Z-component of the magnetic field is zero. The Sun is to the left.

move equatorward. Just before the expansion onset, the open-closed boundaries are located at ~68, ~69, ~66, ~68, and ~66 MLAT for Runs 1, 2, 3, 4, and 5, respectively. This means that the open-closed boundary shifts poleward with increasing *M*. The dependence of the open-closed boundary on *M* is consistent with the previous simulation study (Cnossen et al., 2012). The equatorward boundary of the auroral oval also moves equatorward during the growth phase. Just before the expansion onset, the boundaries are located at ~64, ~66, ~63, ~66, and ~63 MLAT for Runs 1, 2, 3, 4, and 5, respectively. The equatorward boundary shifts poleward by ~2° when *M* is increased by 1.5, and shifts equatorward by ~1° when *M* is decreased by a factor of 1.5. During the expansion phase, the auroral oval expands less poleward and equatorward for stronger *M*. The poleward boundary of the auroral oval does not always coincide with the polar cap boundary during substorm expansion (Birn & Hesse, 2013; Ebihara & Tanaka, 2015a, 2017).

Figure 6 shows the plasma pressure in the equatorial plane at the expansion onset. By the expansion onset, the plasma pressure has been well developed on the nightside as previously pointed out on the basis of observations (Yao et al., 2015a, 2015b) and simulations (Ebihara & Tanaka, 2015b; Tanaka et al., 2010, 2017). The enhancement of the plasma pressure is associated with compression that happens when the earthward flowing plasma is decelerated (Birn & Hesse, 2005; Tanaka et al., 2010, 2017). The deceleration is primarily caused by the pressure force directing outward. The peak of the plasma pressure takes place at $X \cong -7$ R_E (Run 1), -8 R_E (Runs 2 and 4), and -6 R_E (Runs 3 and 5). The result indicates that the peak of the plasma pressure takes place at greater distance for larger *M* because the outward pressure force increases with increasing *M*. The force balance is highly variable near the substorm expansion onset as suggested by Ebihara and Tanaka (2013) but is not shown here. The earthward flow results from the formation of a near-Earth neutral line, which can be identified from the earthward part of the black contour indicating that the *Z*-component of the magnetic field is zero. They are located at $X \cong -13$ R_E (Runs 1, 2, and 4), and $\cong -9$ R_E (Runs 3 and 5). The near-Earth neutral line forms at greater distance for larger *M*, but the reason is not known.

In the ionosphere, the Joule dissipation rate increases significantly in a latitudinally confined region during the substorm expansion (Ahn et al., 1983; Kamide et al., 1996; Lu et al., 2016). To supply the magnetic energy into such confined region, the energy transfer should be associated with FACs, and the magnetic energy must be generated somewhere in the field line connecting to the auroral region. The global MHD simulation results show that the near-Earth dynamo is the one that is responsible to supply the magnetic energy into the auroral region during the substorm expansion (Birn & Hesse, 2005; Ebihara & Tanaka, 2015a, 2015b; Tanaka, 2015). Figure 7 summarizes perspective views of the magnetosphere at the expansion onset. The light and dark greenish surfaces indicate the regions where $\mathbf{J} \cdot \mathbf{E} = -2 \times 10^{-12}$ and -4×10^{-12} W m⁻³, respectively. This region is called a near-Earth dynamo and is responsible for the generation of the substorm-associated FACs. To evaluate the near-Earth dynamo, we calculated the volume integral of $\mathbf{J} \cdot \mathbf{E}$ for the region where $\mathbf{J} \cdot \mathbf{E} < 0$ in the rectangular region ranging from X = -15 R_E to 0, $Y = -10_E$ to 10 R_E,



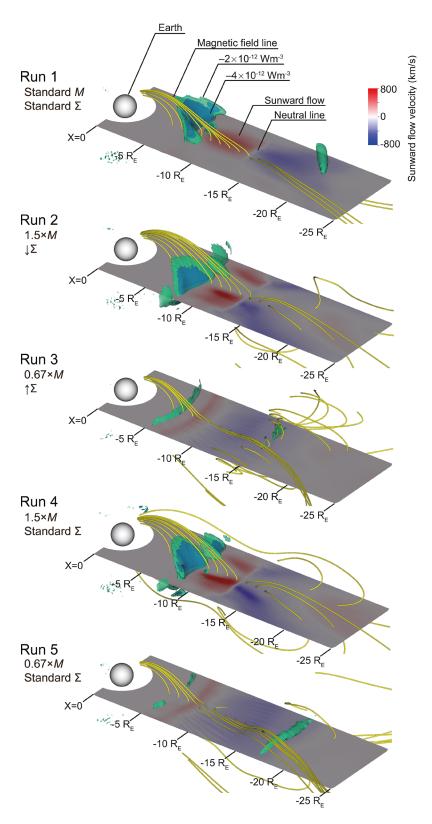


Figure 7. Perspective view of the magnetosphere as if one is looking down the Earth from the dusk-midnight sector above the equatorial plane. All the views were obtained at the expansion onset. The yellow lines indicate the magnetic field lines. The light and dark greenish surfaces indicate the $\mathbf{J} \cdot \mathbf{E}$ values of -2×10^{-12} and -4×10^{-12} W m⁻³, respectively. The sunward flow velocity is indicated by color in the equatorial plane (at Z = 0). The white sphere indicates the Earth.

and $Z = -5 R_E$ to $5 R_E$. The results are summarized in Table 2. The generation rate of the magnetic energy increases with increasing *M*. This gives rise to the rapid development of the net FAC and the auroral electrojet for larger *M*-values. The sunward flow of plasma is indicated by the red color in the equatorial plane. The maximum speed of the sunward flow increases with increasing *M* as summarized in Table 2. For larger *M*-values, the relatively high-speed flow can enhance the plasma pressure (thermal energy) more efficiently because $\partial P/\partial t = -\mathbf{V} \cdot \nabla P - \gamma P \nabla \cdot \mathbf{V}$, where *P*, **V**, and γ are the plasma pressure, the bulk velocity, and the ratio of specific heat (=5/3), respectively (Birn & Hesse, 2005).

We calculated the incident magnetic energy into the magnetosphere by integrating the Poynting flux over the surface of the magnetosphere at the expansion onset. The methodology for the calculation is described by Ebihara et al. (2019) in detail. The results are summarized in Table 2. The incident magnetic energy increases with increasing M probably due to the inflation of the magnetosphere for larger M-value. Table 3 summarizes the net FAC, the Joule heating rate in the ionosphere, the generation rate of magnetic energy in the near-Earth dynamo, and the incident magnetic energy into the magnetosphere at the moment when the net FAC reaches maximum. A noticeable feature is that the maximum value of the net FAC decreases with increasing M, whereas the incident magnetic energy into the magnetosphere increases with increasing M. This implies that the efficiency of the generation of the substorm-associated FAC increases with decreasing M. The ratios of the Joule heating rate in the ionosphere to the incident magnetic energy into the magnetic tosphere are 2.3%, 1.9%, and 3.0% for Runs 1, 2, and 3, respectively, meaning that the ratio increases with decreasing M. For lower M-values, the incident energy into the magnetosphere is efficiently transported to the ionosphere.

4. Discussion

It will take ~830 years to decrease *M* by 50% if the rate of the change in *M* is constant. Thus, the results presented above cannot be verified by observations until the very distant future. We repeated the same simulation except that *M* increases by ~3% (Run 2') and *M* decreases by ~3% (Run 3'). The 6% decrease in *M* corresponds to the recent change in Earth's *M* over ~100 years. The calculated SMU/SML index is shown in Figure S1. The difference is very small. One cannot easily identify the contribution from the secular variation of *M* to the SMU/SML index on the basis of data acquired over ~100 years. For this reason, we have considered larger changes in *M* to more clearly understand and characterize the changes in the development of substorms in response to varying the dipole moment.

The change in M has significant consequences on the development of the auroral electrojet. For the lower M-value, (1) the expansion onset takes place later, (2) the auroral electrojet develops slowly, and (3) the maximum auroral electrojet increases. These consequences are closely related to the FACs that develops during the expansion phase. The first two consequences are probably associated with the slow magneto-spheric convection as manifested by the polar cap potential (i.e., Table 2). The relatively low speed of the sunward flow in the near-Earth plasma sheet (i.e., Table 2) may also cause the slow evolution of the auroral electrojet. In the near-Earth dynamo, the generation rate of the magnetic energy increases with increasing M, which may result from the relatively high speed of the sunward flow in the near-Earth plasma sheet (i.e., Table 2). The pattern of the sunward flow depends on M (i.e., Figure 7), but we have no reasonable explanation for that dependency. The third consequence can be understood to the nonlinear dependence of the substorm-associated FACs on the ionospheric conductivity. The magnetosphere–ionosphere coupling processes is essentially nonlinear, and this nonlinear process is beyond the scope of this paper. We will investigate the influence of the ionospheric conductivity on the whole magnetospheric system in the future.

Regarding the third consequence, the simulation result opposes to the theoretical prediction (Glassmeier et al., 2004) that the magnitude of the auroral electrojet increases with increasing M. Glassmeier et al. (2004) considered that the magnitude of the auroral electrojet depends on the ionospheric conductivity and the electric field. However, our simulation result shows that the FACs also depend on M, and that the maximum value of the net FAC increases with decreasing M. The dependence of the FACs on the conductivity is also shown to be significant.

The dipole magnetic moment of 11.64×10^{22} A m², which is used for Runs 2 and 4, roughly corresponds to that about 2650 years ago, during which the highest dipole moment occurred throughout the Holocene pe-



riod (Knudsen et al., 2008). Coincidently, the earliest records of candidate aurorae are dated back to similar age (Hayakawa, Mitsuma, et al., 2019). The Assyrian cuneiform tablets (~679-655 BCE) described situations of "a red glow [keeps blazing] at [the zenith]," "a red cloud keeps being placed in the sky," and "red covers the sky." These statements are most likely interpreted as reddish aurorae that covered the sky over at Nineveh or Babylon at which the geomagnetic latitude was ~35-41°. With an empirical relationship between the equatorward boundary of the auroral oval and the Dst index (Yokoyama et al., 1998), Hayakawa, Mitsuma, et al. (2019) suggested that large magnetic storms occurred during these occasions. The scale of the storms is estimated to be comparable to that occurred in 1909 (minimum Dst of -595 nT) (Hayakawa, Ebihara, et al., 2019). This empirical relationship was obtained on the basis of data acquired between 1983 and 1991 during which M ranges from 7.88×10^{22} to 7.83×10^{22} A m², according to the international geomagnetic reference field (Thébault et al., 2015). Figure 5 suggests that the equatorward boundary of the auroral oval would shift poleward by $\sim 2^{\circ}$ about 2650 years ago than present for the same solar wind condition. Stronger solar wind driver is required to generate the auroral oval that is located at the same MLAT as today. For Run 2, the equatorward boundary of the auroral oval is located at \sim 66 MLAT just before the expansion onset. We repeated Run 2 except that we increased the solar wind speed from 400 to 500 km/s (data not shown). The equatorward boundary is located at \sim 64 MLAT, which approximately corresponds to that for Run 1. This speculation would be applicable for all the estimation of the solar wind conditions based on historical records back for the last 2650 years from now because the highest dipole moment occurred at that time throughout the Holocene period (Knudsen et al., 2008). From that sense, the poleward shift of $\sim 2^{\circ}$ would be the upper limit throughout the Holocene period for the same solar wind condition.

For possible impacts on space weather, the following prediction can be drawn. If *M* keeps decreasing in future, the maximum intensity of the auroral electrojet will increase. The increased auroral electrojet may increase geomagnetically induced currents that are hazardous to power grid (Boteler et al., 1998; Pirjola & Viljanen, 1989). Once the substorm occurs, the auroral oval is expected to expand more poleward and equatorward. This means that the harmful area will also expand both to the poleward and equatorward. If *M* decreases, hot electrons originating in the nightside plasma sheet will penetrate deeper during the substorm expansion as shown in Figure 7. The hot electrons result in static charge on the surface of satellite, and large discharges and arc (Garrett, 1981; Rosen, 1976). Overall, the decrease in the value of Earth's dipole moment may have a greater impact on human activities than present.

5. Conclusions

We performed the global MHD simulation with different dipole magnetic moment *M*. We increased and decreased *M* by a factor of 1.5 and changed the ionospheric conductivity in accordance with *M*. When we imposed southward IMF, auroral substorms took place regardless of *M*. However, the evolution of the auroral substorm shows a dependence on *M*. The following conclusions can be drawn.

- 1. For the lower *M*-value, the expansion onset takes place later, the auroral electrojet develops slowly, and the maximum auroral electrojet increases. The first two consequences are probably associated with the slow magnetospheric convection as manifested by the polar cap potential. The third consequence can be understood to the nonlinear dependence of the substorm-associated FACs on the ionospheric conductivity.
- 2. For the lower *M*-value, the amount of incident magnetic energy into the magnetosphere decreases, whereas the maximum auroral electrojet increases. This implies that the efficiency of the generation of the substorm-associated FACs increases with decreasing *M*.
- 3. For the lower *M*-value, the auroral oval shifts equatorward during the growth phase and expands more equatorward and poleward during the expansion phase.

The results presented here are not definitive because we omitted the kinetic processes related to particle precipitation, we excluded the contribution from the quadrupole component, and we used simple relations between the ionospheric conductivity and *M*. Careful investigations are needed to fully understand the future substorms.



Data Availability Statement

The simulation data used in this study are available at http://doi.org/10.5281/zenodo.4003656.

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